

**Federal Service for Hydrometeorology
and Environmental Monitoring**



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Since 1849



Information that global modelers need from regional and sub-grid scale

Vladimir Kattsov

Workshop “NEA land surface properties and change and its role in the global earth system and models”, AGCI, Aspen, 12-17 August 2007

Projection credibility aspects

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- ✓ Sensitivity to external forcing
- ✓ Systematic errors

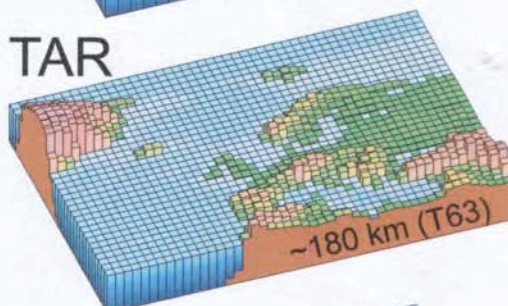
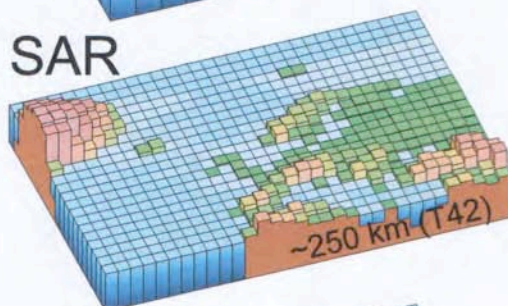
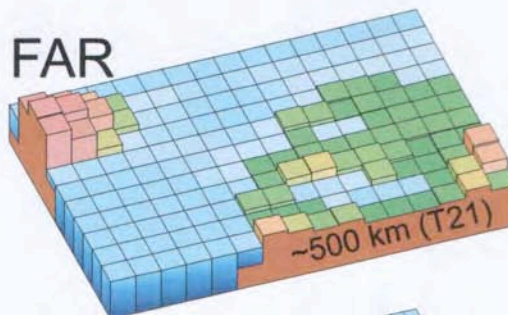
Priorities

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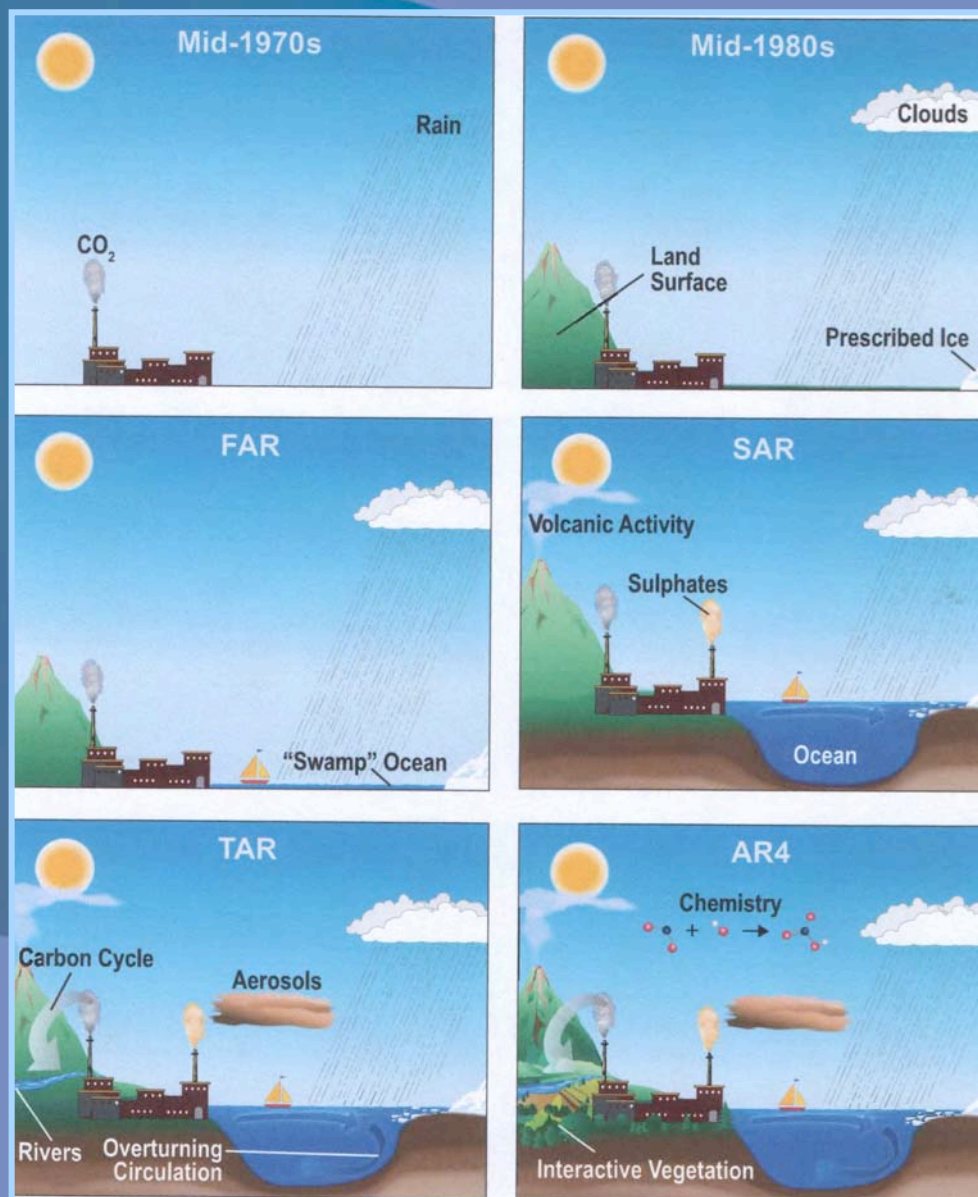


