

Mechanisms of changes in precipitation and droughts driven by natural variability and external forcings

Celine Bonfils

Sept 12, 2018

AGCI, Aspen, Colorado

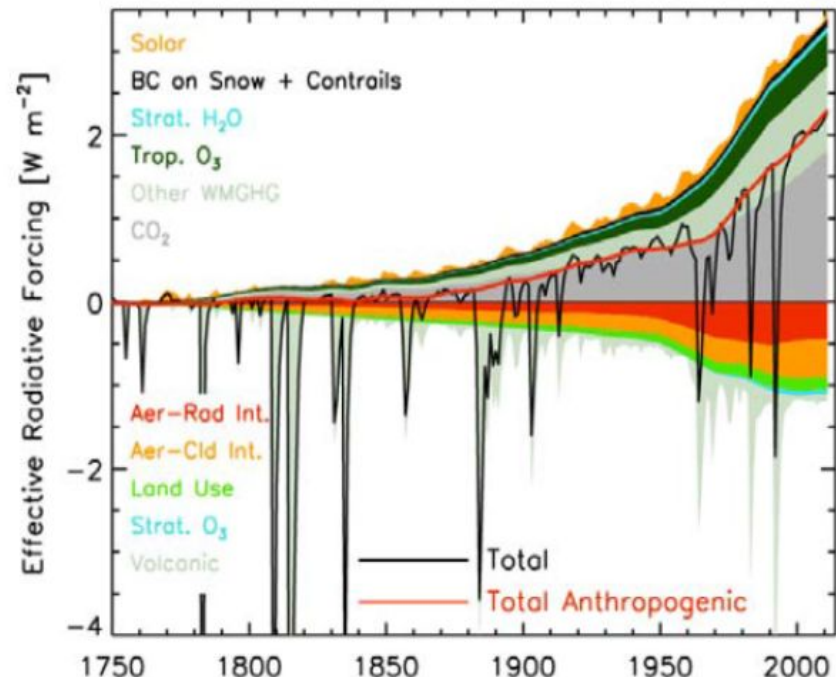
When the Rain Stops: Drought on Subseasonal and Longer timescales

 Lawrence Livermore
National Laboratory

LLNL-PRES-740731

This work was performed under the auspices of the U.S. Department of Energy by Lawrence Livermore National Laboratory under Contract DE-AC52-07NA27344. Lawrence Livermore National Security, LLC

Variation with Time of Climate Forcings:



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When the Rain Stops: Drought on Subseasonal and Longer timescales

Young Buck



“when the, rain stops
I'm doughnut ridin'
and I can't stop”

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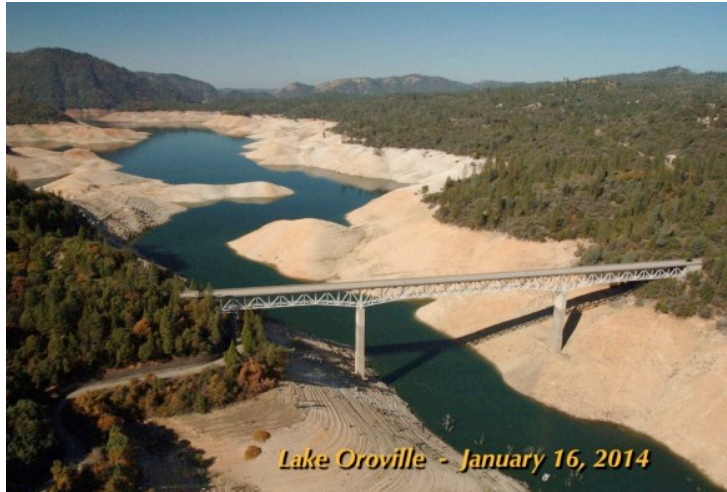


"when the, rain stops
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and I can't stop"



Starts with a flood in a desert region in Australia in 2039 "About self-identity, forgiveness and love"

Motivation

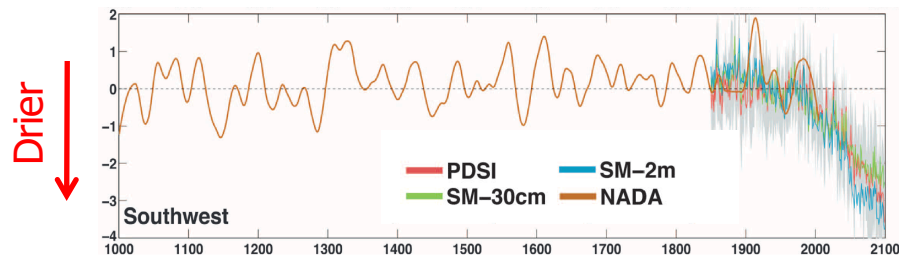


Climate models project unprecedented drought risk and increased aridity

Science questions:

- Are changes in drought behavior becoming increasingly driven by human forcing?
- Can we identify these human-induced changes in observed climate records?

Unprecedented 21st century drought risk in the American SW (Cook et al. 2015)



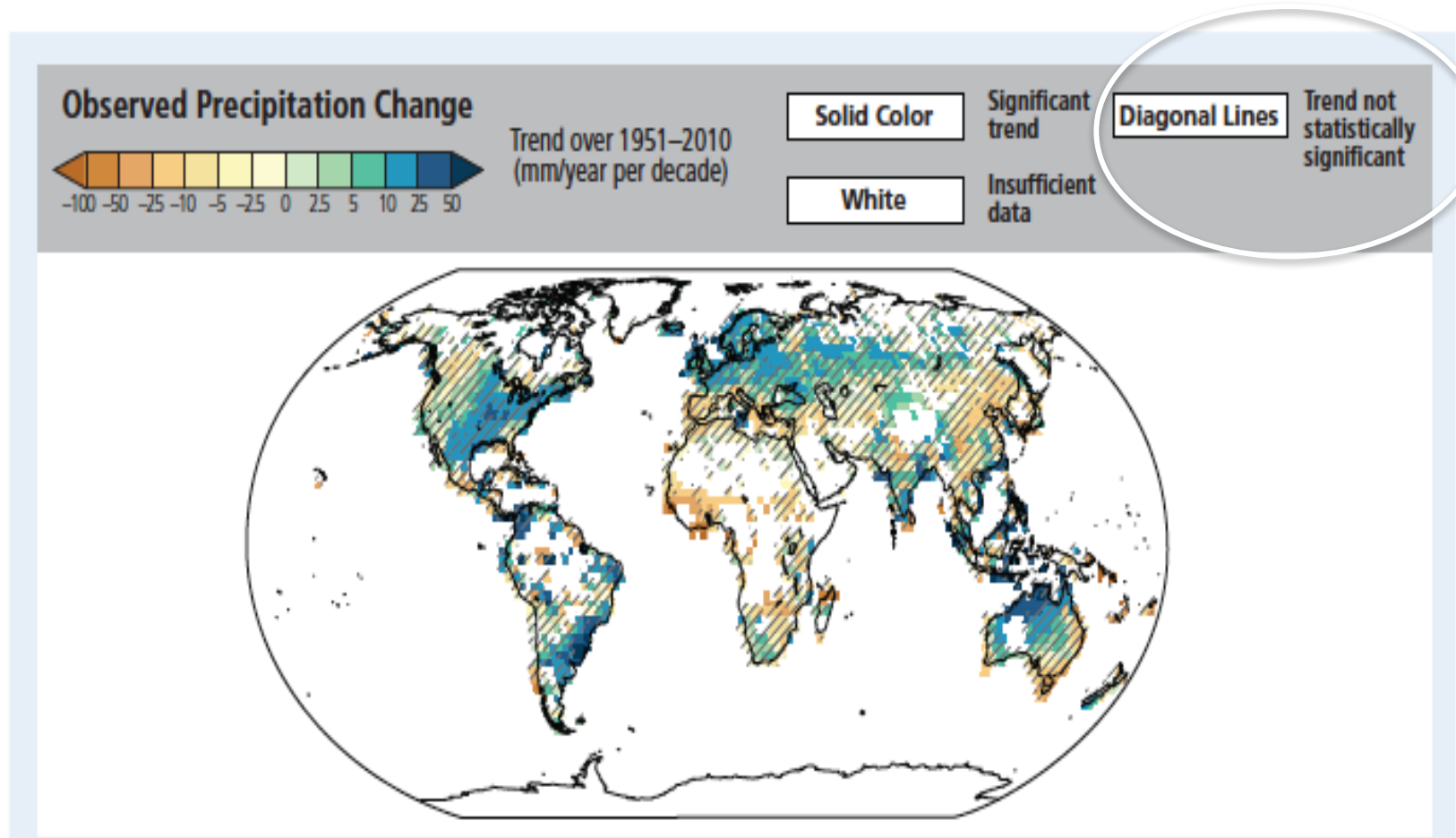
→ Focused first on large-scale precursors of droughts and long-term trends

Improve our understanding of the nature and causes of past/future droughts

Influences on droughts/aridification

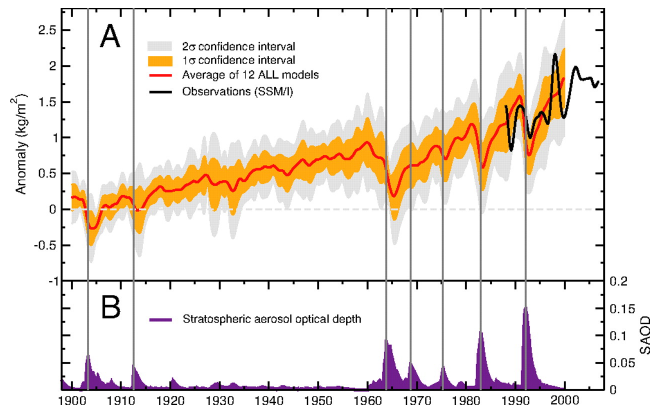
	Moisture supply: Precipitation	Evaporative demand
Gradual change in mean	#1. Discernable human-induced changes in the observed rainfall patterns Marvel and Bonfils 2013	
Gradual change in variance		

How do we identify a human influence on global precipitation?



