

# Separating the Coupled Climate Response to Sea Ice Loss and Greenhouse Warming

Paul Kushner<sup>1</sup>, Russell Blackport<sup>2</sup>, Kelly McCusker<sup>3</sup>, Stephanie Hay<sup>1</sup>, Thomas Oudar<sup>1</sup>, John Fyfe<sup>4</sup>, Cecilia Bitz<sup>3</sup>

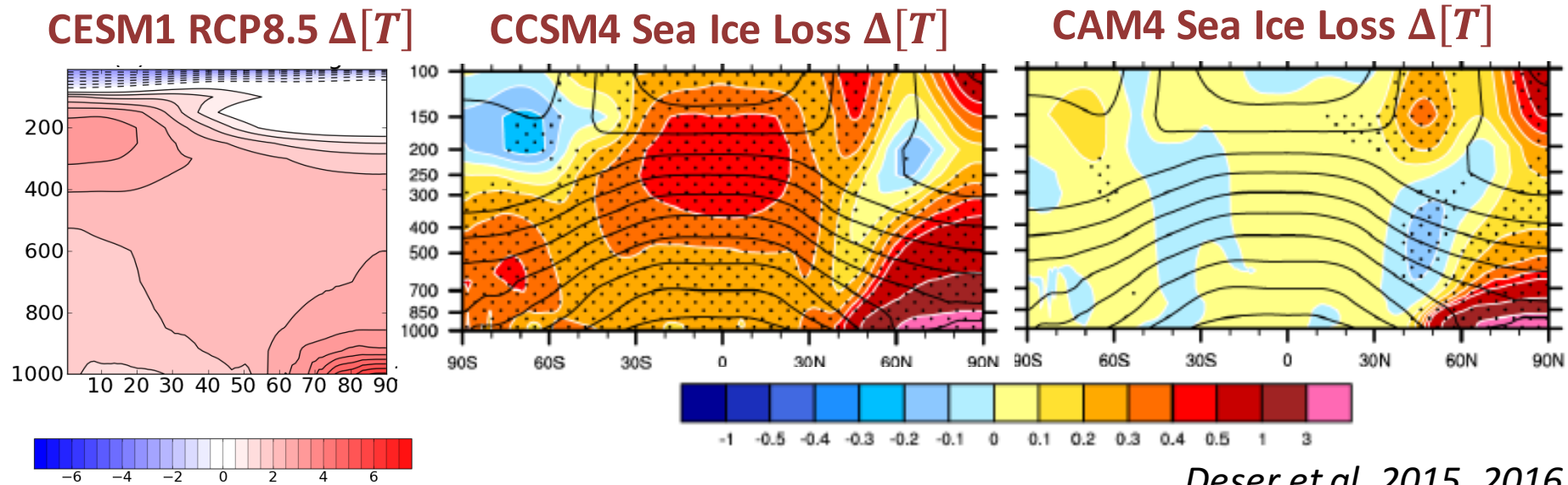
1: University of Toronto

2: University of Exeter

3: University of Washington

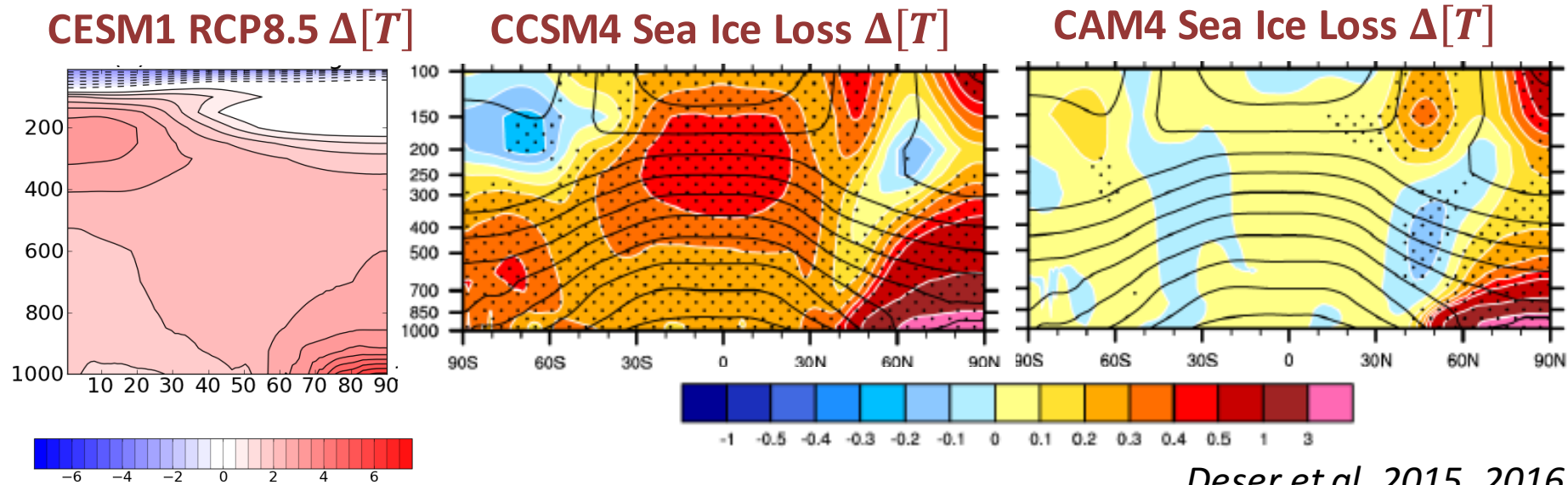
4: Environment and Climate Change Canada

# Can We Isolate the Impact of Sea Ice Loss on Circulation?



*Deser et al. 2015, 2016*

# Can We Isolate the Impact of Sea Ice Loss on Circulation?

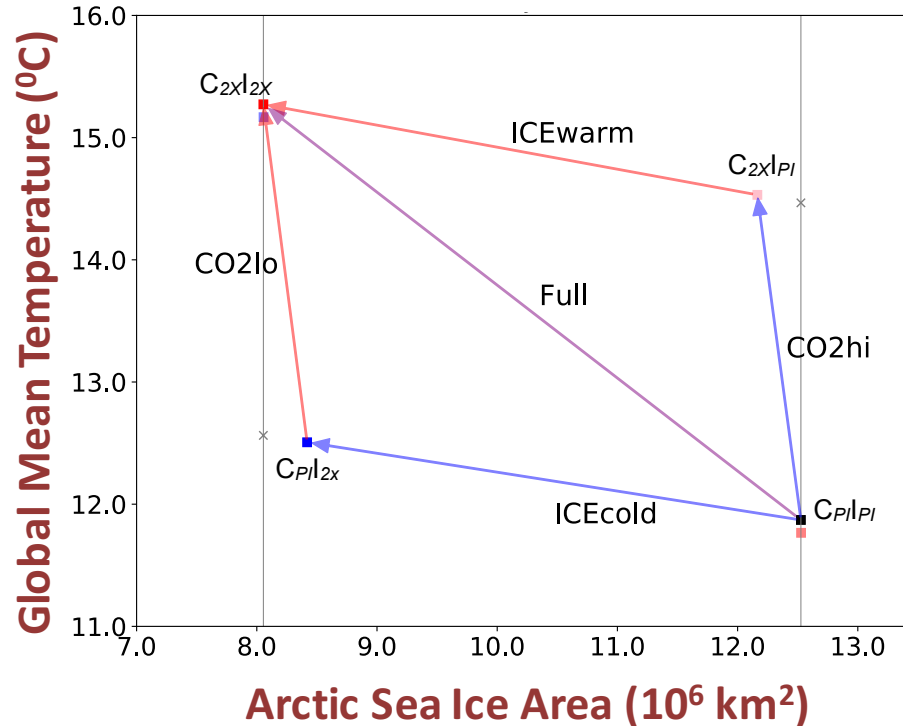


In very well sampled coupled climate simulations:

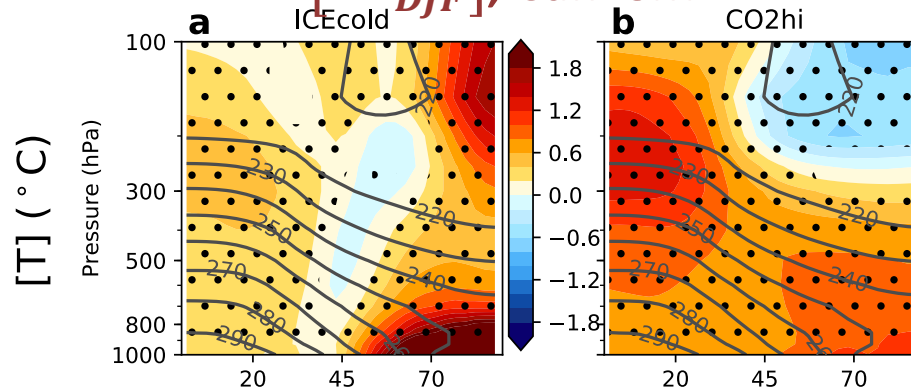
- The greenhouse warming response is remarkably separable from the sea ice loss response.
- *Multivariate pattern scaling* helps divide the response between sea ice loss and low latitude warming.
- The response to sea ice loss is more robust than the low latitude warming response.