- ✓ Higher resolution
- ✓ More components
- ✓ Larger ensembles

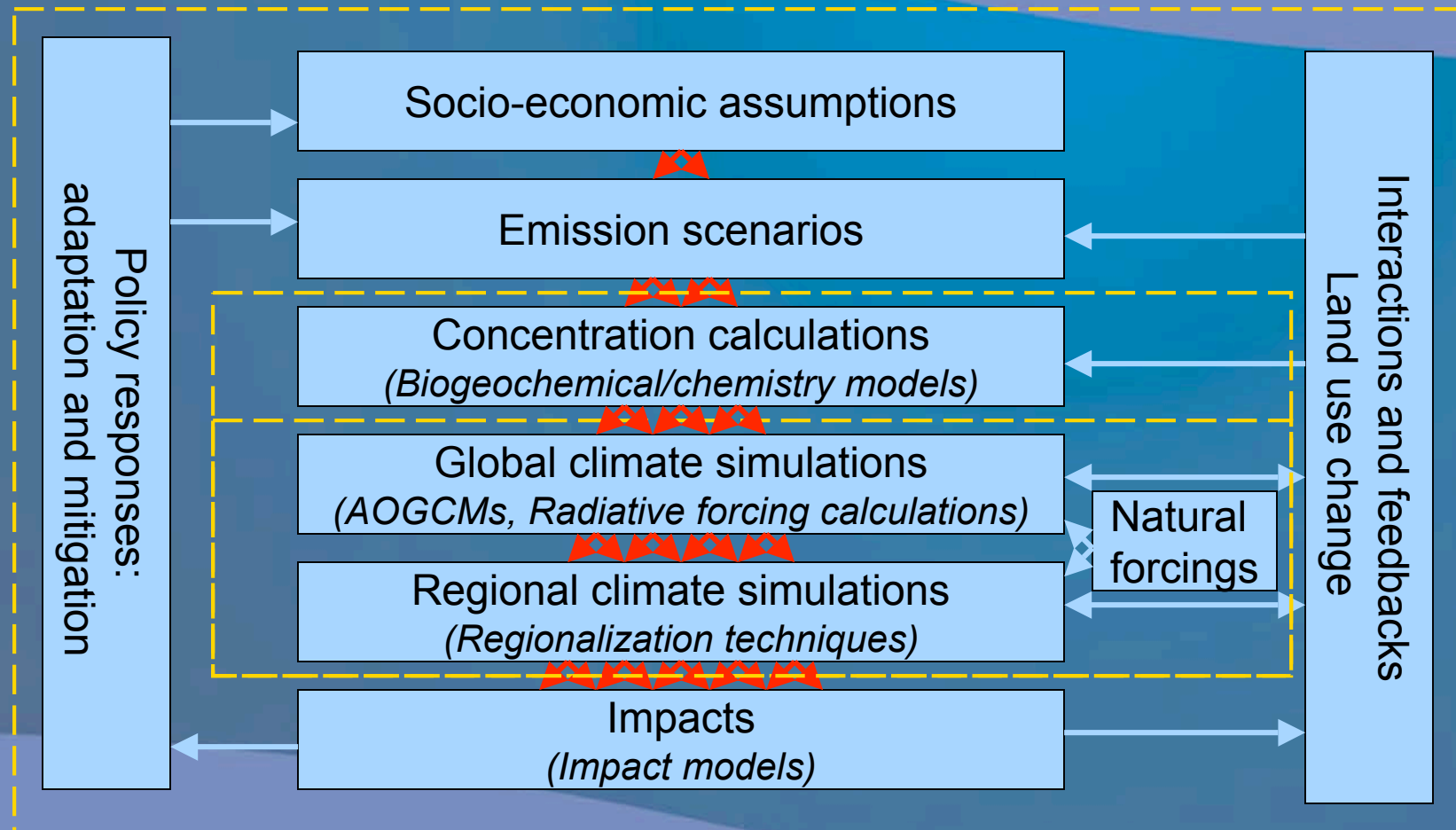
Climate model resolution



Climate model components



Projection uncertainty cascade



WCRP CMIP3 multi-model dataset

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Model name	Country
BCCR-BCM2.0	Norway
BCC-CM1	China
CCSM3	USA
CGCM3.1(T47)	Canada
CGCM3.1(T63)	Canada
CNRM-CM3	France
CSIRO-Mk3.0	Australia
FGOALS-g1.0	China
ECHAM5/MPI-OM	Germany
ECHO-G	Germany/Korea
GFDL-CM2.0	USA
GFDL-CM2.1	USA
GISS-AOM	USA
GISS-EH	USA
GISS-ER	USA
INM-CM3.0	Russia
IPSL-CM4	France
MIROC3.2 (hires)	Japan
MIROC3.2 (medres)	Japan
MRI-CGCM2.3.2	Japan
PCM	USA
UKMO-HadCM3	UK
UKMO-HadGEM1	UK

Unprecedented model experiment:
23 models, 16 centers, 11 countries, 32 Tb,
1000+ scientists.