First, we need to identify some mechanisms that are expected to drive the future changes in P

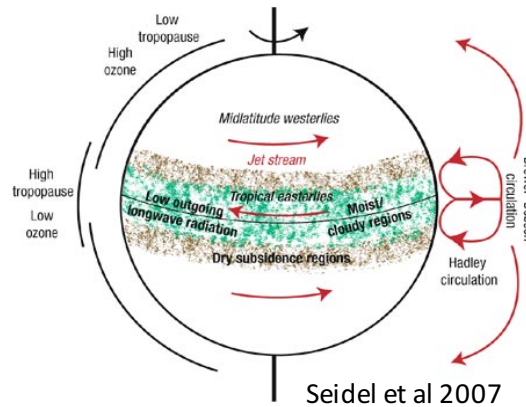
Thermodynamic



Santer et al 2007

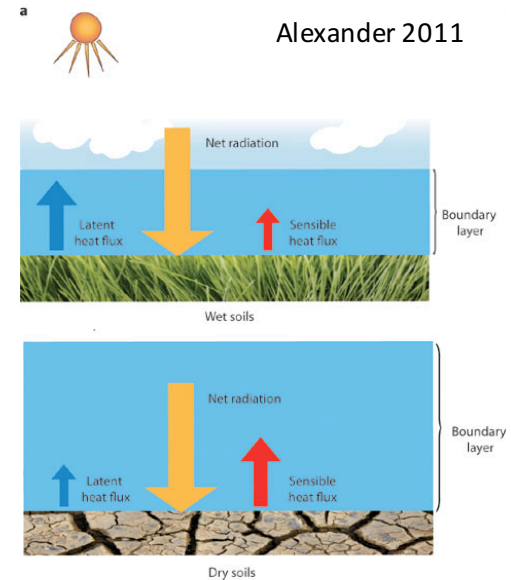
Intensification of current zonal wet-dry patterns

Dynamic



Latitudinal redistribution of global precipitation

Coupled



Land-atm feedbacks

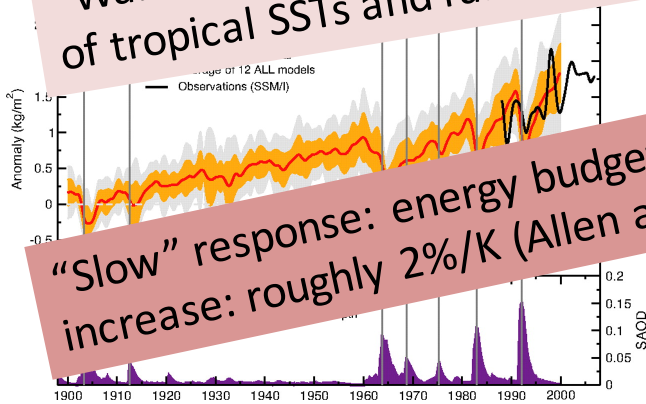
First, we need to identify some mechanisms that are expected to drive the future changes in P

Thermodynamic

“Warmer-gets-wetter” mechanism (link the patterns of tropical SSTs and rainfall changes) (e.g., Solomon et al. 2010)

Fast P response to CO₂ forcing (e.g., Richardson et al. 2016)

“Slow” response: energy budget constraints on P increase: roughly 2%/K (Allen and Ingram 2002)



Santer et al 2007

Intensification of ENSO-driven P responses in the 21st century (Power et al. 2013, Bonfils et al. 2015)

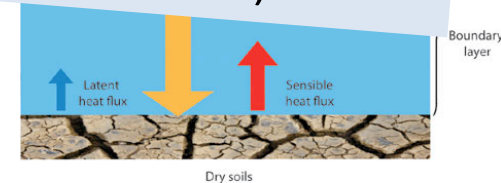
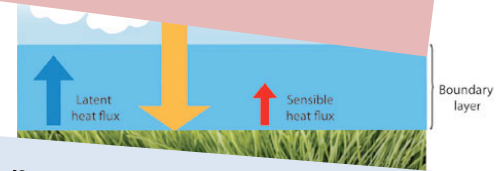
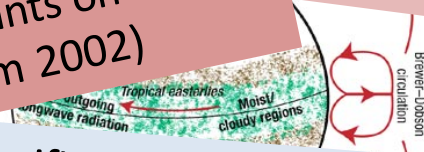
Intensification of ENSO-driven P responses in the 21st century (Power et al. 2013, Bonfils et al. 2015)

WWDD theory does not work well over land (Greve et al. 2014)...

Changes in annual cycle of P, shift in ITCZ...

... unless seasonality is considered (Kumar et al. 2015)

Coupled



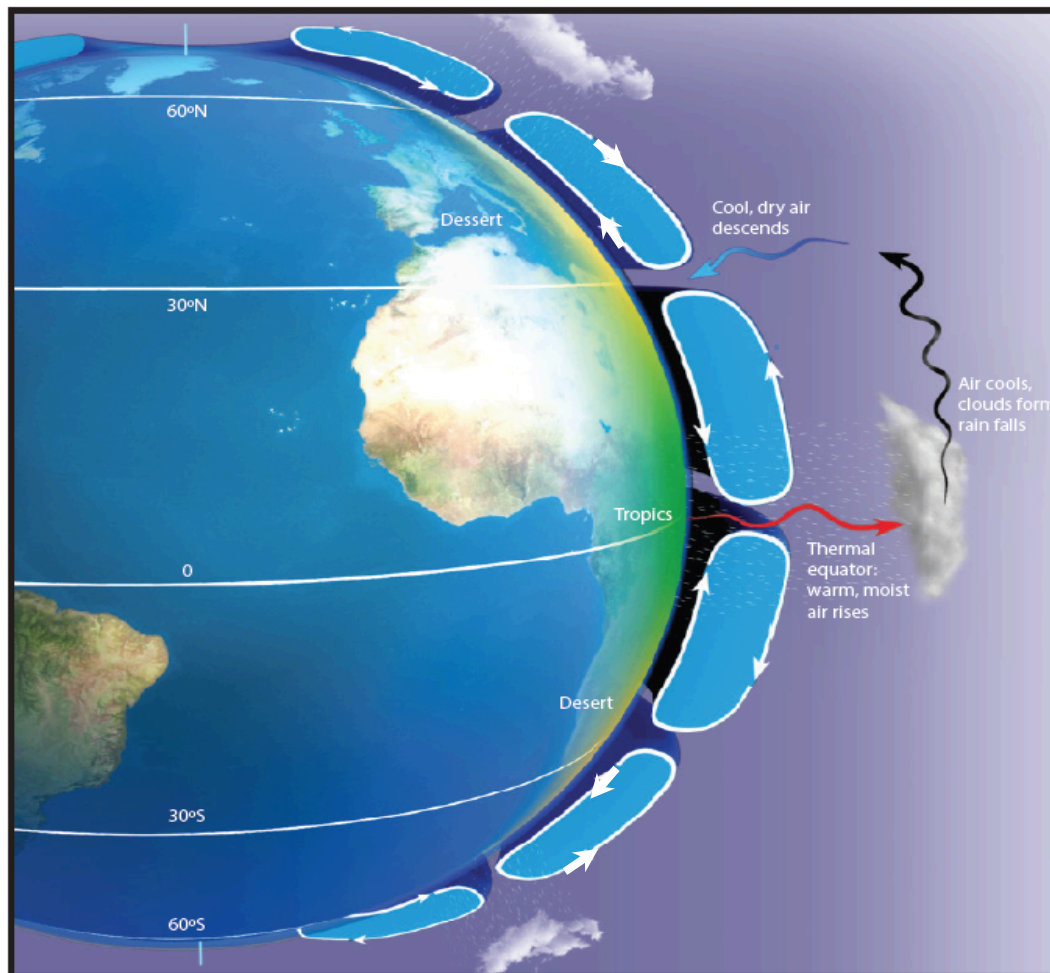
Dry soils

feedbacks

Then, we pick a mechanism and tried to detect a human influence

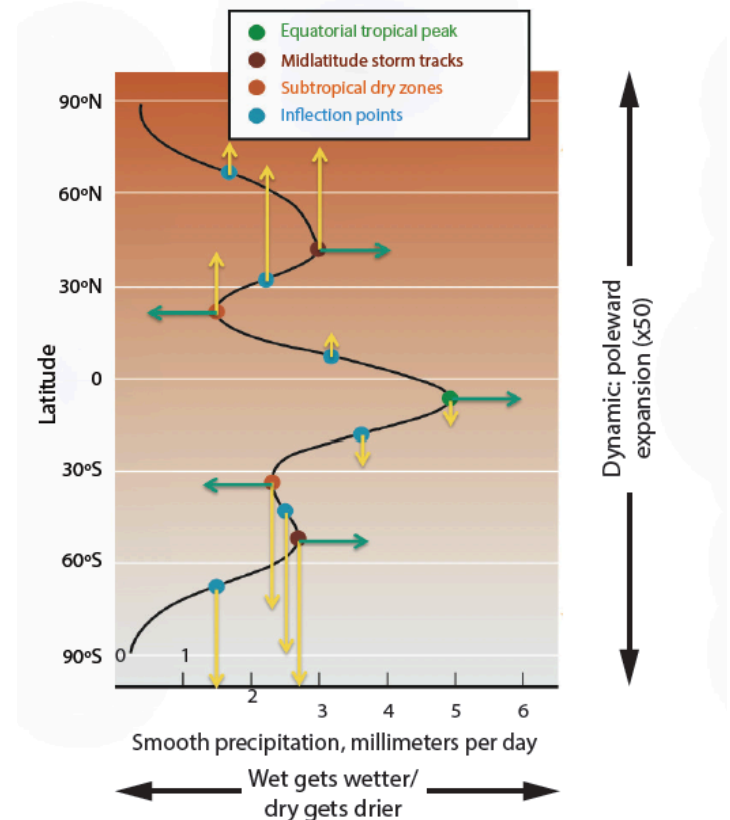
Eg: Natural and Human Influences on Changing Zonal-Mean Precipitation (Marvel & Bonfils, 2013 PNAS)

Theories and climate projections predict a latitudinal **intensification** and **poleward shifts** of global precipitation