# Paired Sea Ice and CO2 Perturbation Experiments

## DJF Response, CanESM2



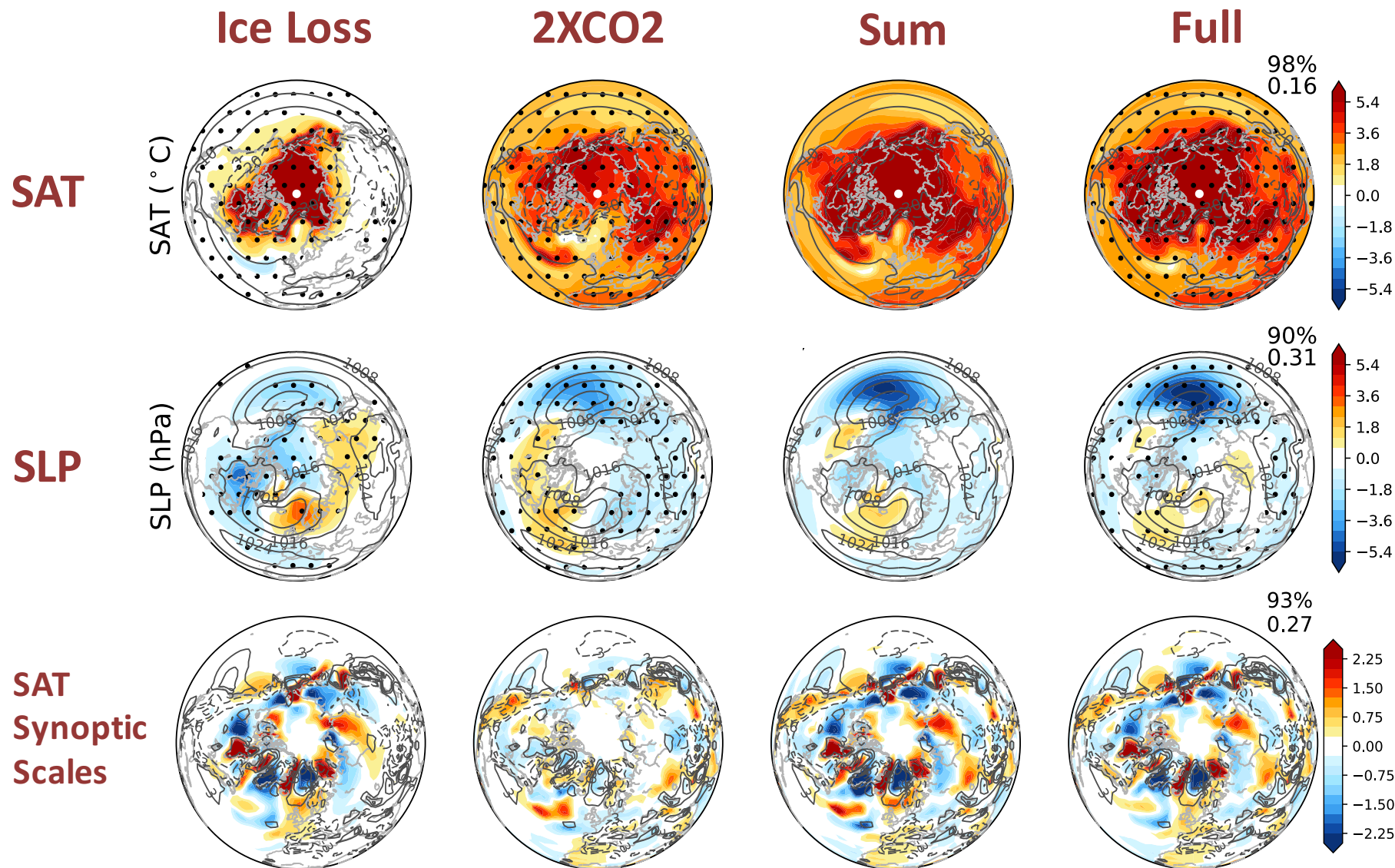
## $[\Delta T_{DJF}]$ , CanESM2

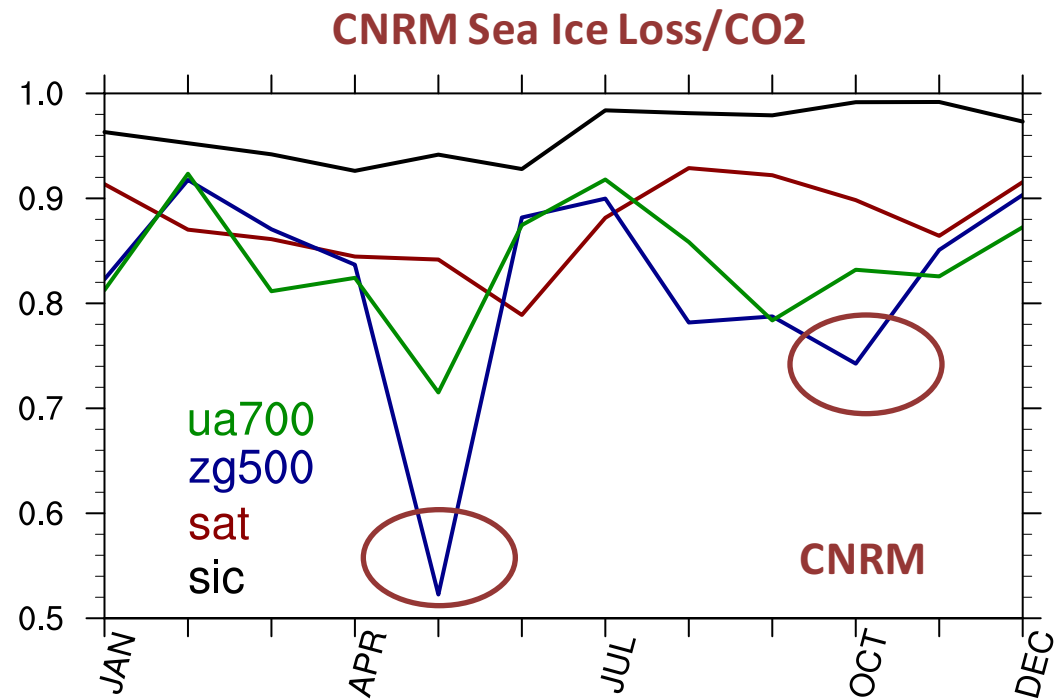
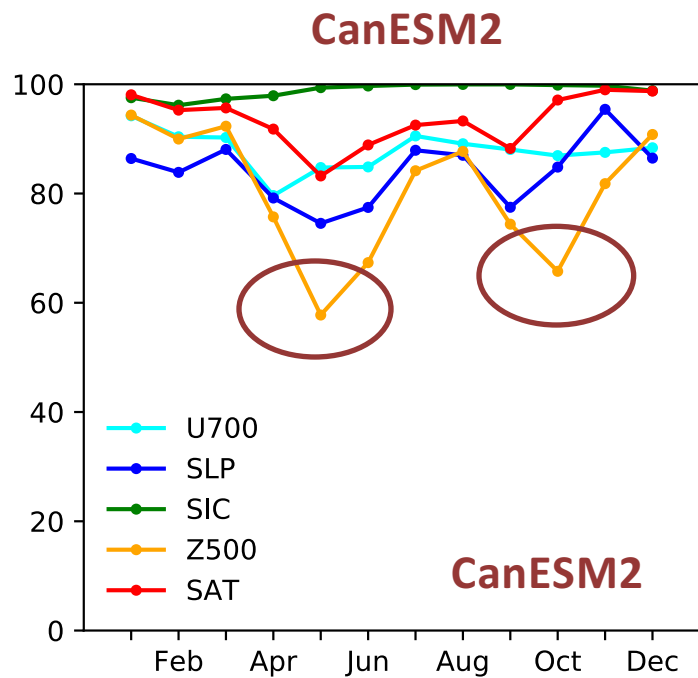


*McCusker et al.  
submitted,  
building on  
Oudar et al.  
2017*



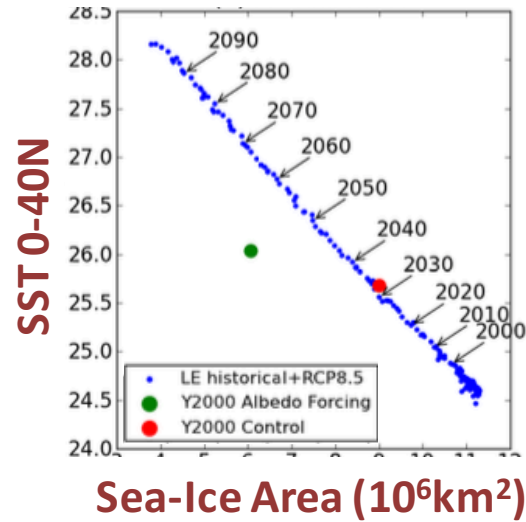
# Additivity of Sea Ice Loss and CO2



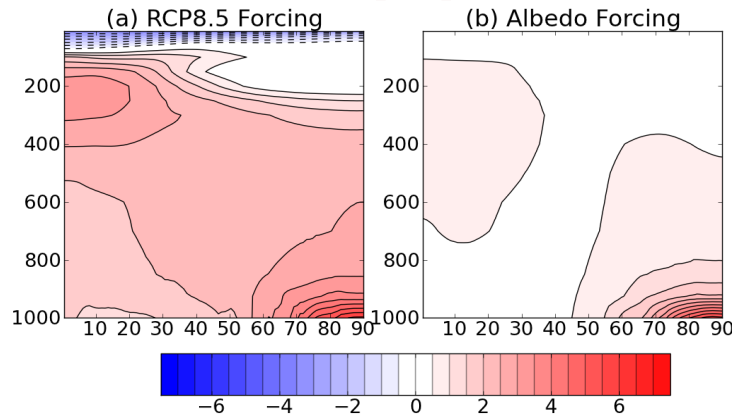


# Sea Ice Loss and Projected Climate Change

## ANN Response, CESM1



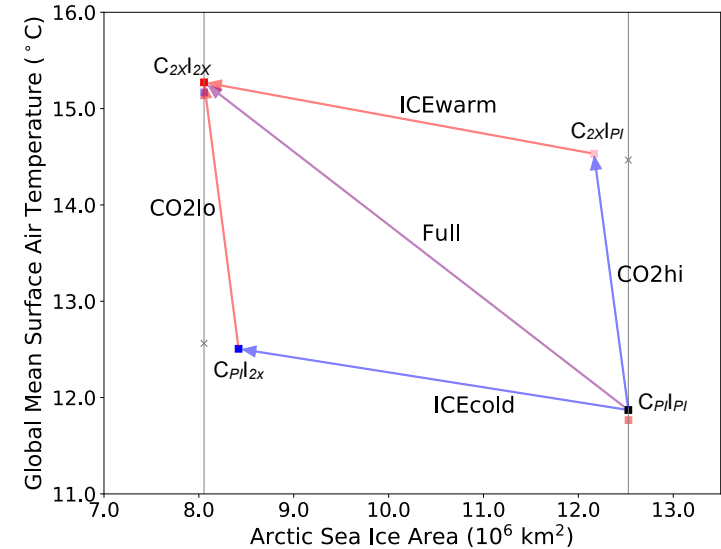
## ANN $[\Delta T]$



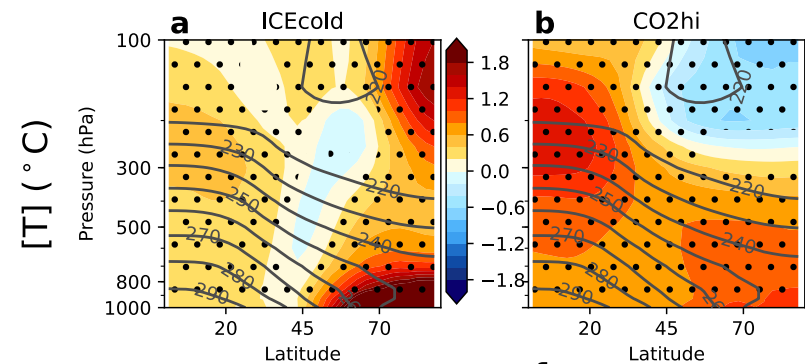
Blackport and Kushner 2017

# Paired Sea Ice and CO2 Perturbation Experiments

## DJF Response, CanESM2



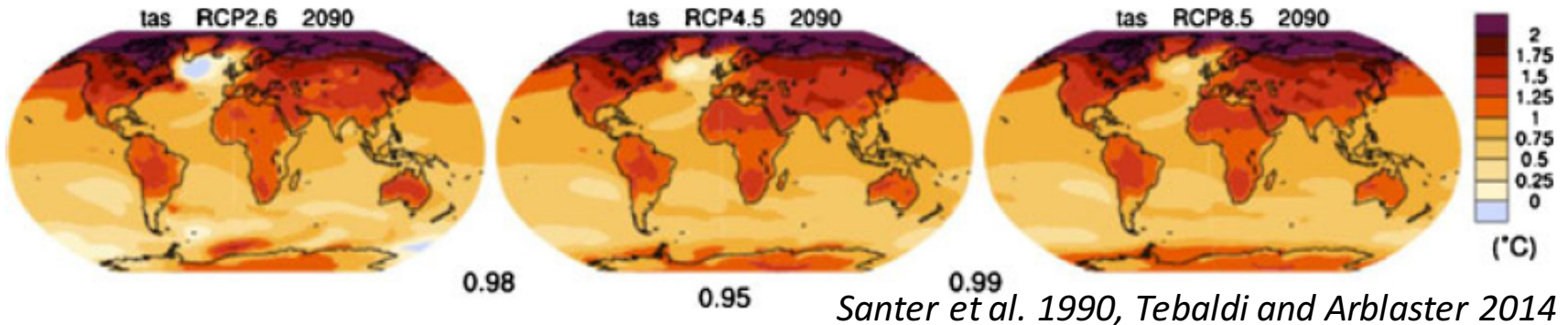
## ANN $[\Delta T]$



McCusker et al. 2017

# Pattern Scaling on Temperature

## ANN Surface Temperature Response Per Global mean temperature, CMIP5

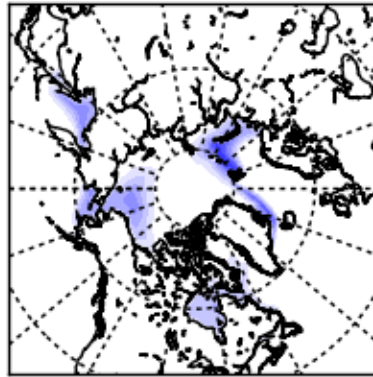


For small change in global mean temperature  $\delta T_g$ ,

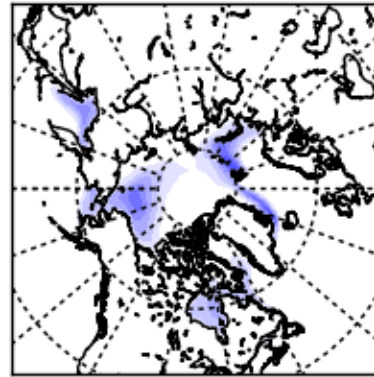
$$\delta Z = Z(T_g + \delta T_g) - Z(T_g) \approx \frac{\partial Z}{\partial T_g} \delta T_g$$

# Can we Pattern Scale on Sea Ice?

NCAR  
CESM1

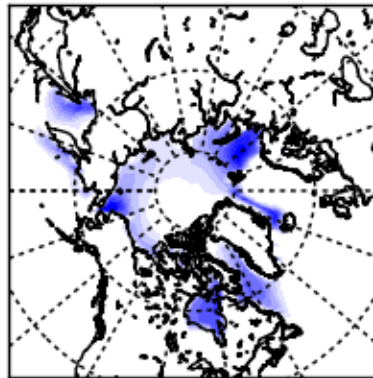


(a) RCP8.5 Forcing

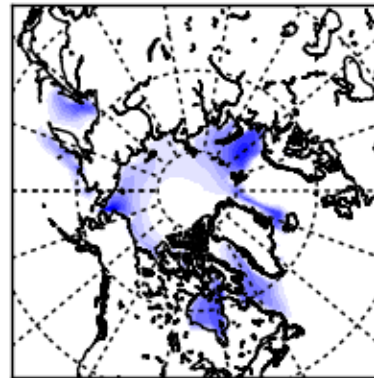


(b) Albedo Forcing

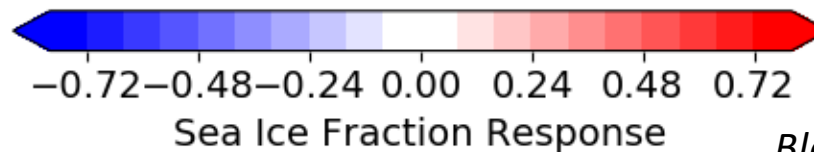
CanESM2



(d) Full



(e) ICEcold



*Blackport and Kushner 2017;  
Hay et al. in prep.*

# Pattern Scaling on Multiple Variables

- Assume field  $Z$  depends parameterically on *internal* integrated variables  $X_i$  (e.g. surface temperature, sea ice extent, ...).
- Want to know sensitivities of  $Z$  to separate change in each  $X_i$ :

$$\left. \partial Z / \partial X_i \right|_{X_j}$$

- Assume coupled simulations generate responses  $\delta Z_m$  with distinctive responses  $\delta X_i$ :

$$\delta Z_m = \sum_i \left. \partial Z / \partial X_i \right|_{X_j} \delta X_{i,m}$$

- We infer sensitivities by inverting this relationship.