<http://www.pcmdi.llnl.gov>

New generation of AOGCMs

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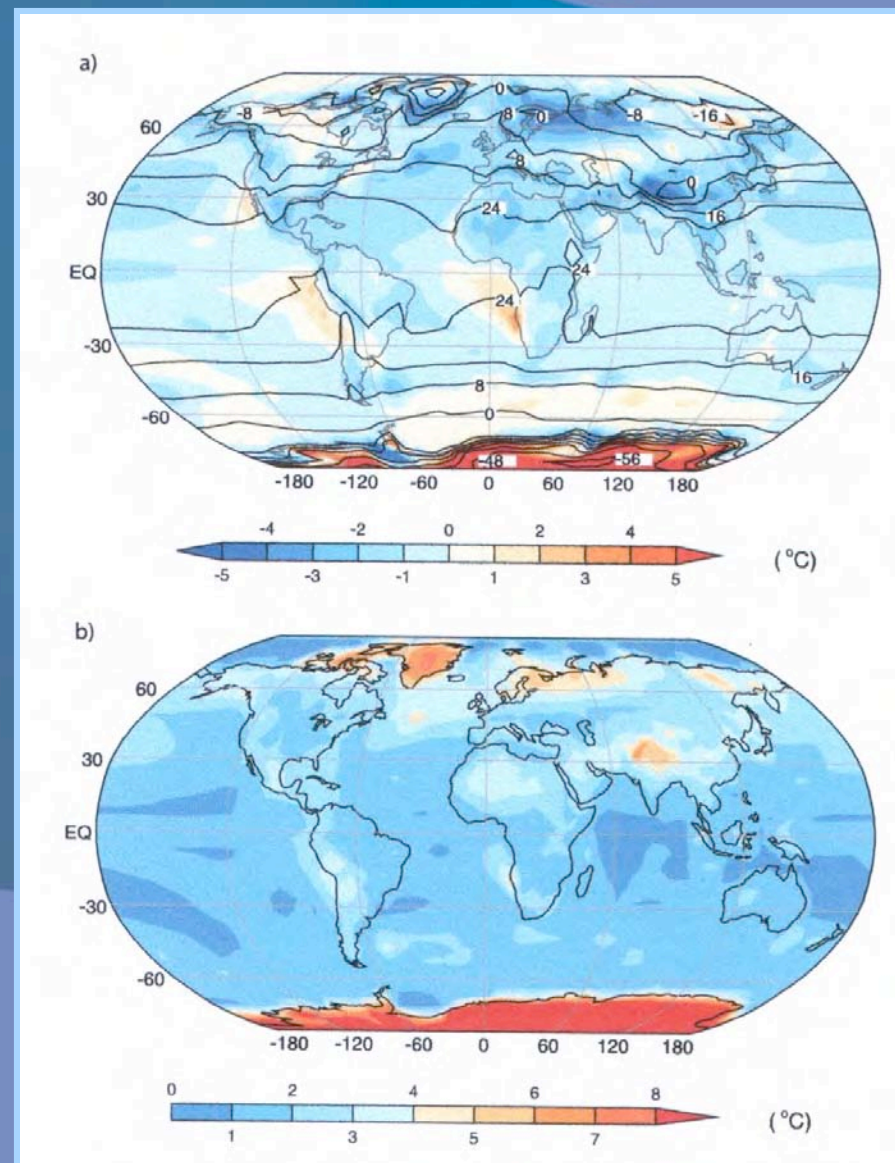


Resolution
Grids
Components
Flux correction
Modes of natural variability
Extremes
Southern Ocean
Sensitivity