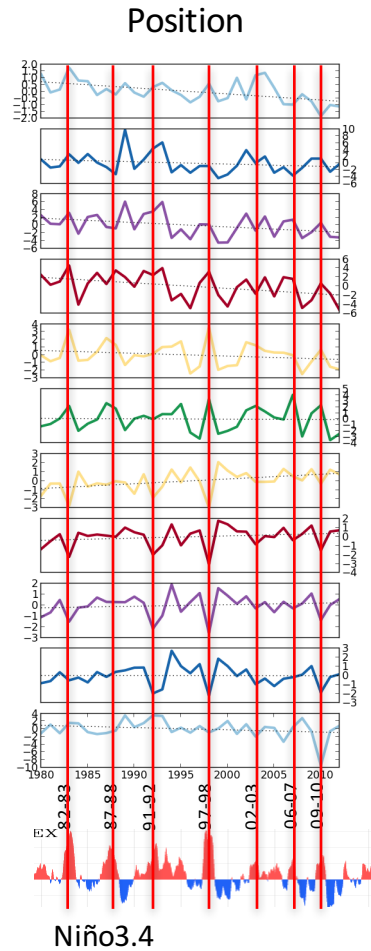
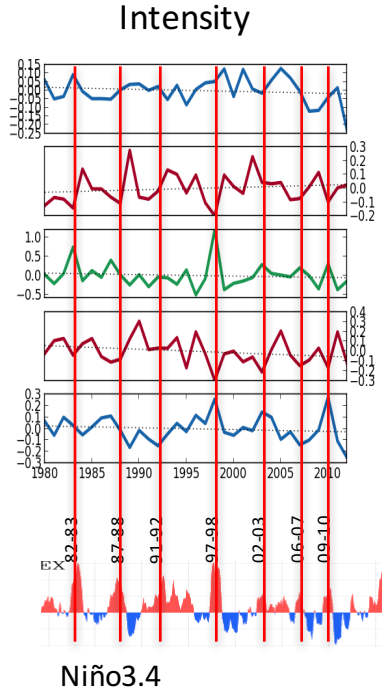


Human fingerprint = expected response to human forcings from 70+ CMIP5 runs of historical climate change (1980-2012)

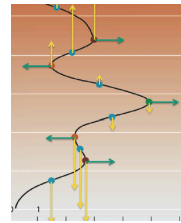
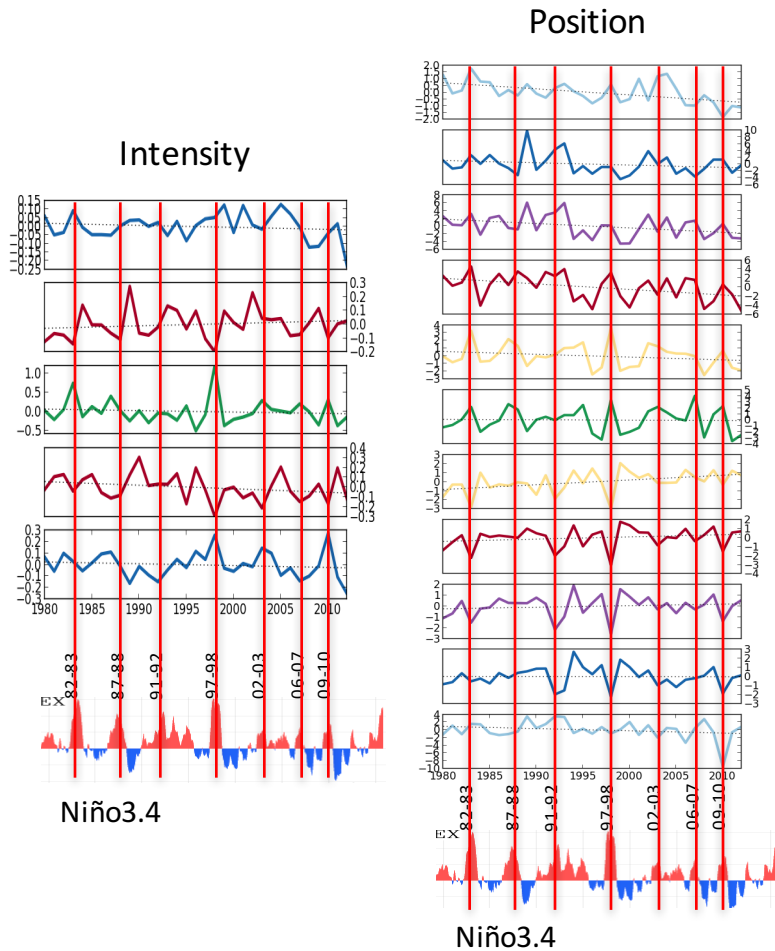
= Distortion of the curve in response to human forcing



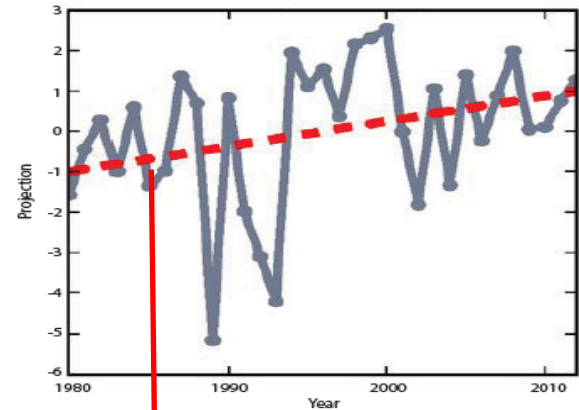
Projection of **observed** Intensity and Position anomalies



Projection of **observed** Intensity and Position anomalies



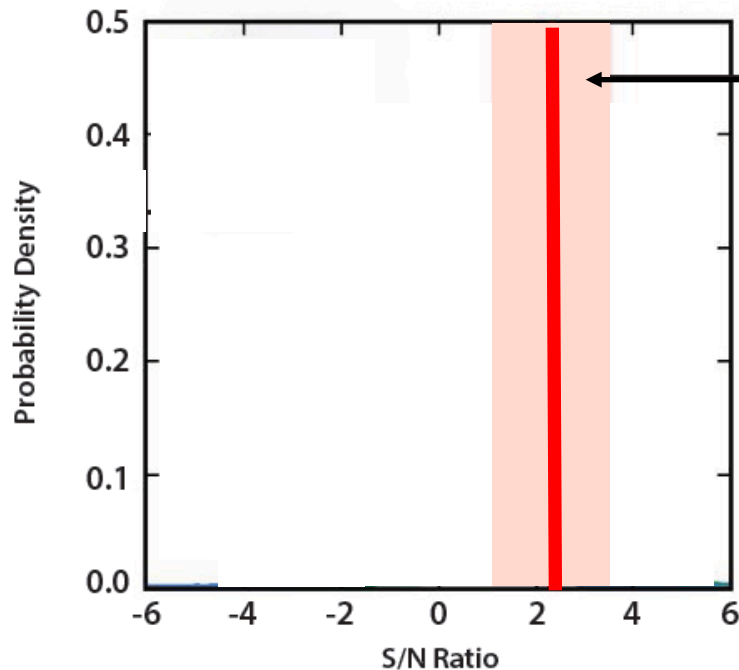
SIGNAL time-series = measure of similarity between obs and fingerprint



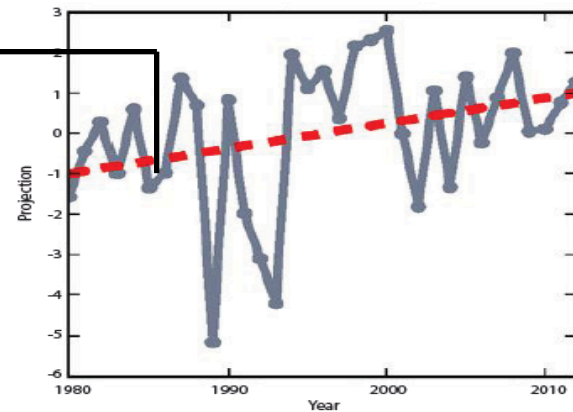
Positive trend: fingerprint is present and growing in the observations

Projection of **observed** Intensity and Position anomalies

Observed/Model Projections on $F_M(D,T)$
(33-year trends)

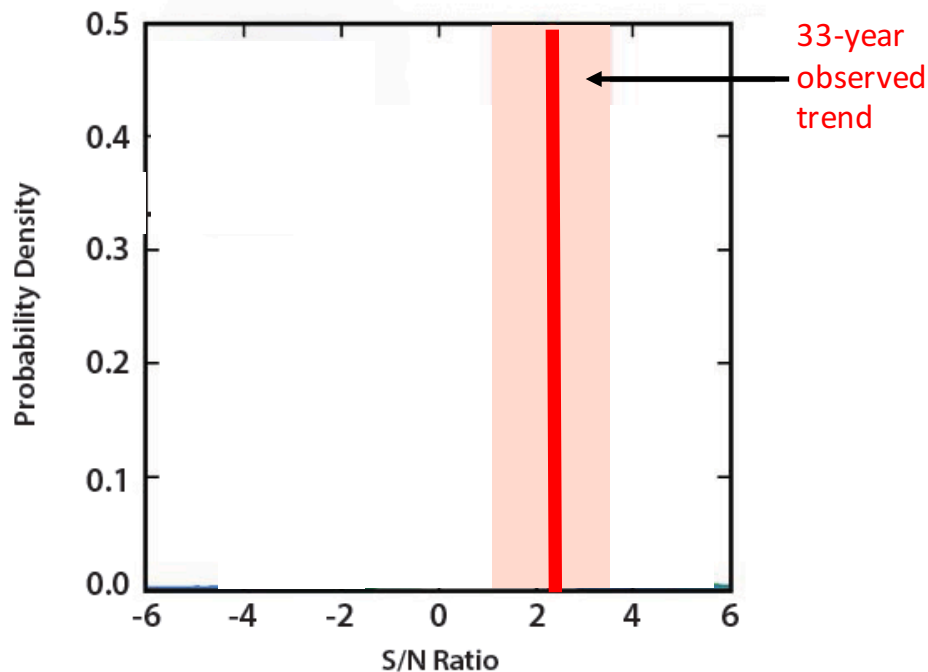


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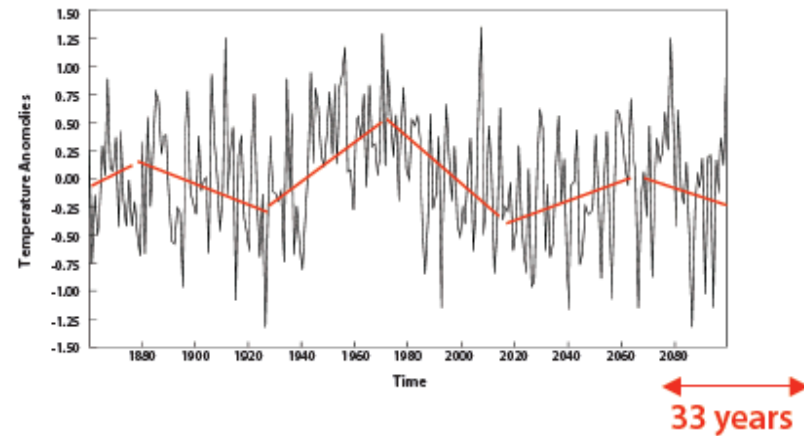