# Pattern Scaling on low-latitude temperature $T_l$ and sea ice $I$

$$Z = Z(T_l, I)$$

$$\delta Z = Z(T_l + \delta T_l, I + \delta I) - Z(T_l, I)$$

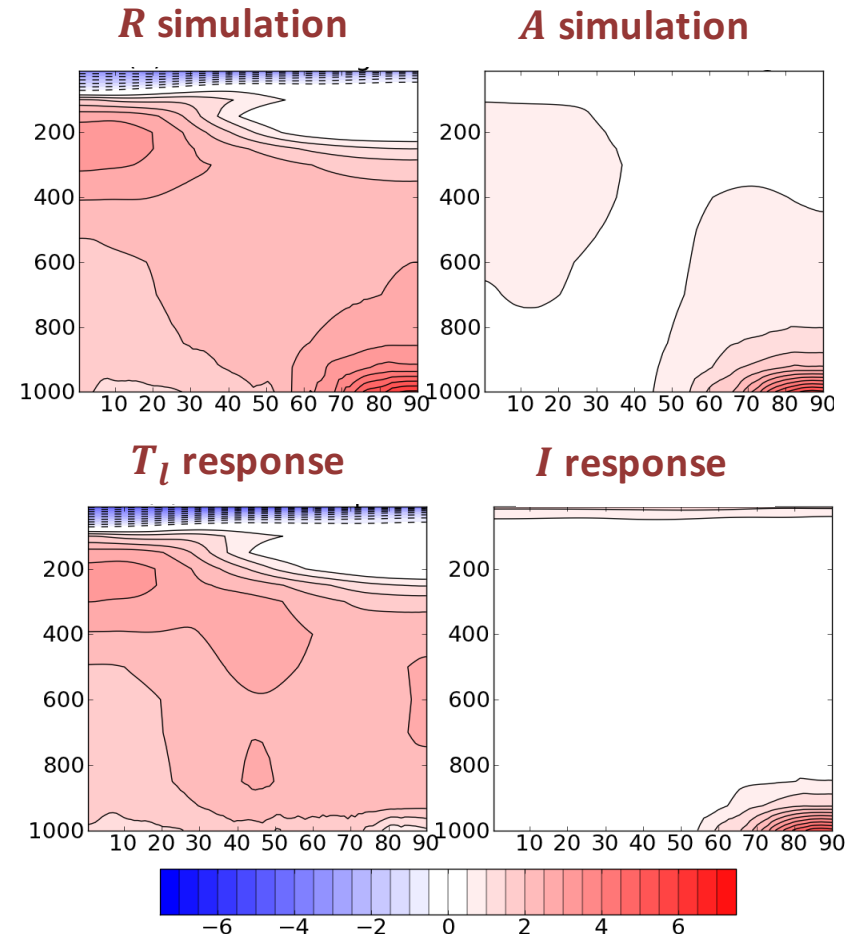
$$\begin{pmatrix} \left. \frac{\partial Z}{\partial T_l} \right|_I \\ \left. \frac{\partial Z}{\partial I} \right|_{T_l} \end{pmatrix} = \begin{pmatrix} \delta T_{l,R} & \delta I_R \\ \delta T_{l,A} & \delta I_A \end{pmatrix}^{-1} \begin{pmatrix} \delta Z_R \\ \delta Z_A \end{pmatrix}$$

Sensitivity to  $T_l$  or  $I$  changes,  
with the other held fixed.

$\delta T_l, \delta I, \delta Z$

- From coupled model simulations  $R$  and  $A$ .

CESM1 [ $\Delta T$ ], ANN



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$$Z = Z(T_l, I)$$

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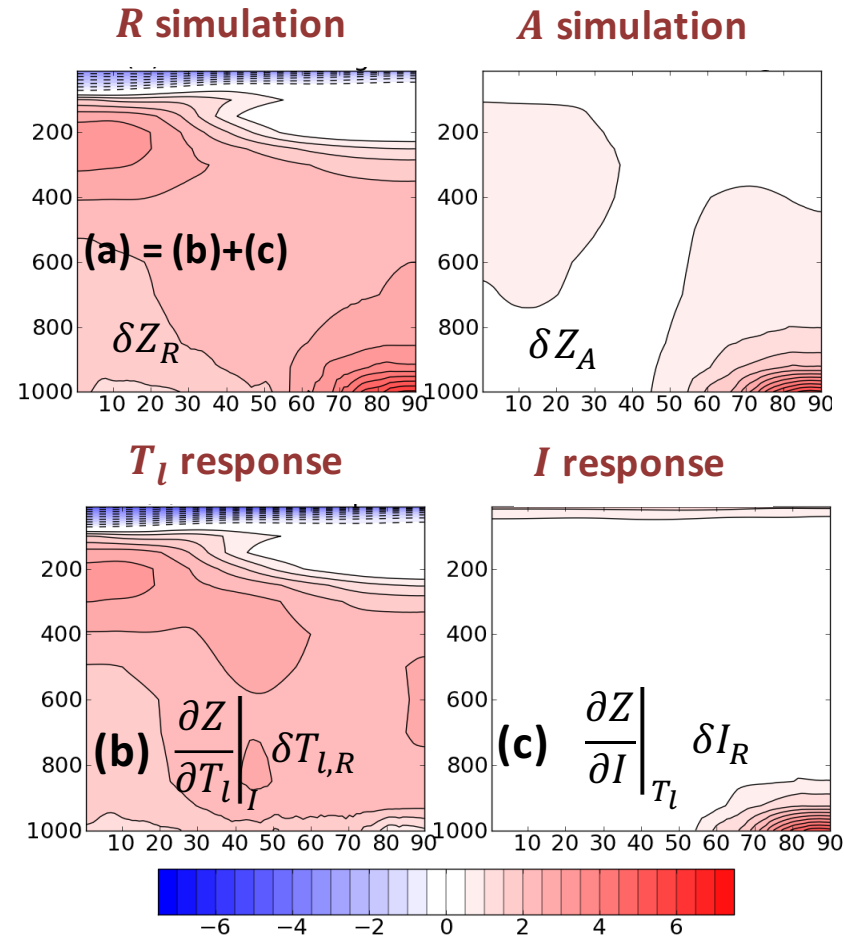
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CESM1 [ $\Delta T$ ], ANN