Dynamic sea ice

Temperature

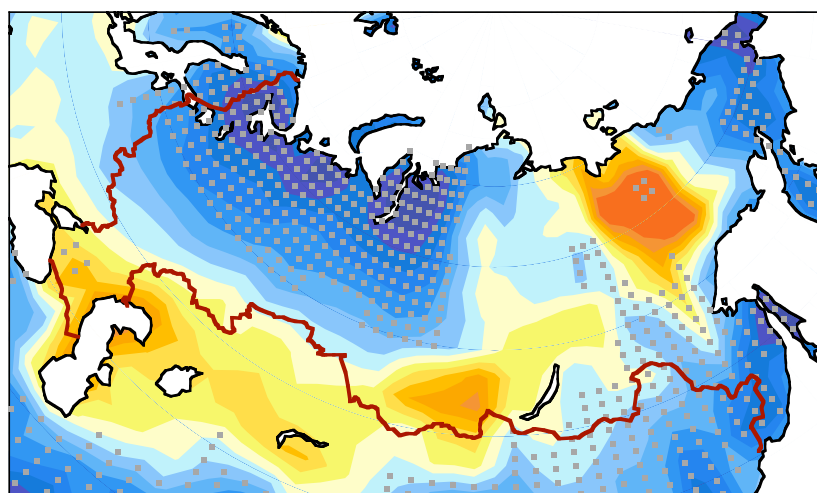
bias

rmse



Temperature

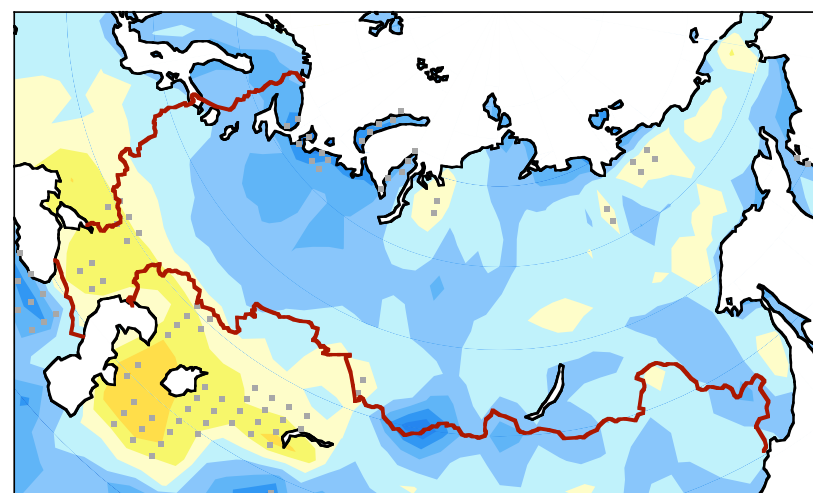
tas C IPCC AR4
DJF 16m Ensemble mean S1 - CRU CLIMATE A2 1980-1999
fild min= -12.8704 max= 9.69561
■ >90% sign_agree