Projection of 20k yrs of **unforced** Intensity and Position anomalies

Observed/Model Projections on $F_M(D,T)$
(33-year trends)

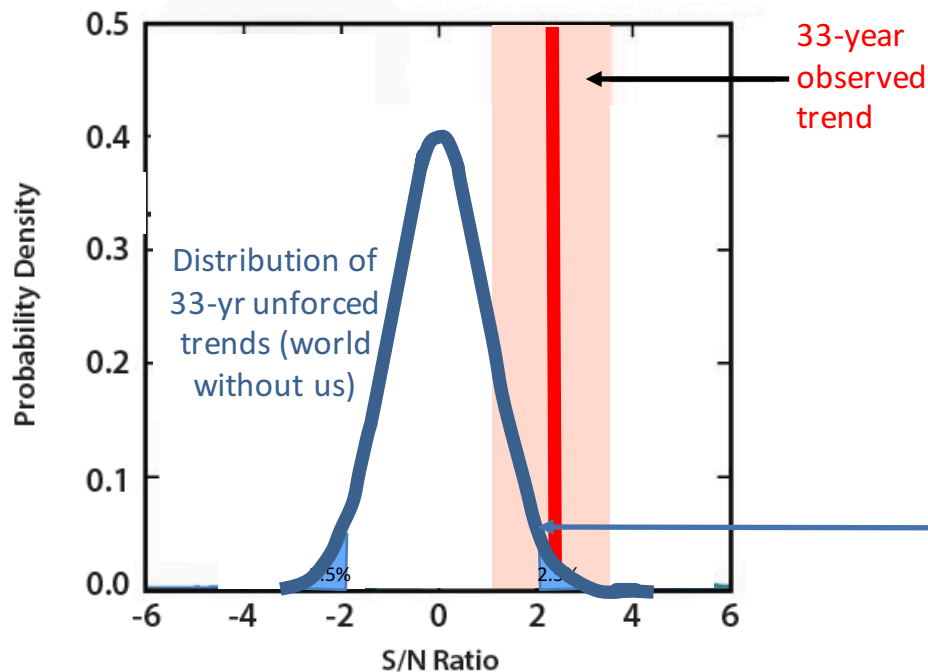


NOISE time-series (no reason to project, except by chance)

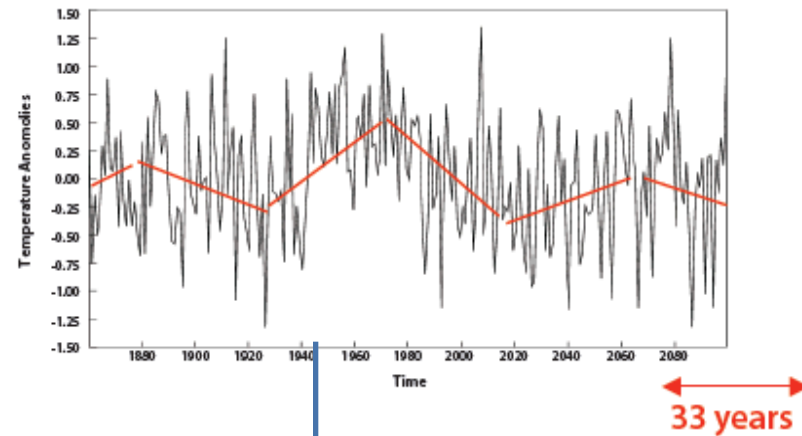


Observed trend cannot be explained by internal variability

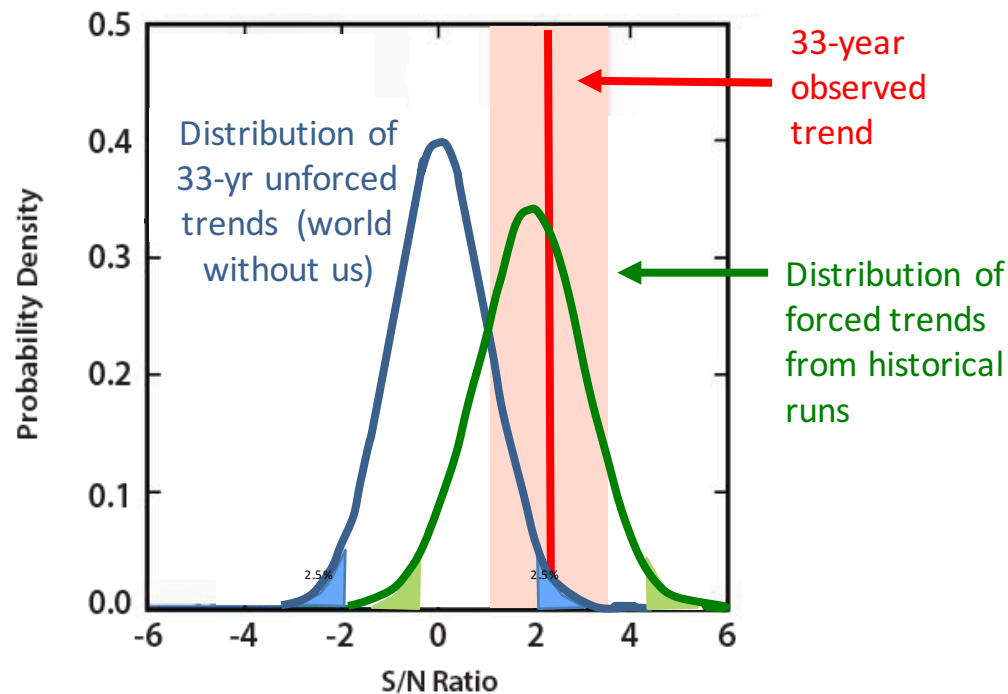
Observed/Model Projections on $F_M(D,T)$
(33-year trends)



NOISE time-series (no reason to project, except by chance)



#1 We detected a human influence on changing zonal-mean precipitation



- ❑ **Detection:** The observed intensification + poleward expansion of zonal P cannot be explained by climate noise alone.
- ❑ **Attribution:** The observations matches predictions from simulations with combined natural and human forcings

Influences on droughts

	Moisture supply: Precipitation	Evaporative demand
Gradual change in mean	Marvel and Bonfils 2013	
Gradual change in variance	#2. Increase in ENSO-driven precipitation variability Bonfils et al. 2015	

In the 21st century, ENSO-driven precipitation variability is intensified

Goal

ENSO is the main trigger of precipitation variability

- Models do not agree on how ENSO characteristics will evolve in the future

Unclear whether the precipitation response to ENSO will change in the future, even if ENSO remains unchanged

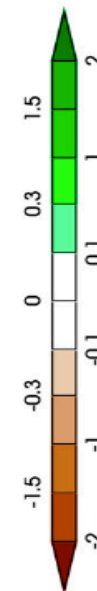
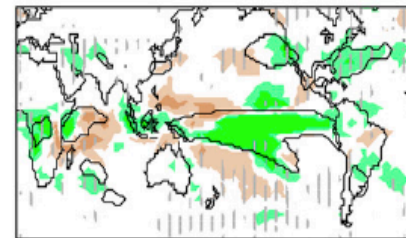
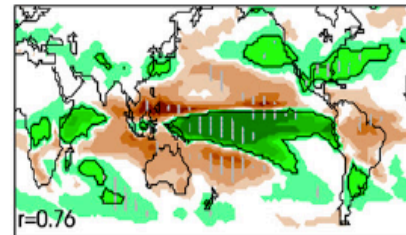
- Non-uniform warming in tropical Pacific Ocean
- Atmospheric circulation change
- Moister atmosphere

Results

Basic 20th century
P response to
ENSO events

Additional
P response to
ENSO
in the future

Intensification



Influence on droughts

	Moisture supply: Precipitation	Evaporative demand
Gradual change in mean	Marvel and Bonfils 2013	#3. Most regions where aridity/moistening is currently regulated by ENSO variability will become more arid in the future Bonfils et al. 2016
Gradual change in variance	Bonfils et al. 2015	

#3 Influence of anthropogenic climate change on regional aridity

Objectives

Investigate the contributions from:

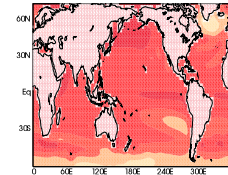
- changes in precipitation vs. evaporative demand
- Changes in mean aridity vs. ENSO variability

Approach

- Identify regions where aridity is historically sensitive to ENSO
 - Find regions where the future changes in mean aridity exceeds the range of ENSO variability
- 6 measures of terrestrial aridity
- $P/(P+E_o)$, CMI, PDSI, 2 SM indices, runoff

Use different sets of experiments to assess the impact of:

1) Ocean warming (+Warming, mean=4K)



2) Plant fertilization to rising CO2 (+VEG)



3) “fast” radiative forcing from enhanced CO2 (+RAD)