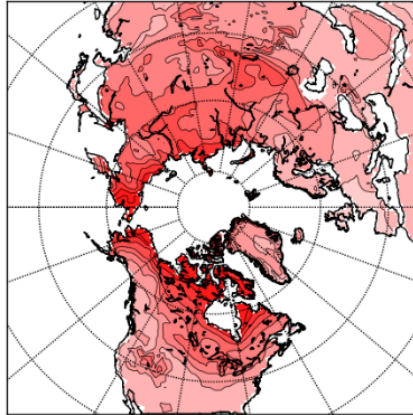




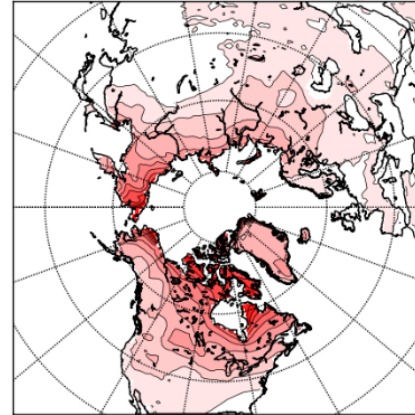
# CESM1 $\Delta T_{2m,DJF}$

Coupled Model  
Experiments

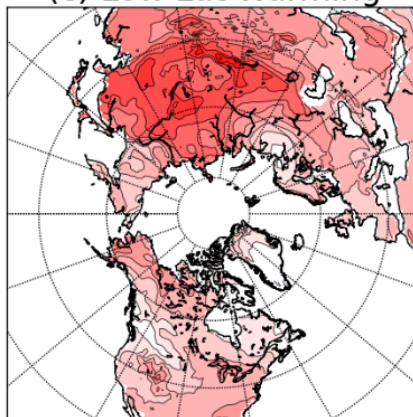
*R* simulation



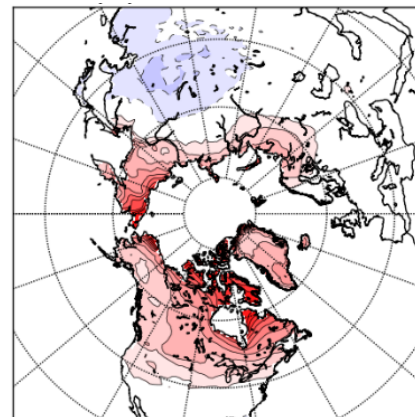
*A* simulation



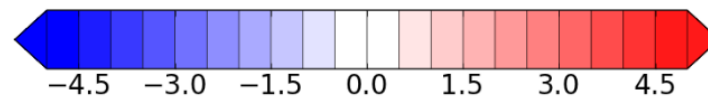
$T_l$  response



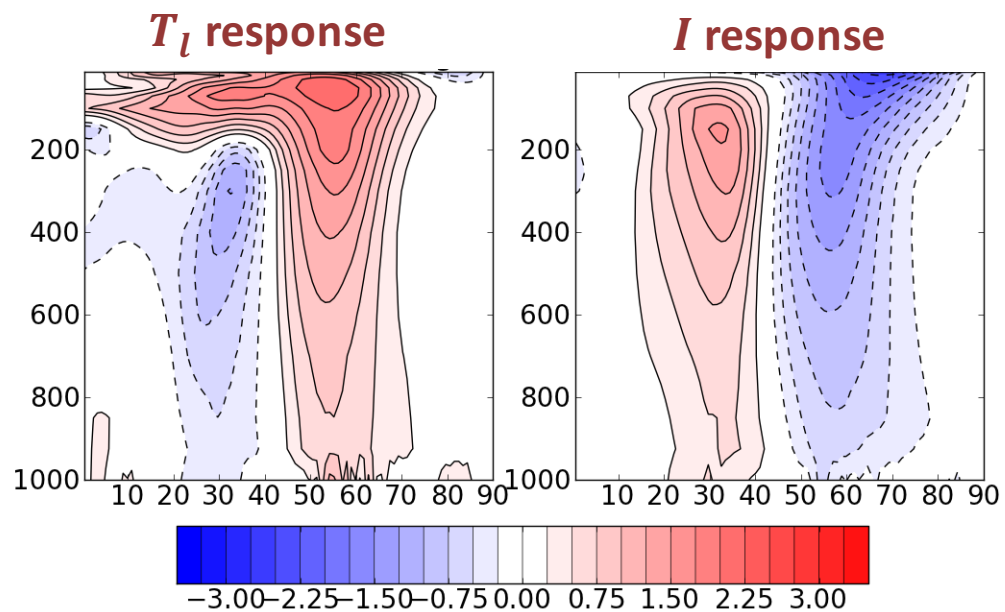
*I* response



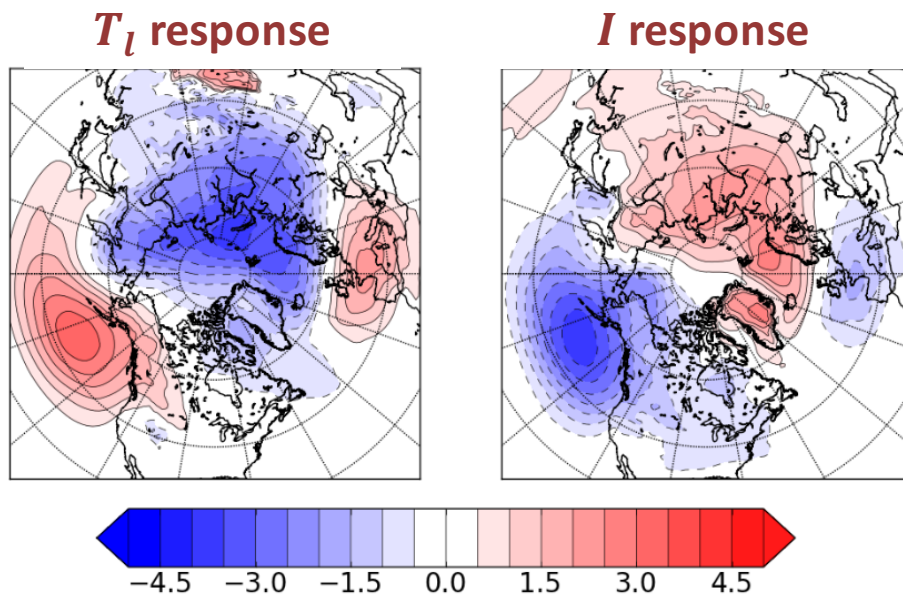
Diagnosed  
Decomposition



CESM1  $\Delta[U_{\text{DJF}}]$



CESM1  $\Delta SLP_{\text{DJF}}$



# $\Delta SLP_{DJF}$ (hPa)

$T_I$  response

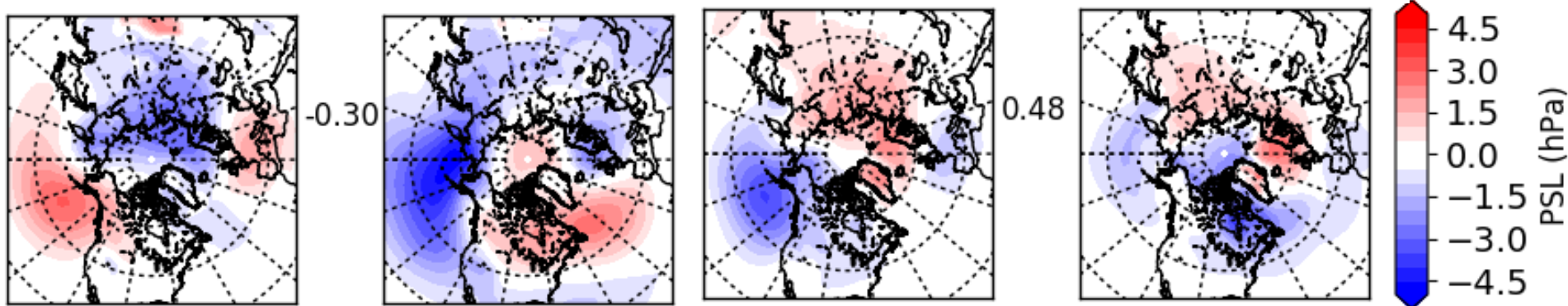
$I$  response

CESM1

CanESM2

CESM1

CanESM2

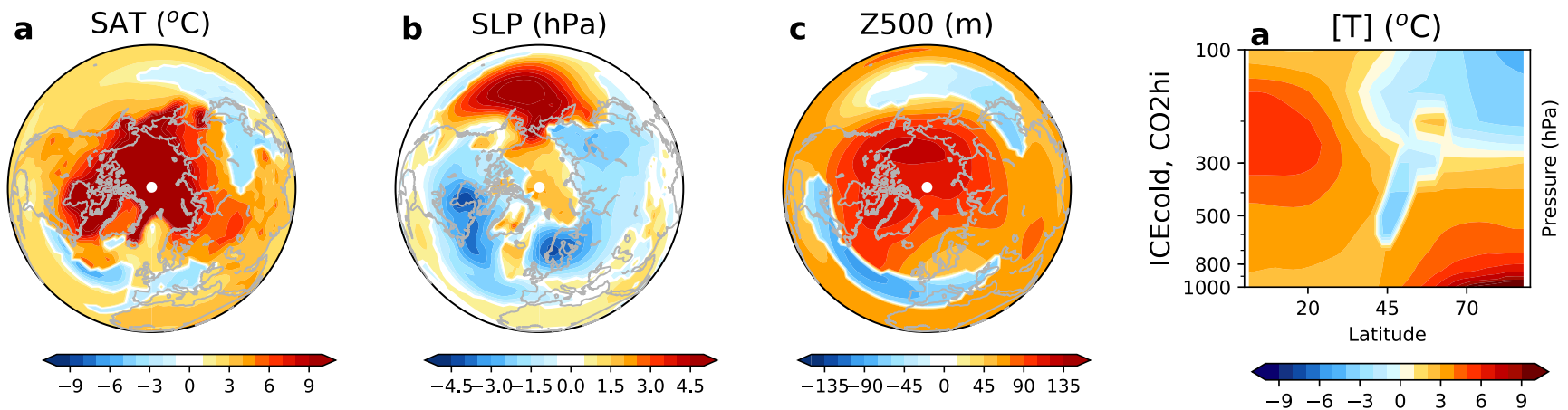


- Strengthened Aleutian Low and Eurasian high seen in several papers: Sun et al. 2015, Deser et al. 2016, Blackport and Kushner 2017, Oudar et al. 2017
- Robust sea ice loss response seen in several other fields: surface temperature, u850, precipitation
- Response to low latitude warming is less robust except in summer.

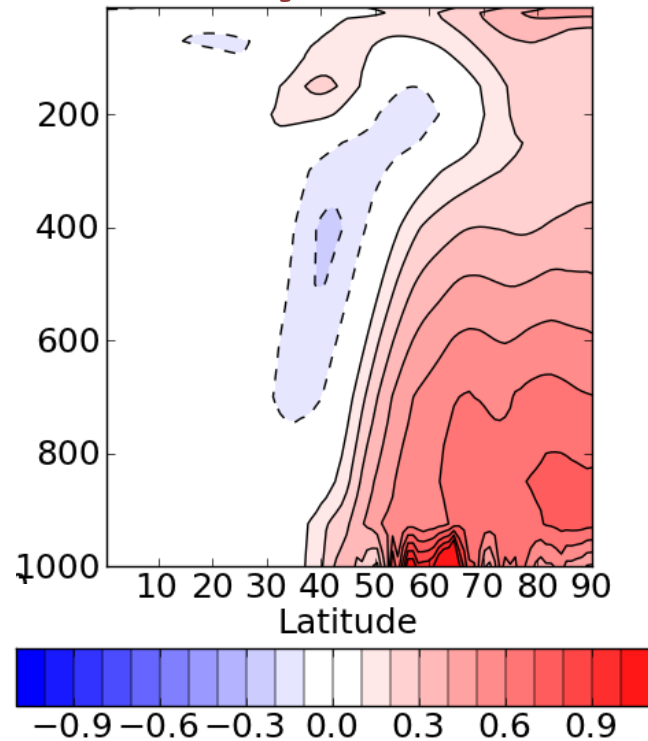
- **Key point #1:** AOGCM simulations of sea ice loss are worth carrying out despite the computational cost, and we can use pattern scaling for further insight.
- **Key point #2:** Sea ice loss and radiative forcing responses are remarkably separable, enabling attribution.
- **Key point #3:** The response to sea ice loss is more robust across models than the response to low latitude warming, especially in the North Pacific.
- **Topics of current interest: a few slides.**

# Diagnosing Feedbacks in CanESM2

$$\text{“Agreement” } A = (|I| + |C|)\text{sign}(I)\text{sign}(C)$$

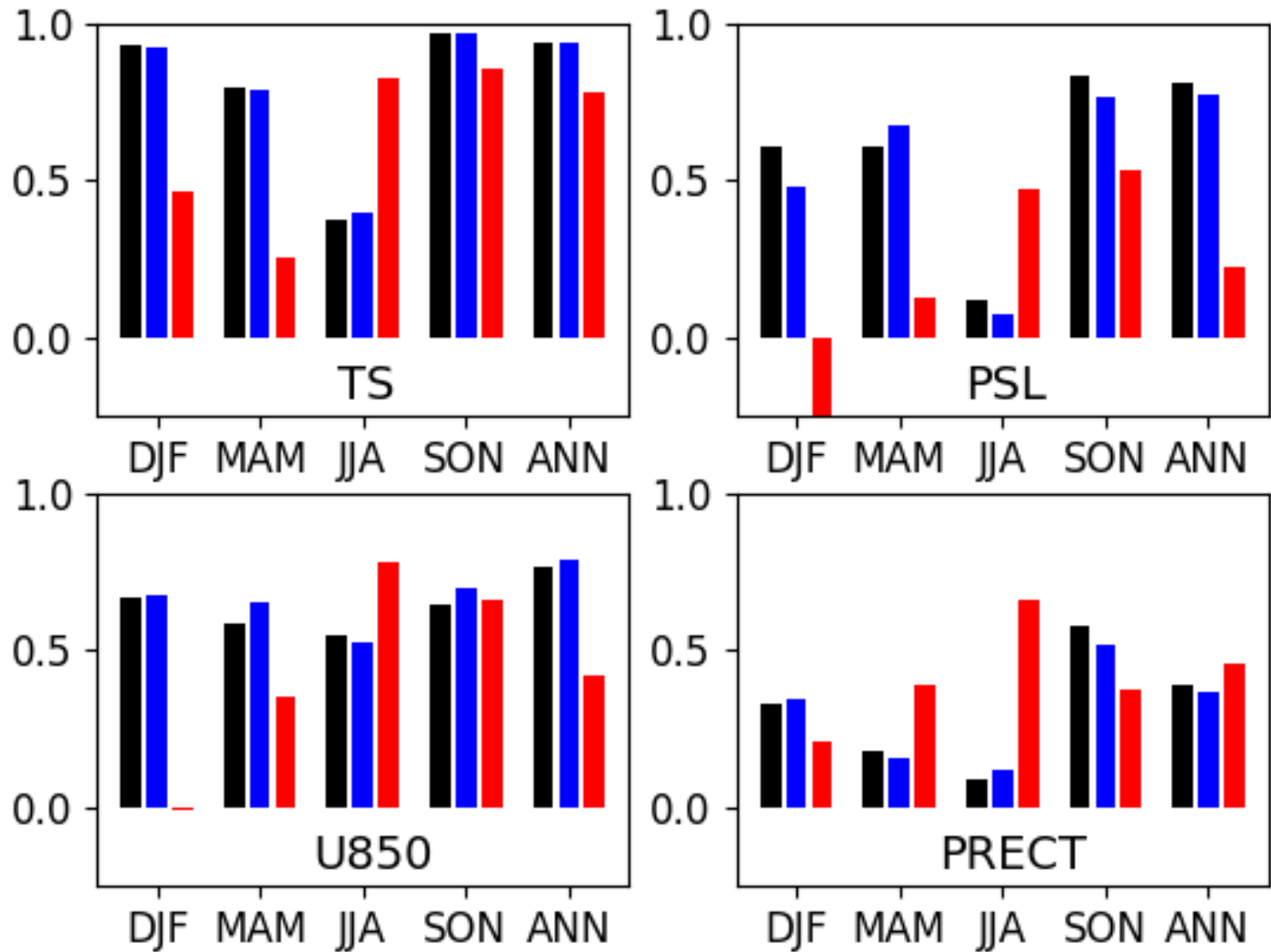


## $[\Delta T_{\text{DJF}}]$ from SST warming caused by sea ice loss



# Are CESM1 and CanESM2 consistent?

Spatial pattern correlation between CESM1 and CanESM2 responses



CESM Sea Ice Albedo and CanESM2 Sea Ice Loss ( $A$  simulations)



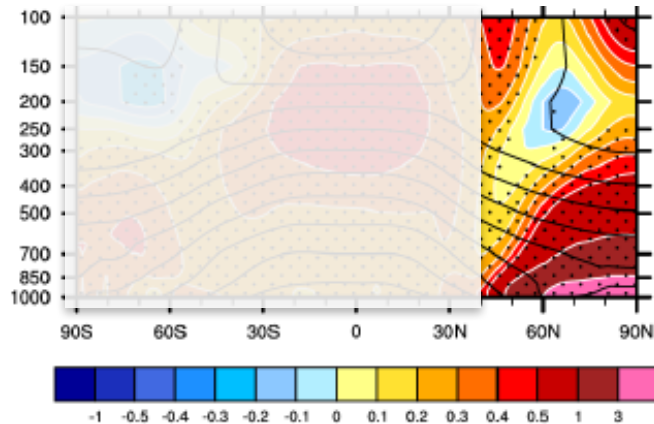
$I$  response



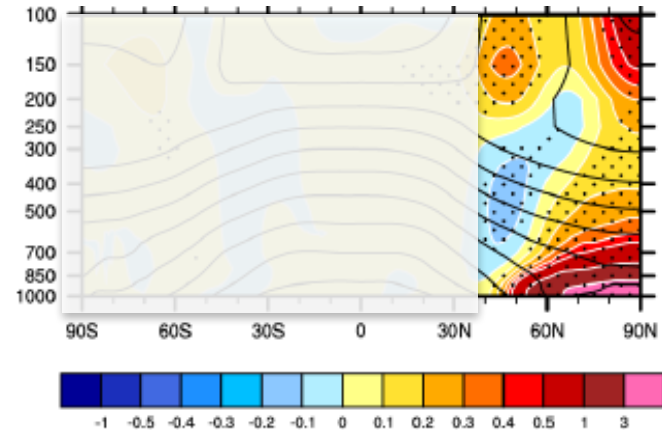
$T_l$  response

# Zonal mean DJF $T$ Response to Sea Ice Loss

Coupled ocean atmosphere GCM



AGCM

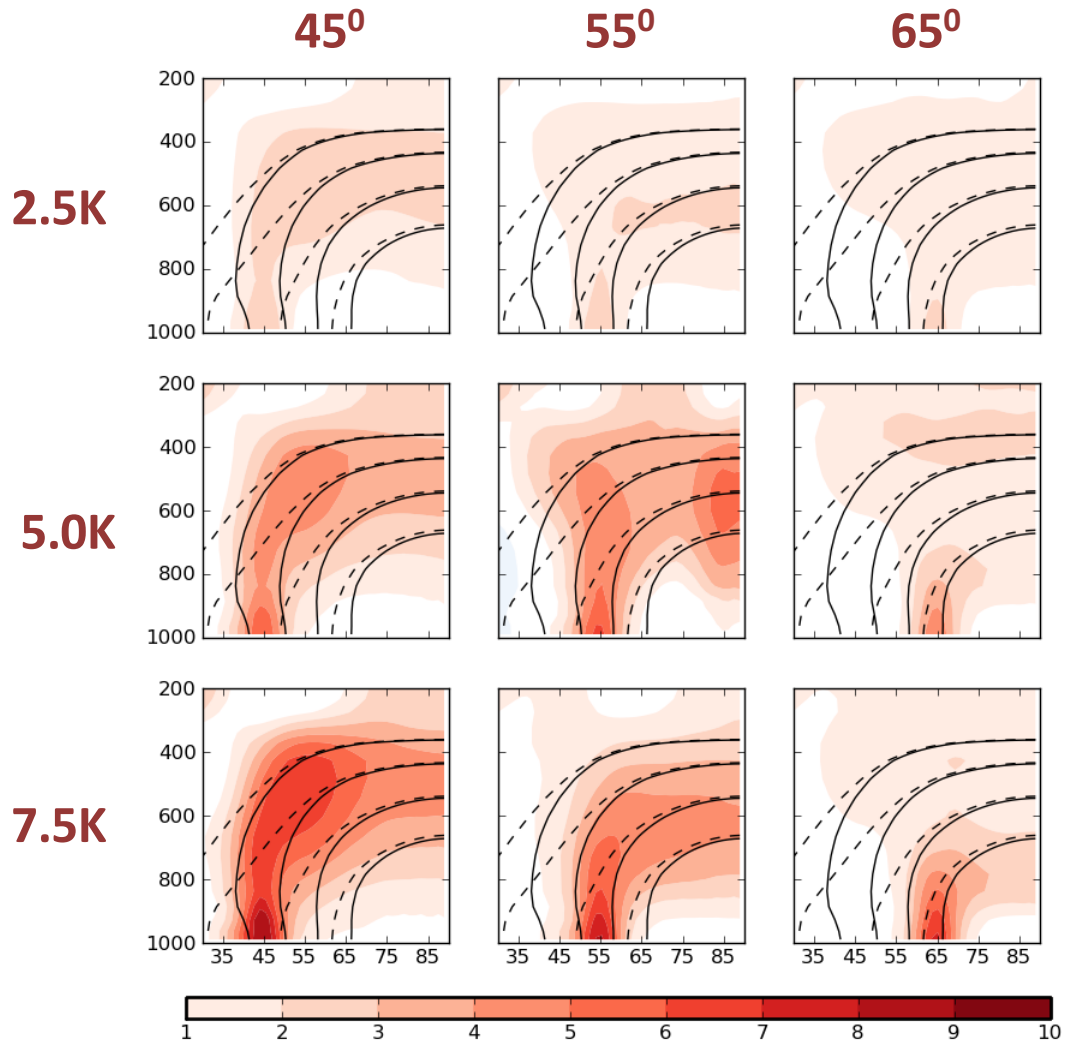


- Why is the Arctic free tropospheric response to sea ice loss amplified in the coupled model?



# Response to Midlatitude Surface Heating

- Simplified aquaplanet GCM.
- Transient (10-20) day response.
- Response roughly organized by  $\theta_e$  distribution.
- Response warms and stabilizes the polar troposphere.



Solid: Control  $\theta$   
Dashed: Control  $\theta_e$

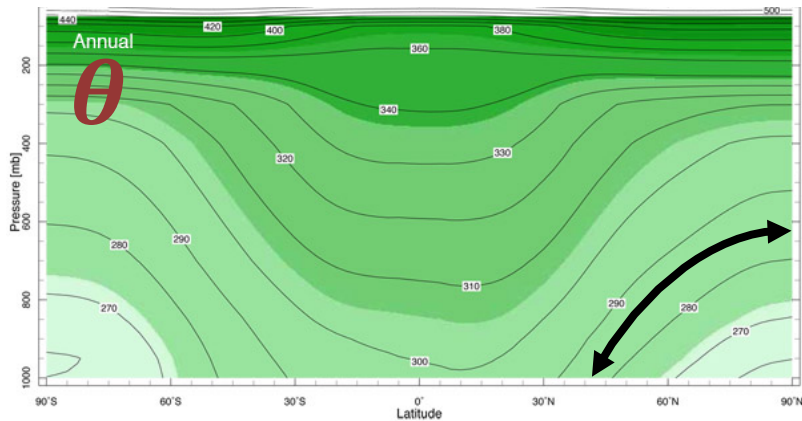
*Fajber, Kushner, Laliberte in prep.*

- **Topics of current interest:**
  - Consistent strengthening of the Aleutian Low
  - The “tug of war” on the circulation response over the North Atlantic
  - The non-additivity in the shoulder seasons.
  - A positive feedback from the North Pacific SST to the Arctic mid troposphere.

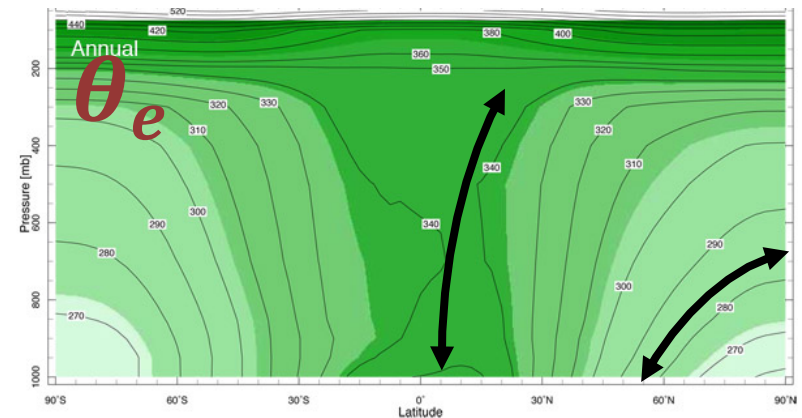
# Extra slides

# Entropy Perspective

Annual Mean Potential Temperature



Annual Mean Moist Potential Temperature

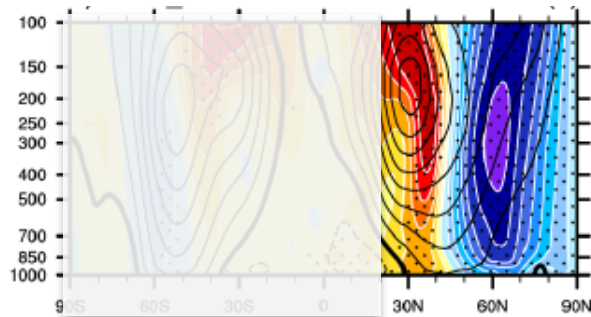


<http://paoc.mit.edu/labweb/notes/chap5.pdf>

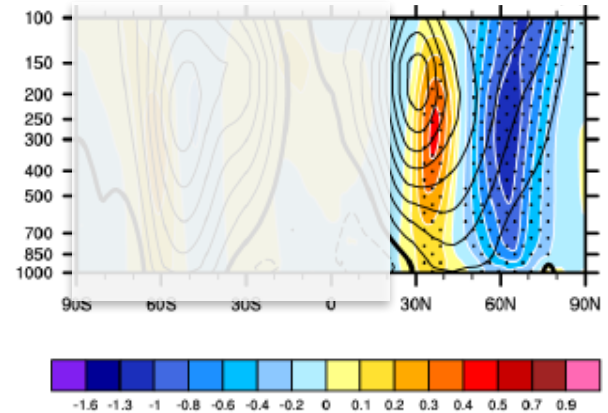
- Moist isentropic circulations in extratropical dynamics (Juckes 2000, Frierson 2008, Wu and Pauluis 2014, Laliberté and Kushner 2014)
- Isentropes link the Arctic troposphere to the midlatitude surface.