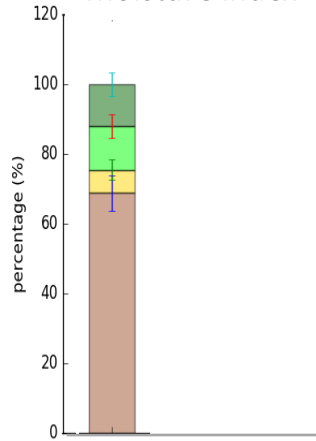
#3 Influence of anthropogenic climate change on regional aridity

■ Always wetter

■ Wetter/Drier with El Niño

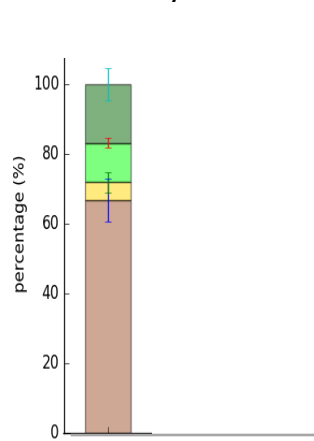
■ Always drier

Climate
Moisture Index



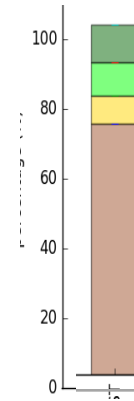
+Warming +Warming
+VEG
+RAD

Palmer Drought
Severity Index



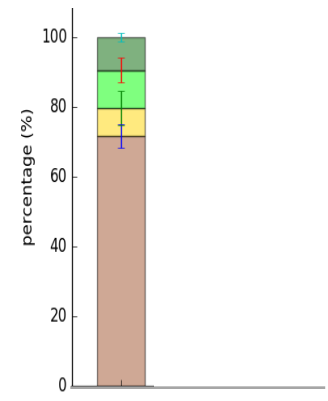
+Warming +Warming
+VEG
+RAD

Surface soil moisture



+Warming +Warming
+VEG
+RAD

Total soil moisture



+Warming +Warming
+VEG
+RAD

Results

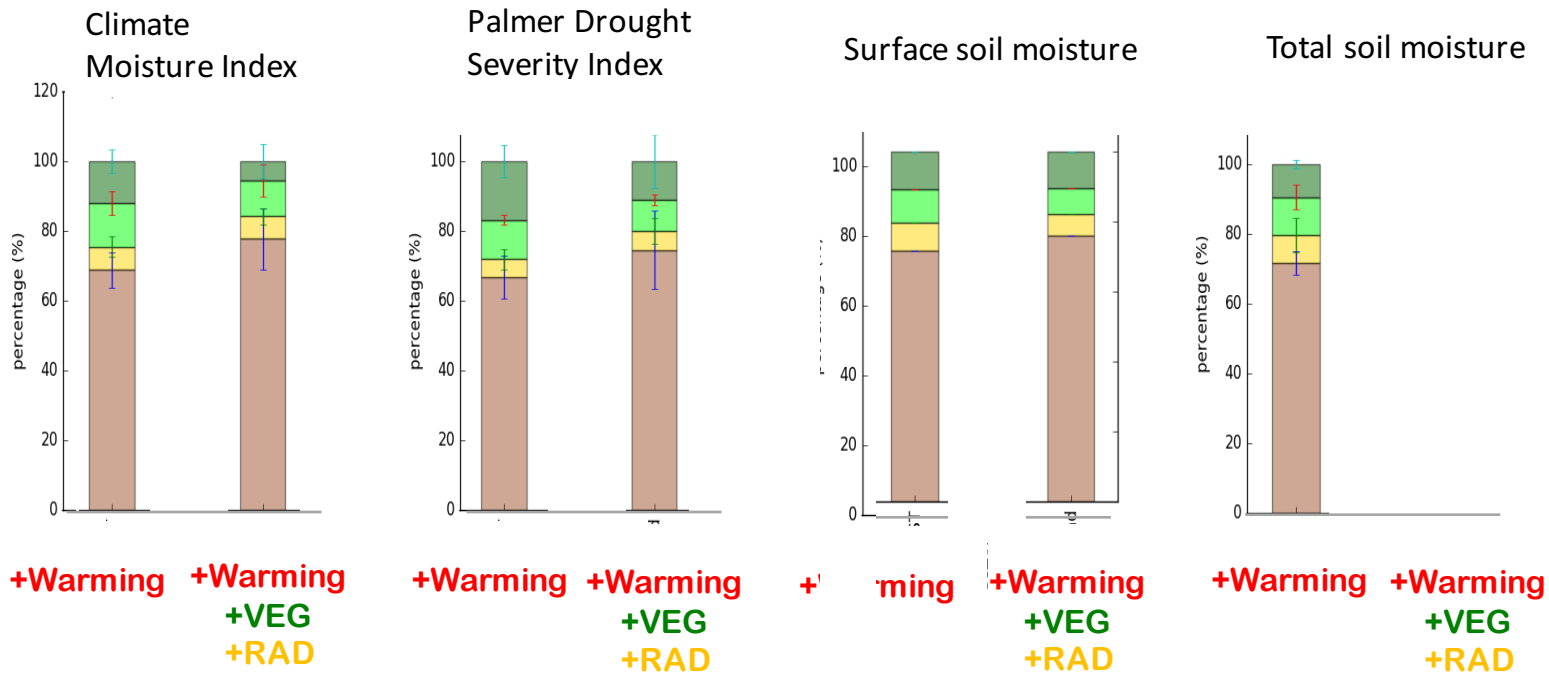
- Future aridity predicted in ~67-72% of the regions where aridity is currently driven by ENSO variability

#3 Influence of anthropogenic climate change on regional aridity

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Results

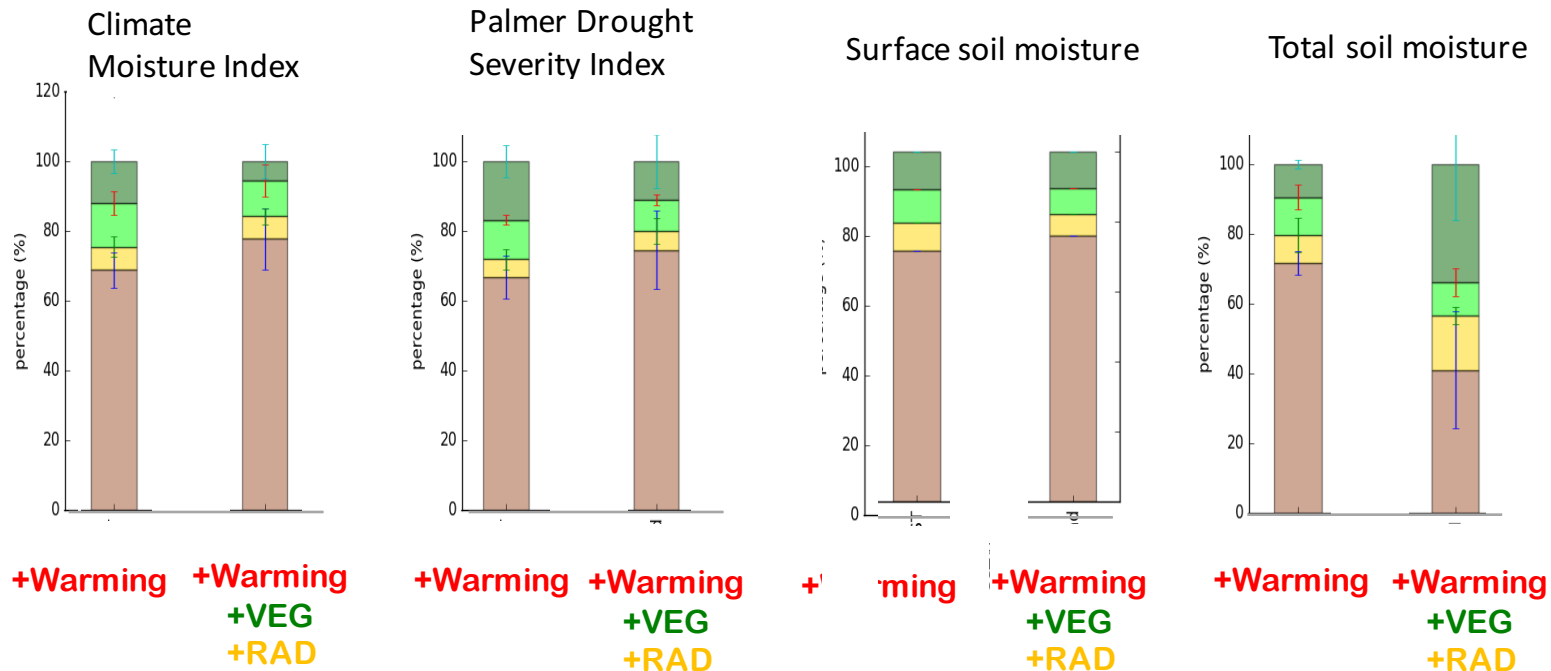
- Future aridity predicted in ~67-72% of the regions where aridity is currently driven by ENSO variability
- It reaches ~75-78% when the vegetation and radiative effects are included

#3 Influence of anthropogenic climate change on regional aridity

■ Always wetter

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■ Always drier

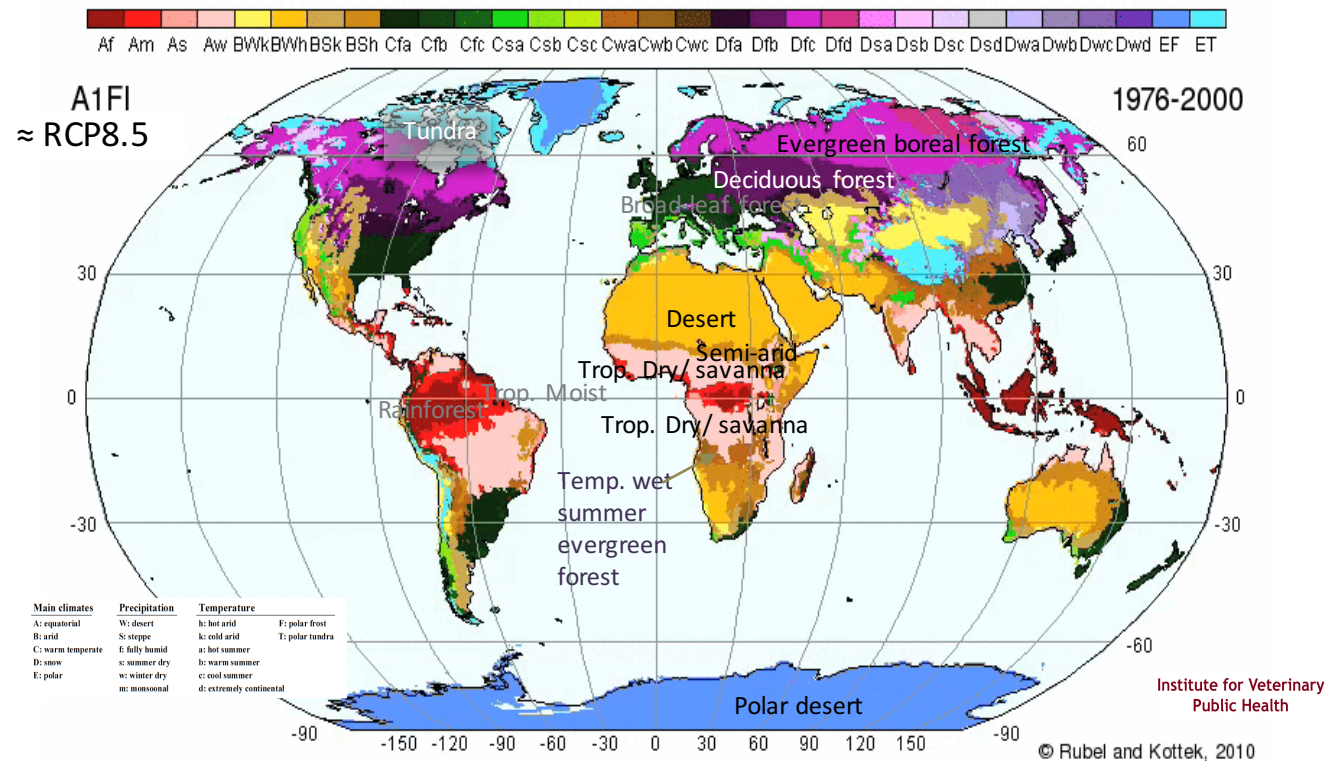


Results

- Future aridity predicted in ~67-72% of the regions where aridity is currently driven by ENSO variability
- It reaches ~75-78% when the vegetation and radiative effects are included
- This prediction is much weaker when total soil moisture is considered (41%): stomatal closure prevents soil desiccation

On the natural and human influences in changes in aridity /droughts?

- **Changes in latitudinal distribution of Köppen vegetation types. Single metric that:**
 - Summarizes the changes in climate that are ecologically relevant
 - Is sensitive to thresholds and features of the seasonal cycle in temperature and precipitation



On the natural and human influences in changes in aridity /droughts?

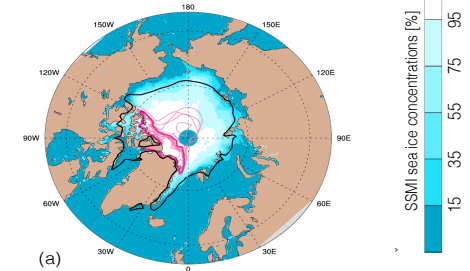
- Changes in latitudinal distribution of Köppen vegetation types.
- Arctic sea ice loss could favor drying in California (Cvijanovic et al 2017)

CESM ensemble of simulations with seasonally ice free Arctic (by sampling model uncertainty in 3 sea-ice physics parameters + initial conditions) -

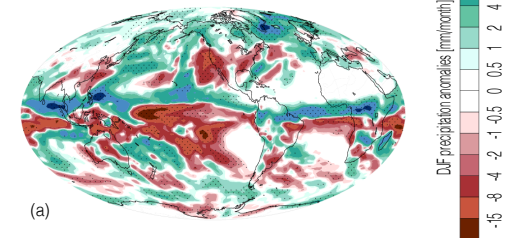
2-step teleconnection:

1. Northward shift in ITCZ
2. Tropical convection reorganization favors a persistent ridge over North Pacific coast

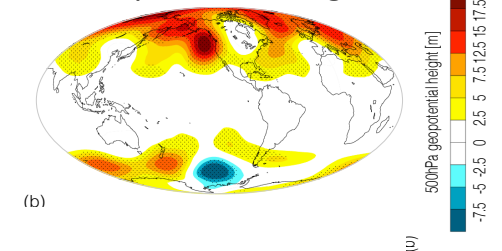
Sea ice (September)



Winter precipitation



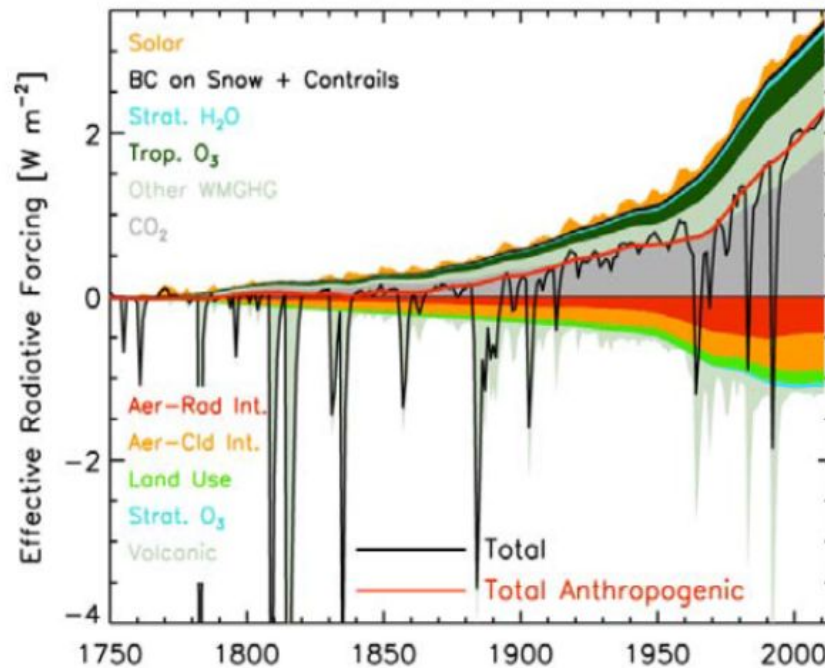
Geopotential height



On the natural and human influences in changes in aridity /droughts?

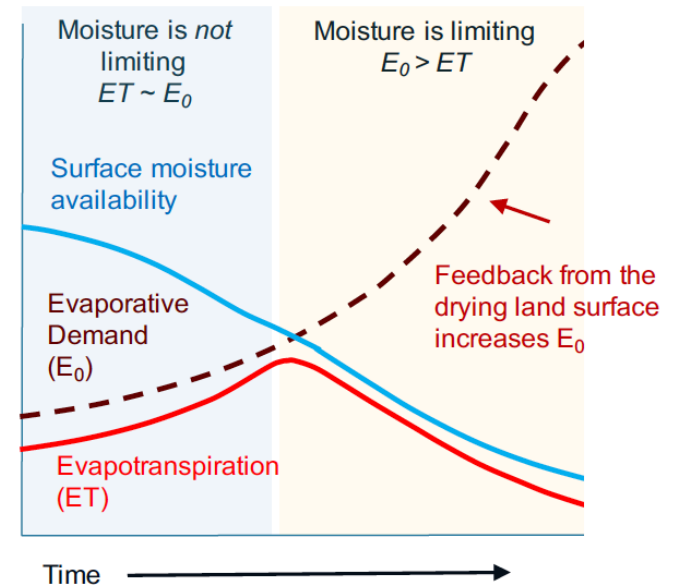
- Changes in latitudinal distribution of Köppen vegetation types.
- Arctic sea ice loss could favor drying in California (Cvijanovic et al 2017)
- D&A considering the impact of GHG, AA, volcanic activity on large-scale changes in P, T & drought index

Variation with Time of Climate Forcings:



Flash droughts

- We find that the characterization of flash droughts can be regionally dependent
- However, within a same region, the characterization of flash droughts could change with time (i.e. forcings).
- A change in SM or E_o mean state could trigger flash droughts that could occur earlier and / or more frequently



Lukas, Hobbins, Rangwala

Large-scale precursors → changes occurring over land, at regional scale

Monthly → daily data

Change in variable → Change in frequency, duration and/or intensity

Ideal coverage period: ideally 1948 – 2017

Need a metric easy to capture in both CMIP simulations and observations

Supplemental Materiel

Another mechanisms: Arctic sea ice loss could favor drying in California

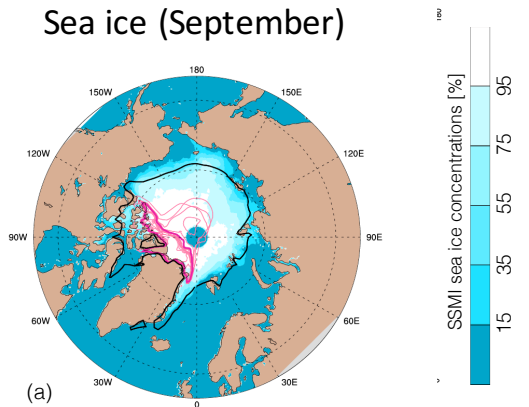
Approach

- CESM ensemble of simulations with seasonally ice free Arctic (by sampling model uncertainty in 3 sea-ice physics parameters + initial conditions)
- Framework allows coupling between sea-ice, ocean and atmosphere in an energetically consistent way