# Zonal mean DJF $U$ Response to Sea Ice Loss

Coupled ocean atmosphere GCM

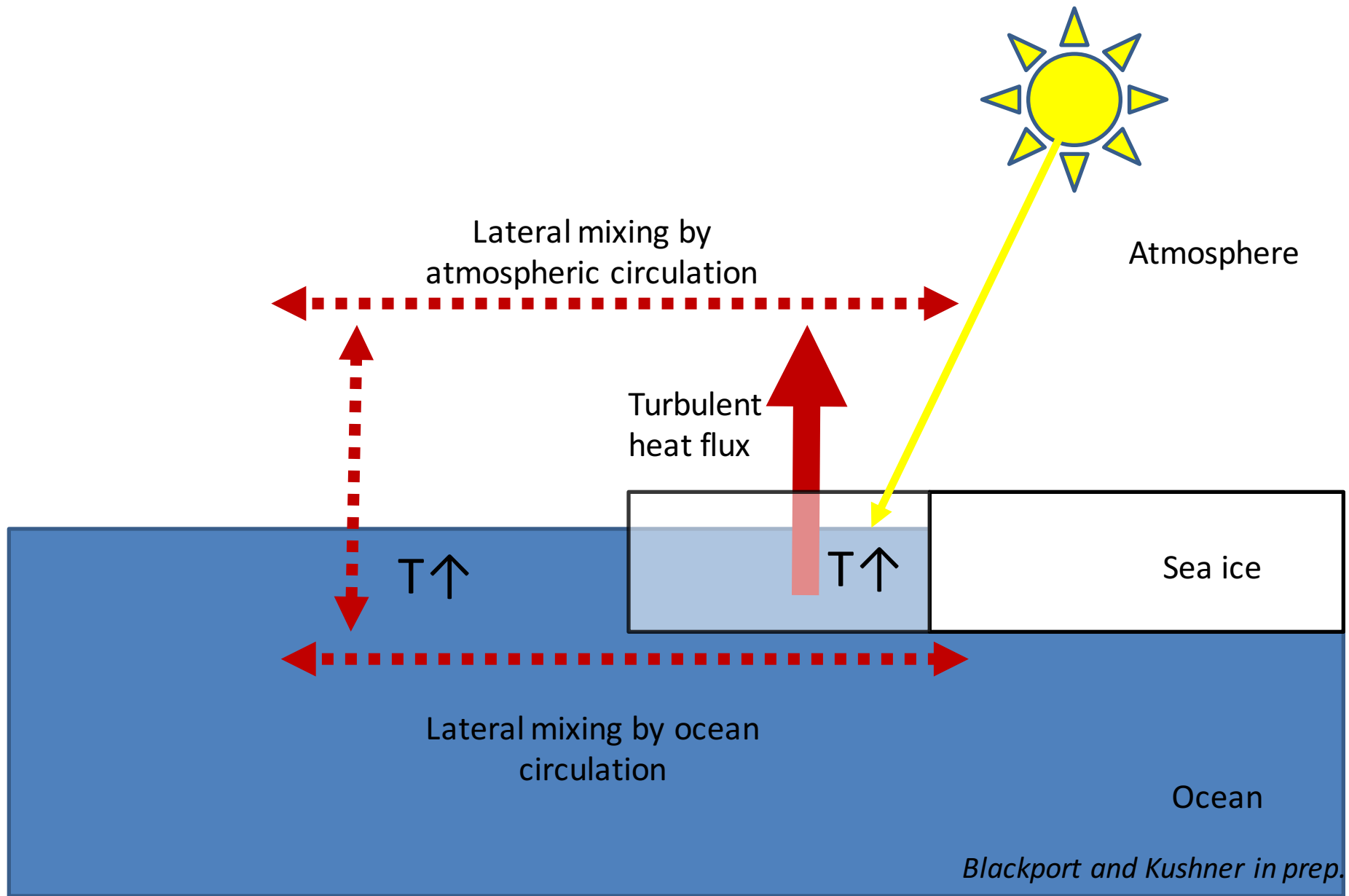


AGCM



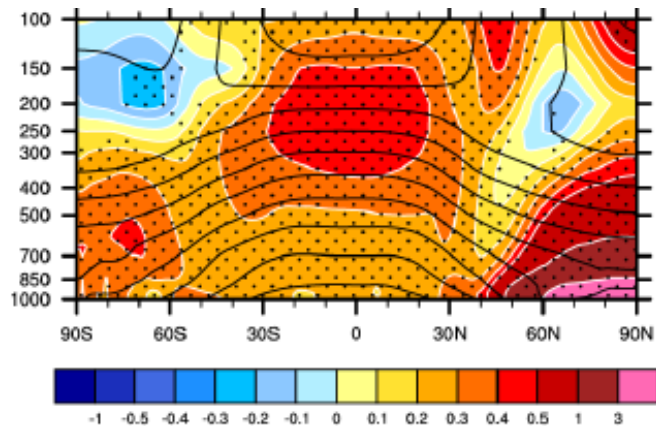
- Circulation is also impacted by coupling.

# Ocean warming from sea-ice loss

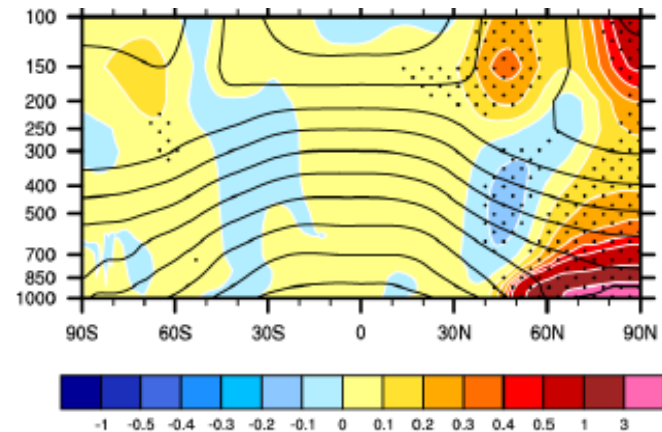


# Zonal mean DJF $T$ Response to Sea Ice Loss

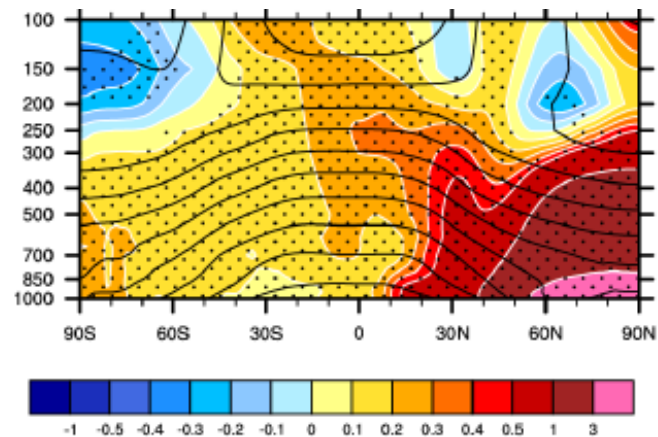
Coupled ocean atmosphere GCM



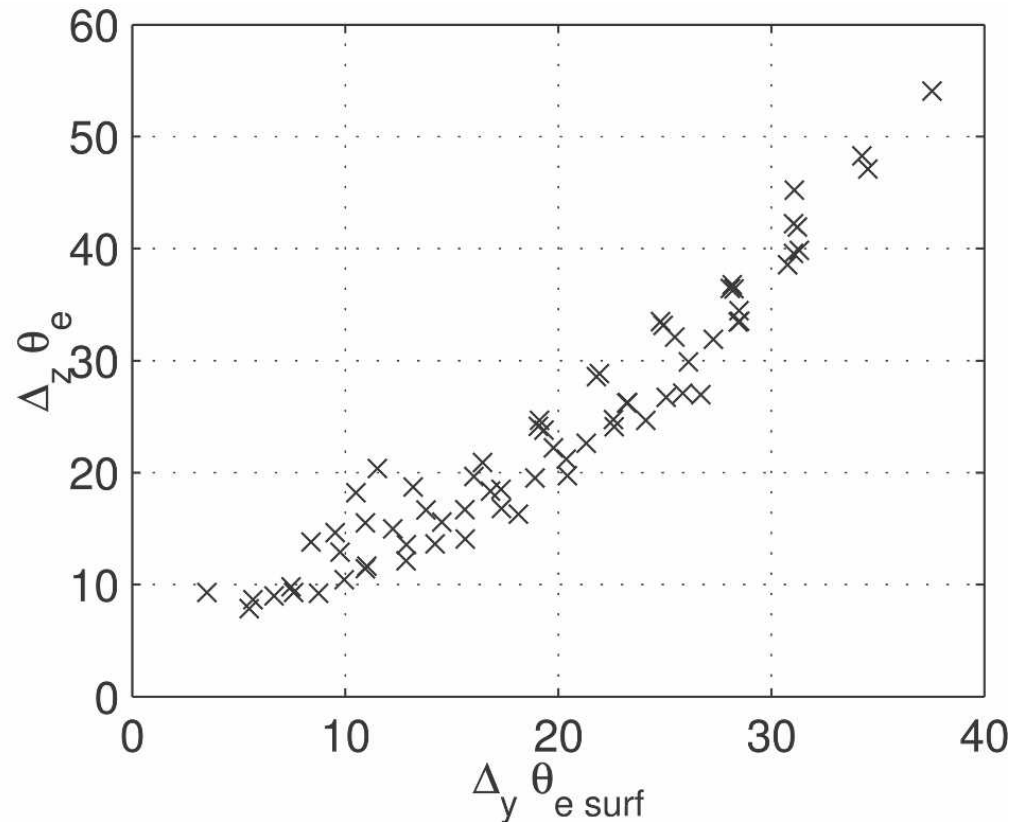
AGCM



Slab Ocean Model + AGCM



# Midlatitude Responses in GCMs (Frierson 2008)



- Bulk stability up to tropopause versus  $\theta_e$  gradient of baroclinic zone in aquaplanet comprehensive GCMs.

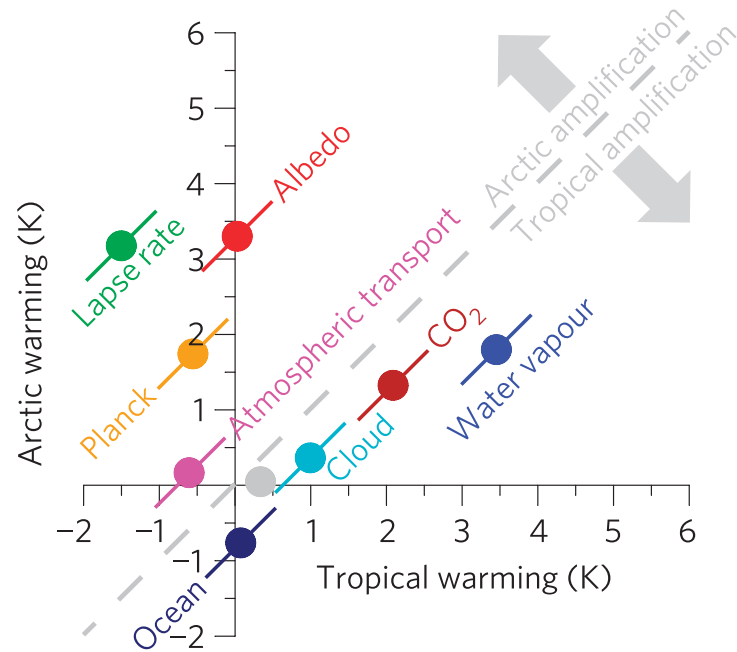


# Questions

- Can we improve on the prescribed SST framework to study this feedback process?
- Moist isentropic mechanisms are better established in summer/shoulder seasons. In winter, can it be disentangled from other mechanisms, such as tropical moisture driving surface polar amplification (Sukyoung Lee et al.)?
- What are the broader implications of this enhanced Arctic tropospheric warming?

# How might Arctic free tropospheric warming impact surface Arctic amplification?

**a** Annual warming (TOA perspective)

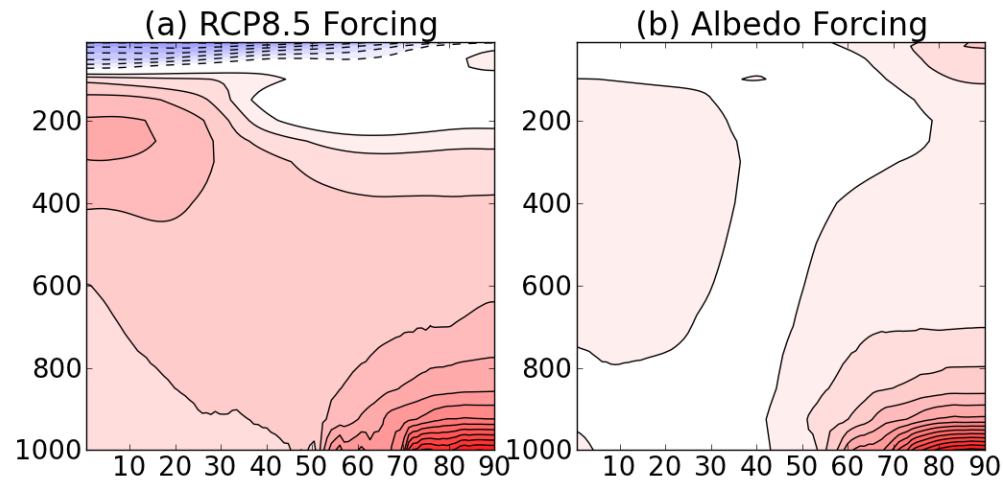


*Pithan & Mauritsen 2014*

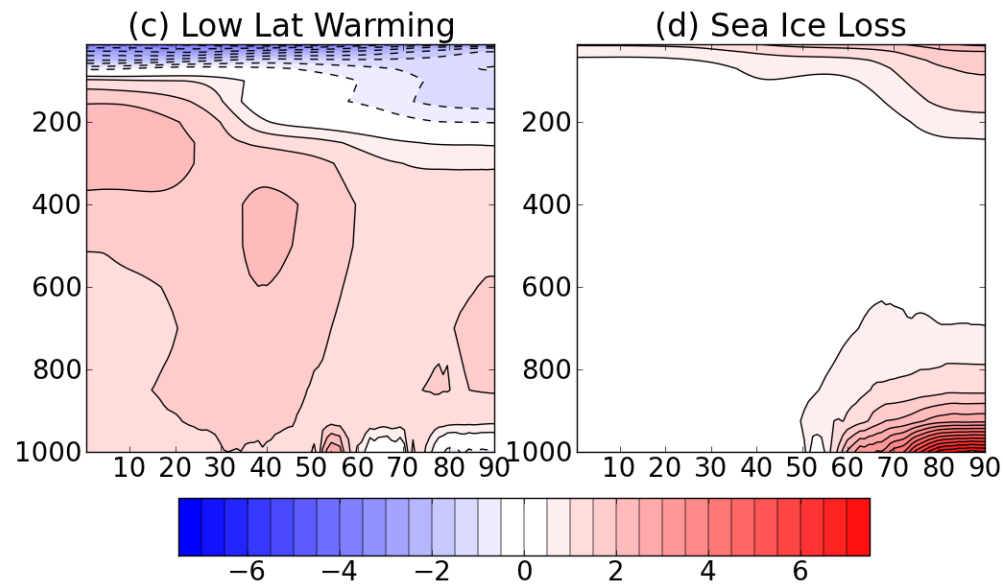
- Coupled response to sea ice loss influences Arctic lapse rate and related radiative feedbacks.

# $\Delta[T]$ , DJF

**Coupled Model  
Experiments**



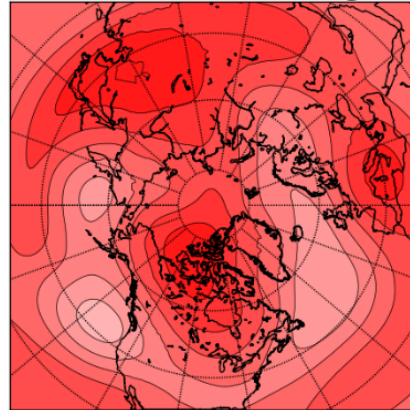
**Diagnosed  
Decomposition**



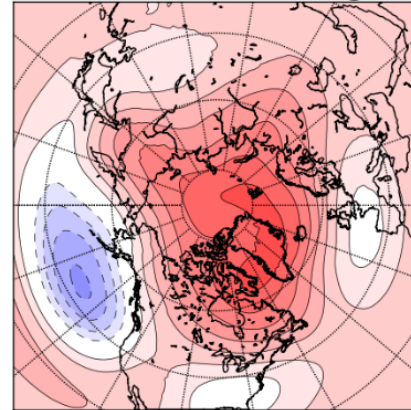
# $\Delta Z500$ , DJF

Coupled Model  
Experiments

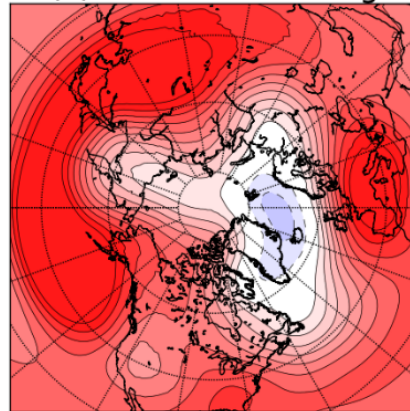
(a) RCP8.5 Forcing



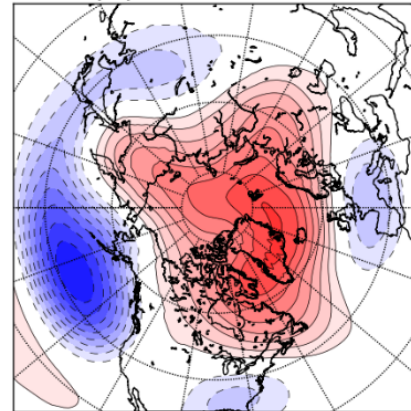
(b) Albedo Forcing



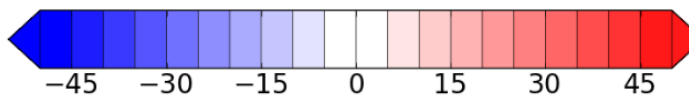
(c) Low Lat Warming



(d) Sea Ice Loss



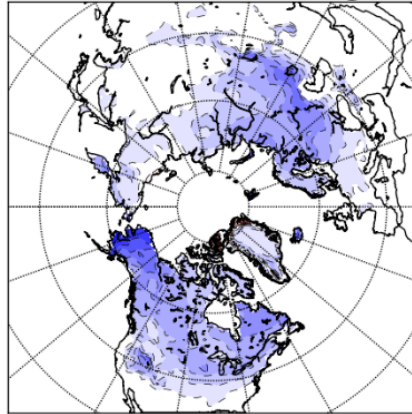
Diagnosed  
Decomposition



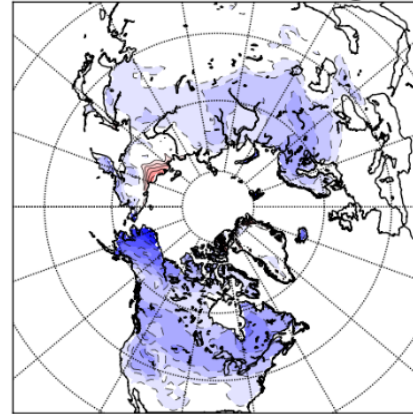
# Change in subseasonal $T_{2m}$ variability, DJF

Coupled Model  
Experiments

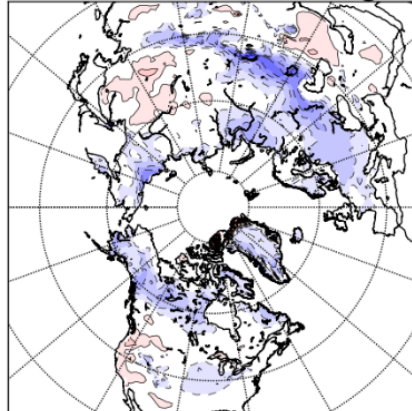
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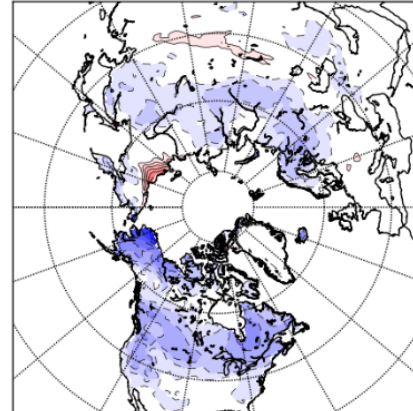
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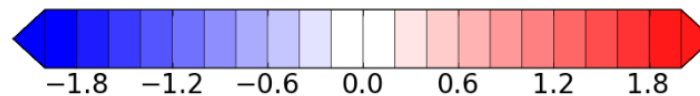
(c) Low Lat Warming



(d) Sea Ice Loss

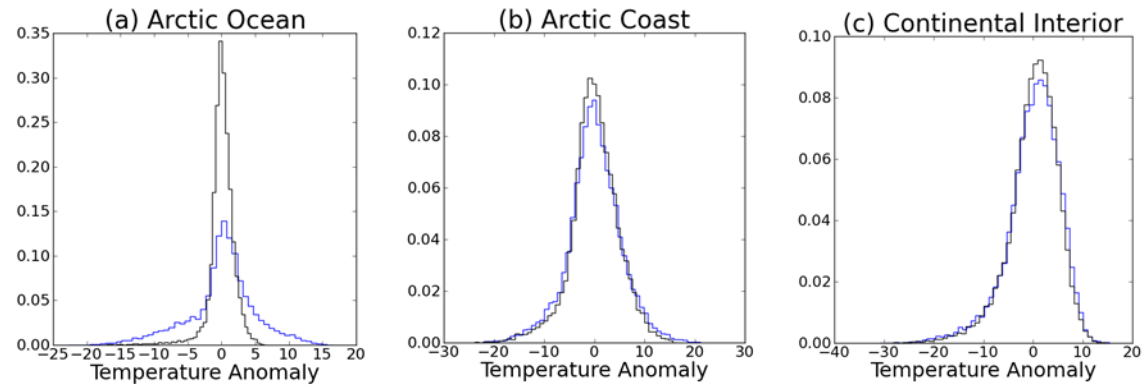


Diagnosed  
Decomposition

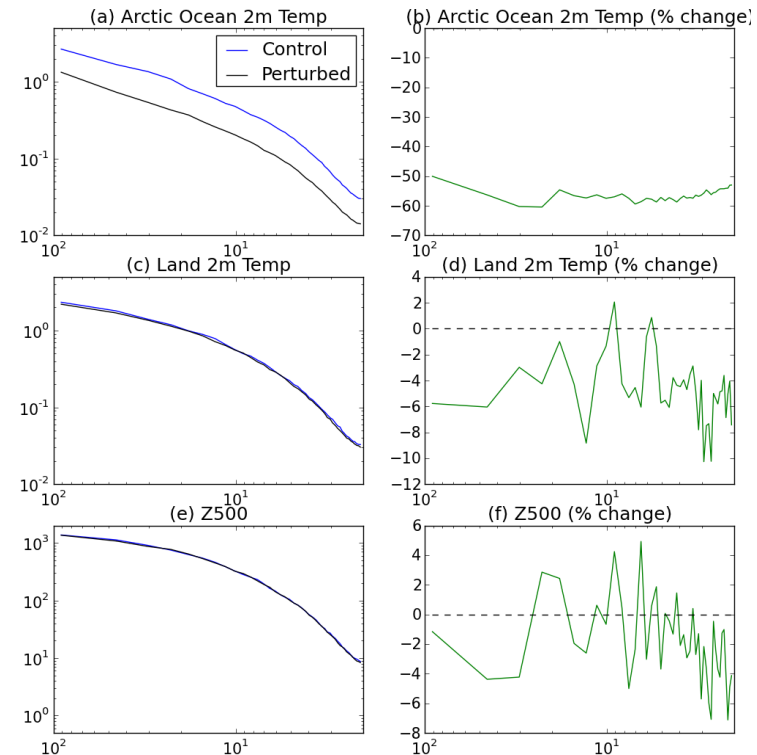


## Distributions of T2m at Selected Points, CCSM4

- Sea ice loss creates a more maritime environment.
- Stronger ocean-atmosphere coupling and polar amplification reduce temperature variability over the ocean and close to coasts.
- We generally find that polar amplification leads to a weakening of the storm tracks, but with lots of structural details.