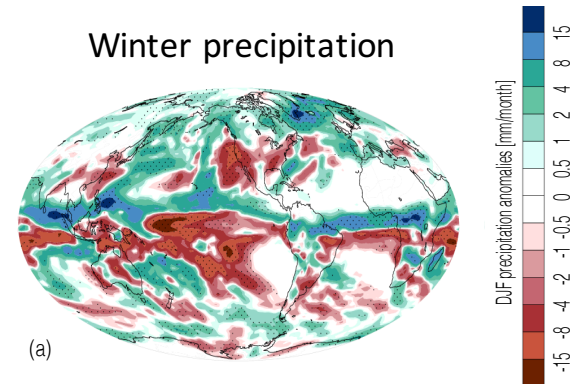
Impact

- ❑ 2-step teleconnection:
 1. Northward shift in ITCZ (Chiang and Bitz 2005)
 2. Tropical convection reorganization favors a persistent ridge over North Pacific coast
- ❑ A misrepresentation of future sea-ice changes has implications for the prediction of future drought risks

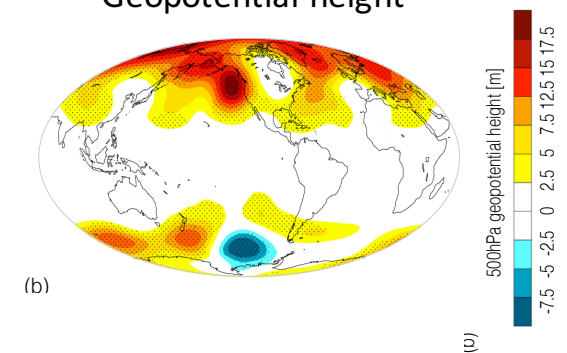
Sea ice (September)

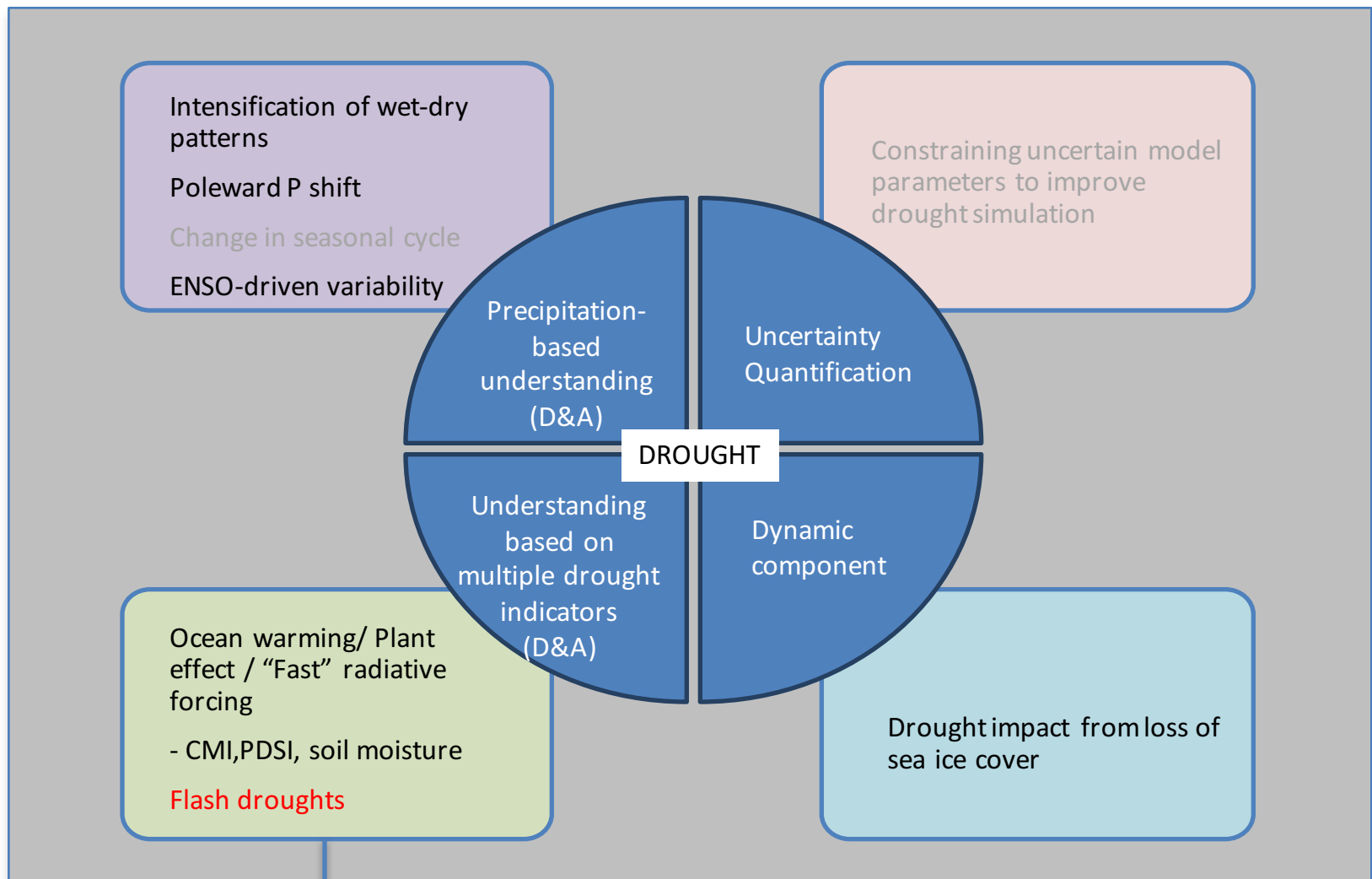


Winter precipitation



Geopotential height





Large-scale precursors of droughts → changes at regional scale
Change in aridity compared to historical mean state → change relative to contemporaneous climatology
Change in variable → Change in frequency, duration and/or intensity
Observations : the same metrics should be easy to compute from model simulations and observations
Opportunities to make progress toward defining adapted flash droughts metrics

#5 Quantifying the effect of parameter uncertainties in simulations of drought in the Western United States

Goal

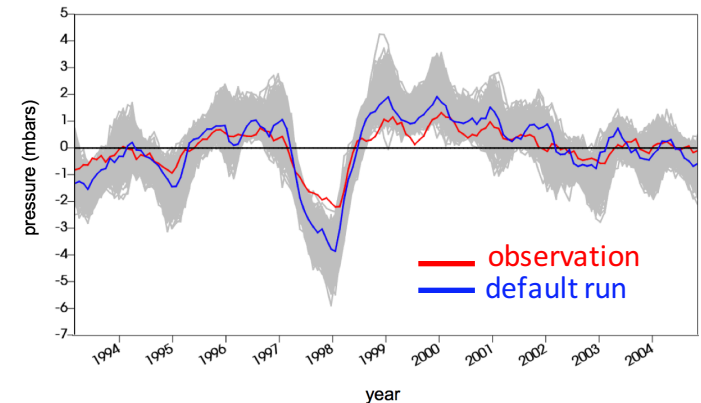
Can we more successfully simulate key features of observed drought behavior?

Approach

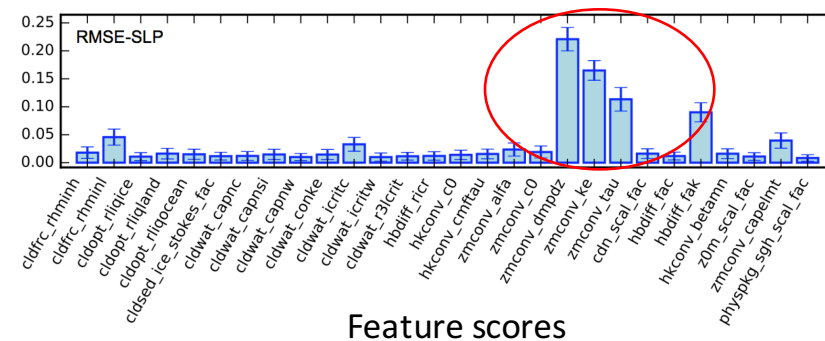
1317-member perturbed physics ensemble

- Latin Hypercube Sampling to vary the values of 28 input parameters (e.g., clouds, P, convection & boundary layer) over allowable ranges
- Set of metrics that best characterize the drought and its drivers (tropical forcing, spatial extent of P and aridity bias)
- We perform a sensitivity analysis to identify the key parameters influencing drought metrics

Forcing metric derived from the difference in pressure anomalies in tropics



Forcing metric is most sensitive to deep convection parameters

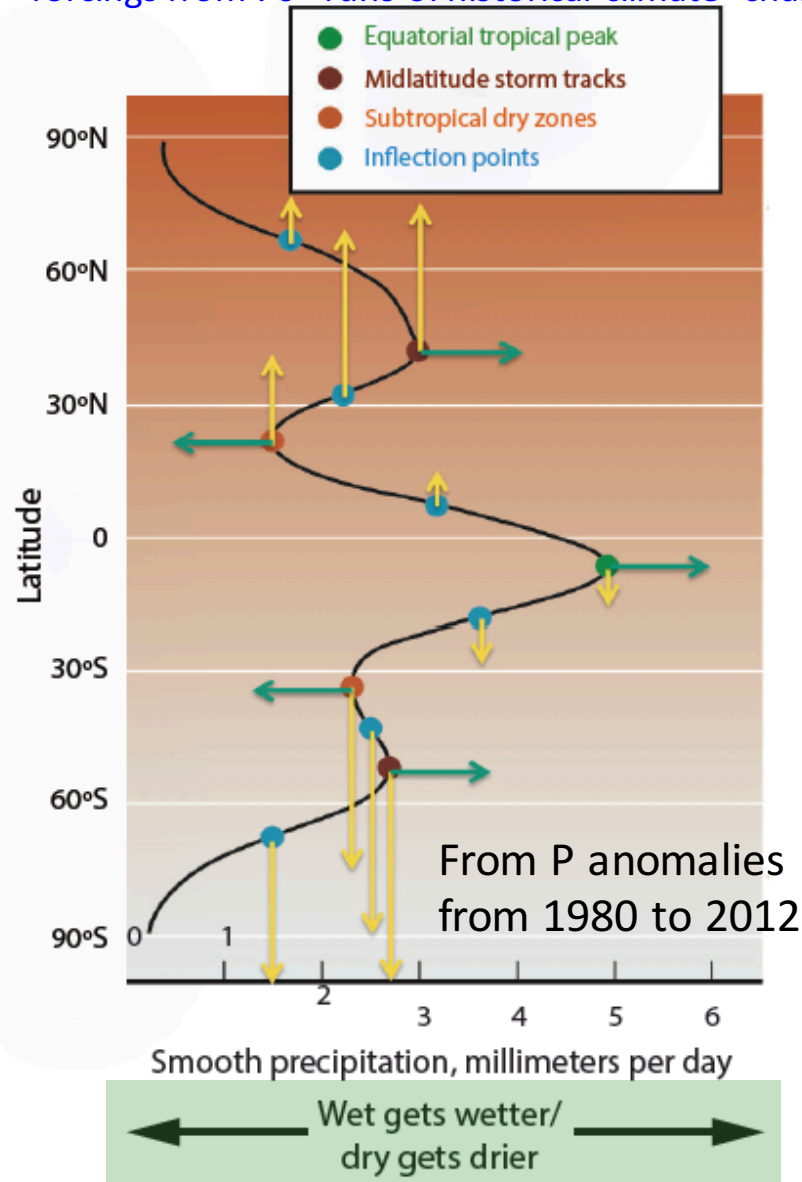


Köppen vegetation map after Ganadesikan and Stouffer (2006)