## Time Spectra for CCSM4 Experiments



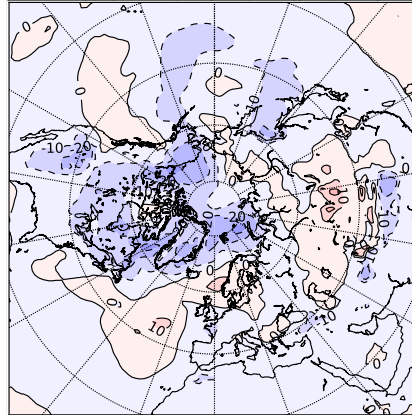
*Blackport and Kushner 2016*



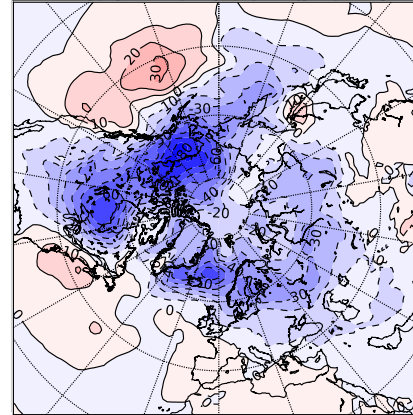
# Change in 2-6 Day Band Pass SLP, DJF

**Coupled Model  
Experiments**

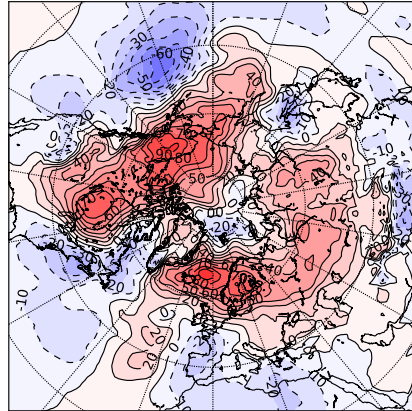
(a) RCP8.5 Forcing



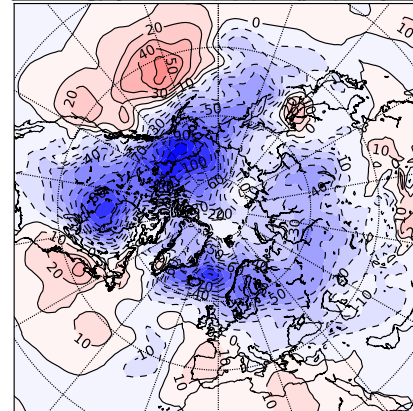
(b) Albedo Forcing



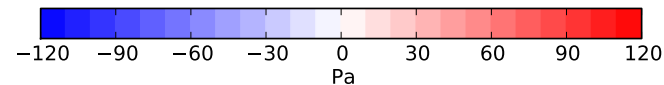
(c) Low Lat Warming



(d) Sea Ice Loss

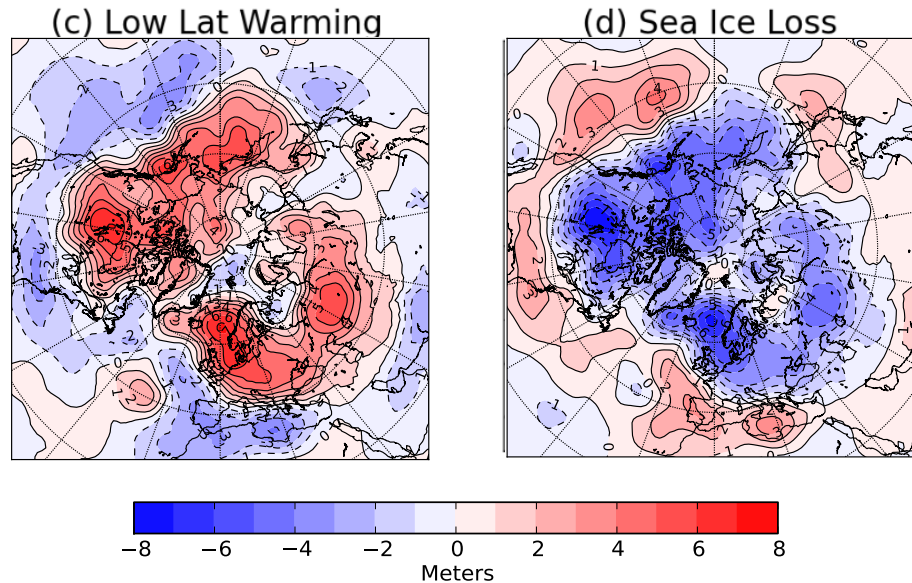


**Diagnosed  
Decomposition**

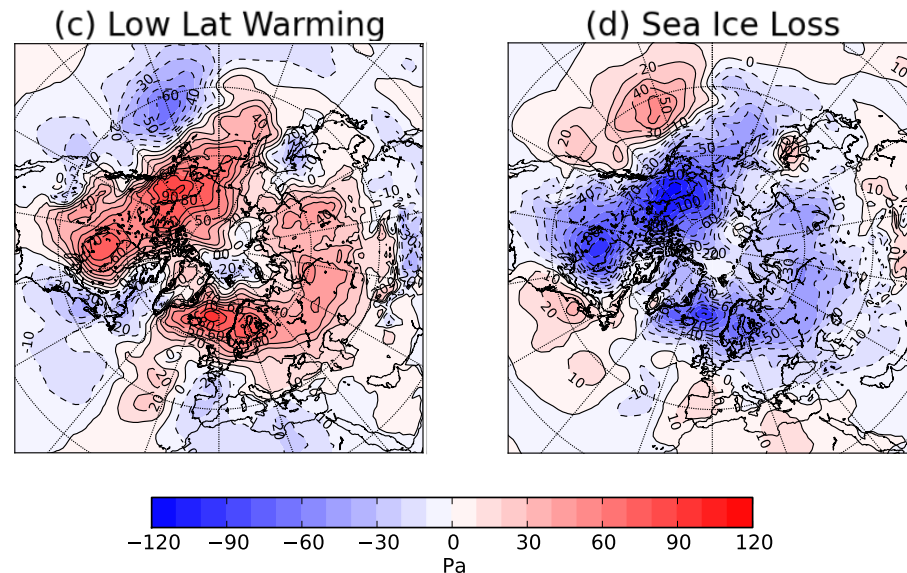


# Change in 2-6 Day Band Pass SLP and Z500, DJF

Diagnosed  
Decomposition  
2-6 Day Band  
Pass Z500



Diagnosed  
Decomposition  
2-6 Day Band  
Pass SLP

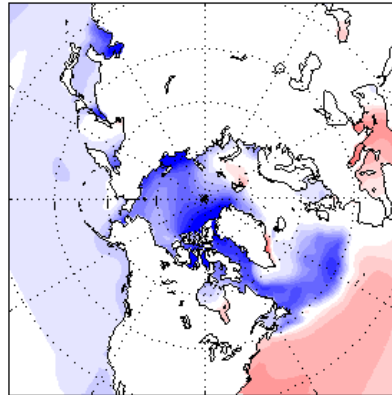




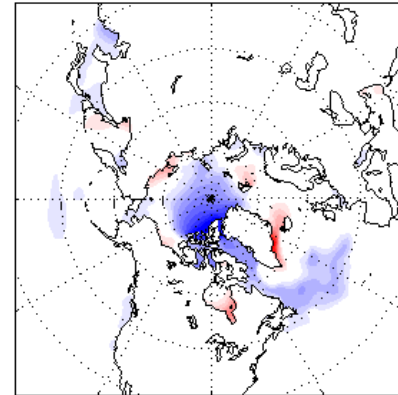
# DJF Sea Surface Salinity Response

**Coupled Model  
Experiments**

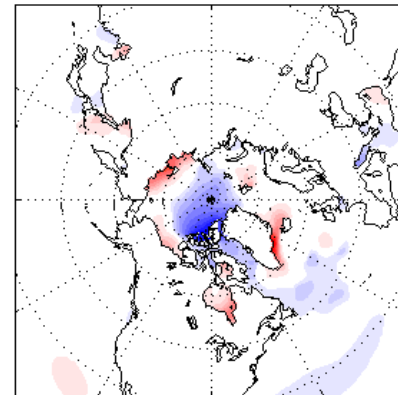
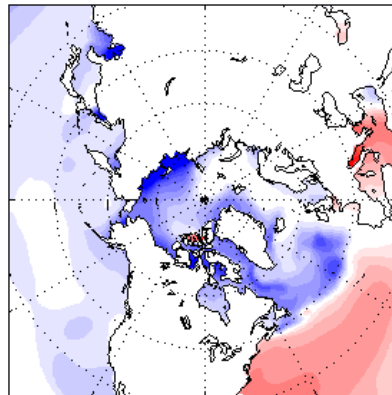
(a) RCP



(b) alb

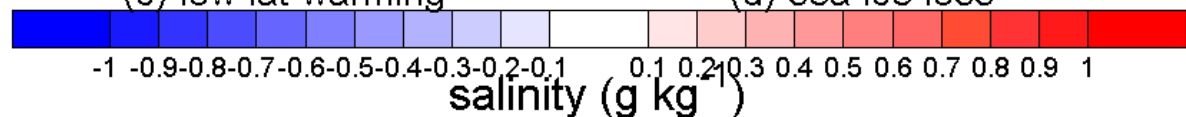


**Diagnosed  
Decomposition**



(c) low lat warming

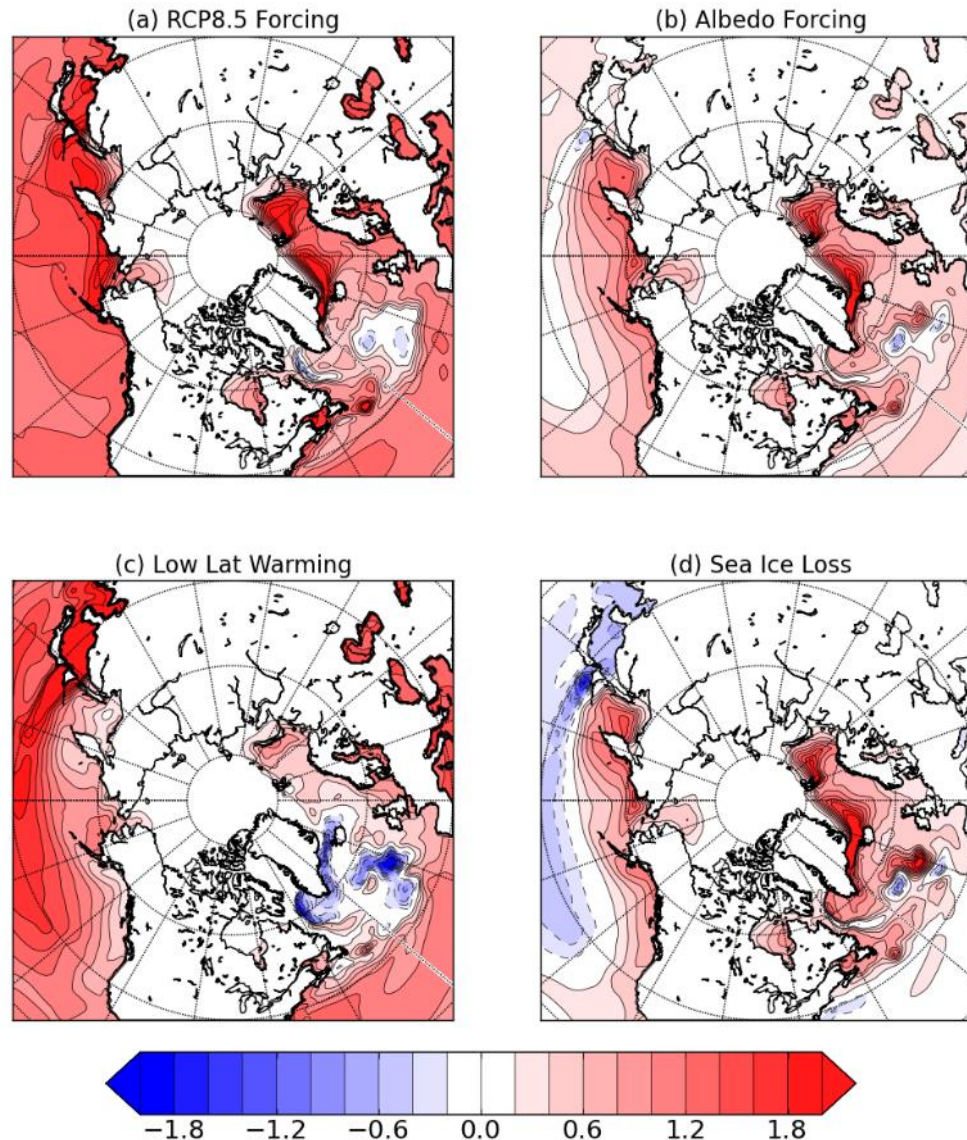
(d) sea ice loss



# DJF Ocean Surface Temperature Response

Coupled Model  
Experiments

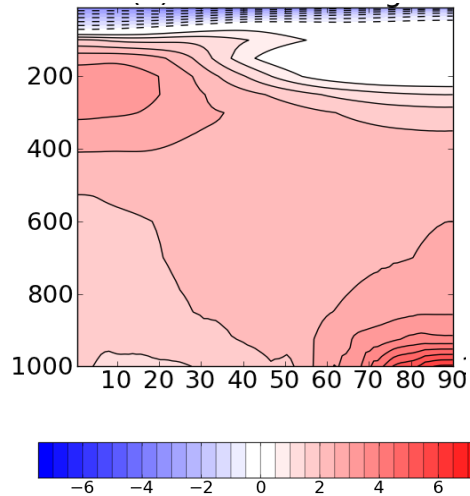
Diagnosed  
Decomposition



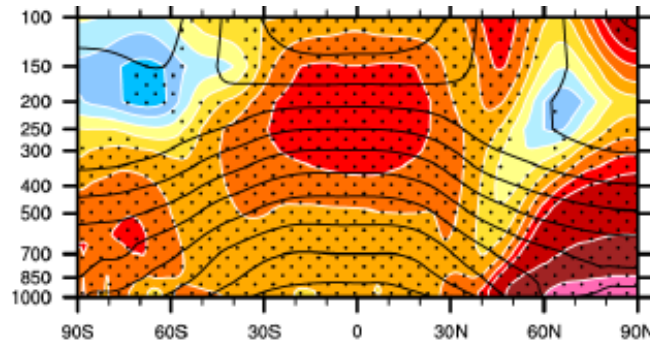
Hay et al. in prep  
Blackport and Kushner in prep.

# Can we Isolate the Impact of Sea Ice Loss on Circulation?

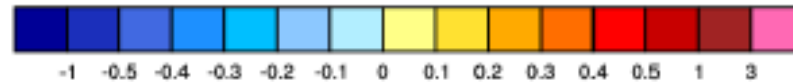
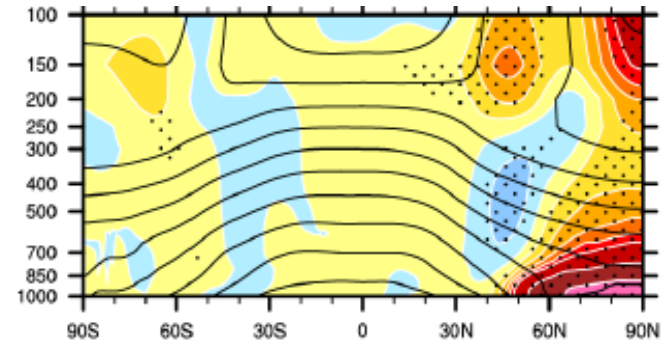
CESM1 RCP8.5  $\Delta[T]$



CCSM4 Sea Ice Loss  $\Delta[T]$



CAM4 Sea Ice Loss  $\Delta[T]$



*Deser et al. 2015*