Köppen Regional Climate (Generic Vegetation Type)	T/P Criteria*
Ef: Polar Desert (scant vegetation)	$T_{\max} < 0$ Celsius (C)
Et: Tundra (dwarf trees, mosses)	$0\text{ C} < T_{\max} < 10\text{ C}$ and $T_{\min} < -3\text{ C}$
Dc: Cold Winters/Cool Summers (evergreen boreal forest)	$T_{\min} < -3\text{ C}$ and < 4 months warmer than 10 C , but not types BS or BW
Dab: Cold Winters/Warm Summers (deciduous forest)	$T_{\min} < -3\text{ C}$, $T_{\max} > 10\text{ C}$ and > 4 months warmer than 10 C , but not types BS or BW
Cw: Temperate, Wet Summers (evergreen forest)	$-3\text{ C} < T_{\min} < 18\text{ C}$ and $P_{\max} > 10P_{\min}$ with P_{\max} occurring in summer and P_{\min} in winter, but not types BS or BW
Cs: Temperate, Wet Winters (evergreen broad-leaf forest)	$-3\text{ C} < T_{\min} < -18\text{ C}$ and $P_{\max} > 3P_{\min}$, with P_{\max} occurring in winter and P_{\min} in summer, but not types BS or BW
Cfc: Temperate, Cool and Moist (needle-tree forest)	$-3\text{ C} < T_{\min} < 18$ and $T_{\max} < 22\text{ C}$, and with < 4 months warmer than 10 C , but not types BS , BW , Cs , or Cw
Cfb: Temperate, Warm and Moist (broad-leaf forest)	$-3\text{ C} < T_{\min} < 18$ and $T_{\max} < 22\text{ C}$, and with > 4 months warmer than 10 C , but not types BS , BW , Cs , or Cw
Cfa: Temperate, Hot and Moist (broad-leaf forest)	$-3\text{ C} < T_{\min} < 18\text{ C}$ and $T_{\max} > 22\text{ C}$, but not types BS , BW , Cs , or Cw
BS: Semiarid (bush or grassland)	$(T_{\text{avg}} + P_{\text{off}}) < P_{\text{year}} < 2(T_{\text{avg}} + P_{\text{off}})$
BW: Desert (wasteland, cactus/seasonal Vegetation)	$P_{\text{year}} < (T_{\text{avg}} + P_{\text{off}})$
Af: Tropical Wet (tropical evergreen rain forest)	$T_{\min} > 18\text{ C}$ and $P_{\min} > 6\text{ cm}$, but not types BS or BW
Am: Tropical Moist (tropical evergreen forest)	$T_{\min} > 18\text{ C}$ and $(250\text{ cm} - P_{\text{year}})/25 < P_{\min} < 6\text{ cm}$, but not types BS or BW
Aw: Tropical Dry (savanna/woodland)	$T_{\min} > 18\text{ C}$ and $P_{\min} < 6\text{ cm}$, $(250\text{ cm} - P_{\text{year}})/25$, but not types BS or BW

#1 Natural and Human Influences on Changing Zonal-Mean Precipitation

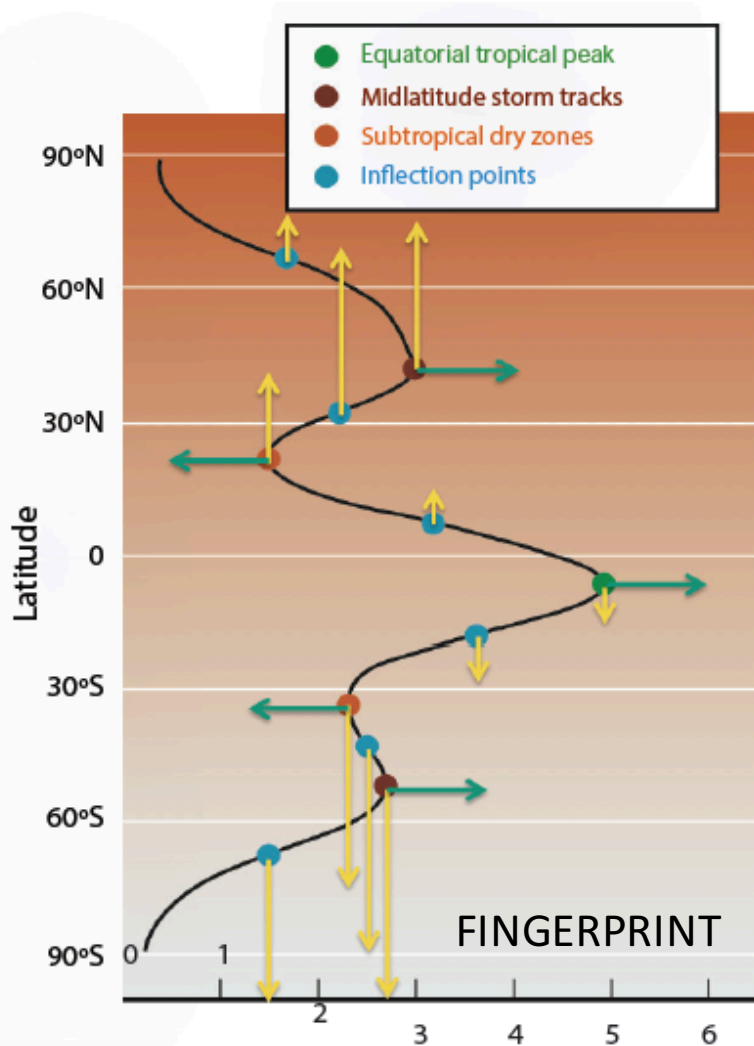
Human fingerprint = expected response to human forcings from 70+ runs of historical climate change



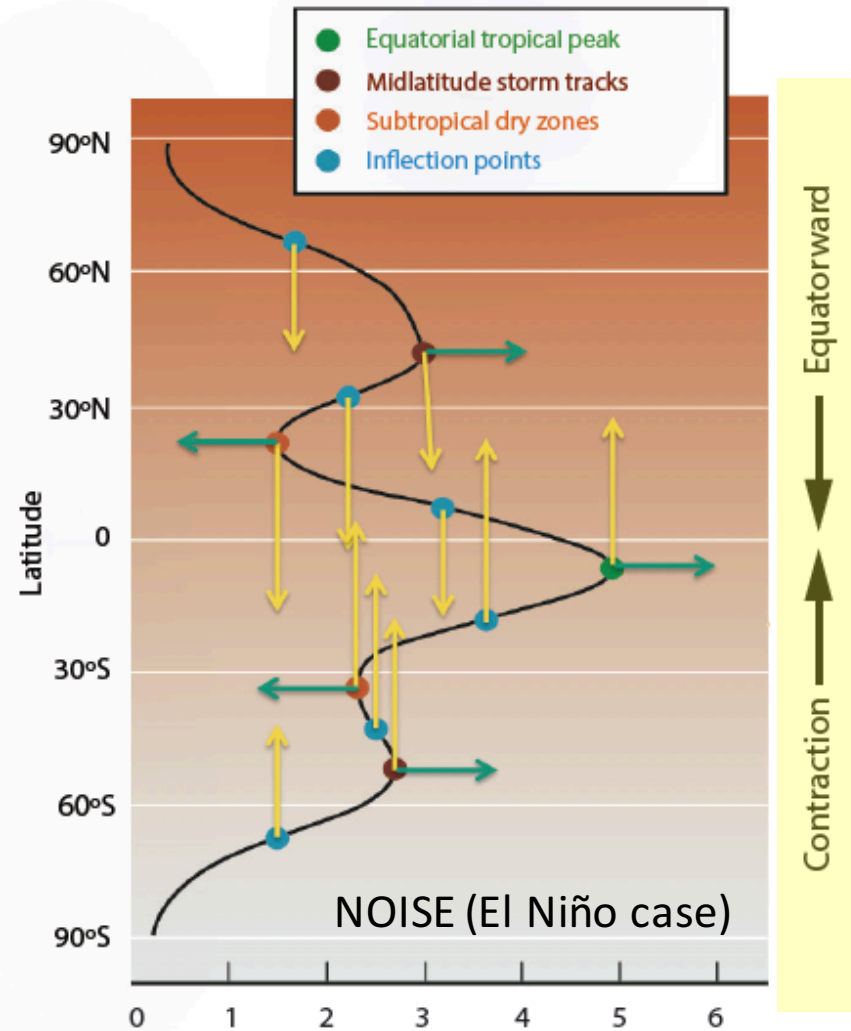
Dynamic: poleward expansion (x50)

FINGERPRINT
= Distortion of the curve in response to human forcing

#1 Natural and Human Influences on Changing Zonal-Mean Precipitation



Dynamic: poleward expansion (x50)



Equatorward
Contraction

Human pattern \neq Noise pattern

The intensification + poleward expansion of zonal P is found in observations and cannot be explained by climate noise alone