20.07.2007 VG

DJF

tas C IPCC AR4
JJA 16m Ensemble mean S1 - CRU CLIMATE A2 1980-1999
fild min= -11.4618 max= 5.1216
■ >90% sign_agree



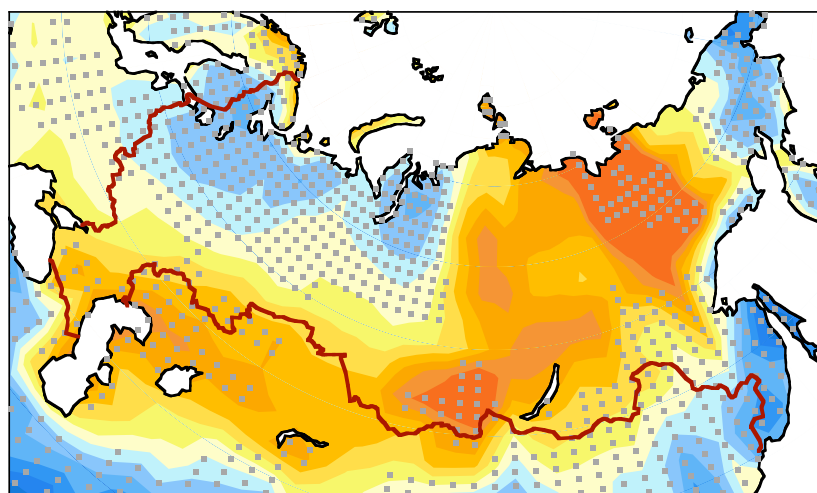
20.07.2007 VG

JJA

Temperature (stand-alone AGCMs)

tas C AMP2
DJF 16m Ensemble mean S1 - CRU CLIMATE 1980-1999
fild min= -10.4061 max= 13.1574

■ >66% sign_agree

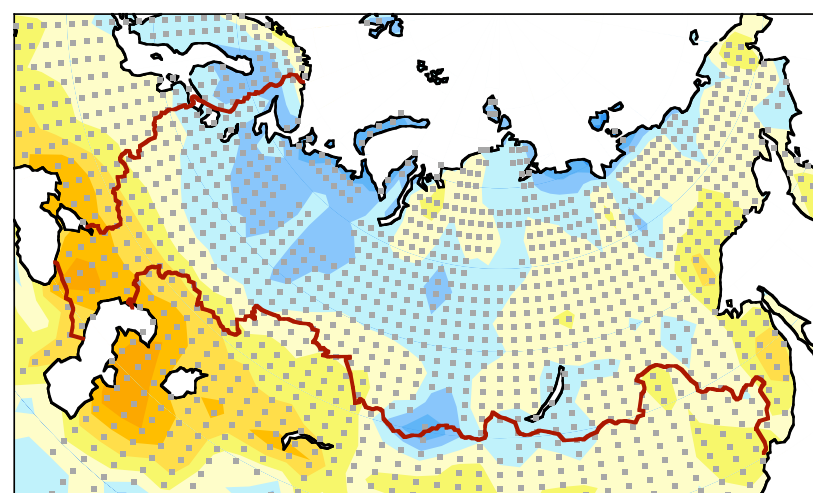


29.06.2007 VG

DJF

tas C AMP2
JJA 16m Ensemble mean S1 - CRU CLIMATE 1980-1999
fild min= -7.56765 max= 6.61344

■ >66% sign_agree

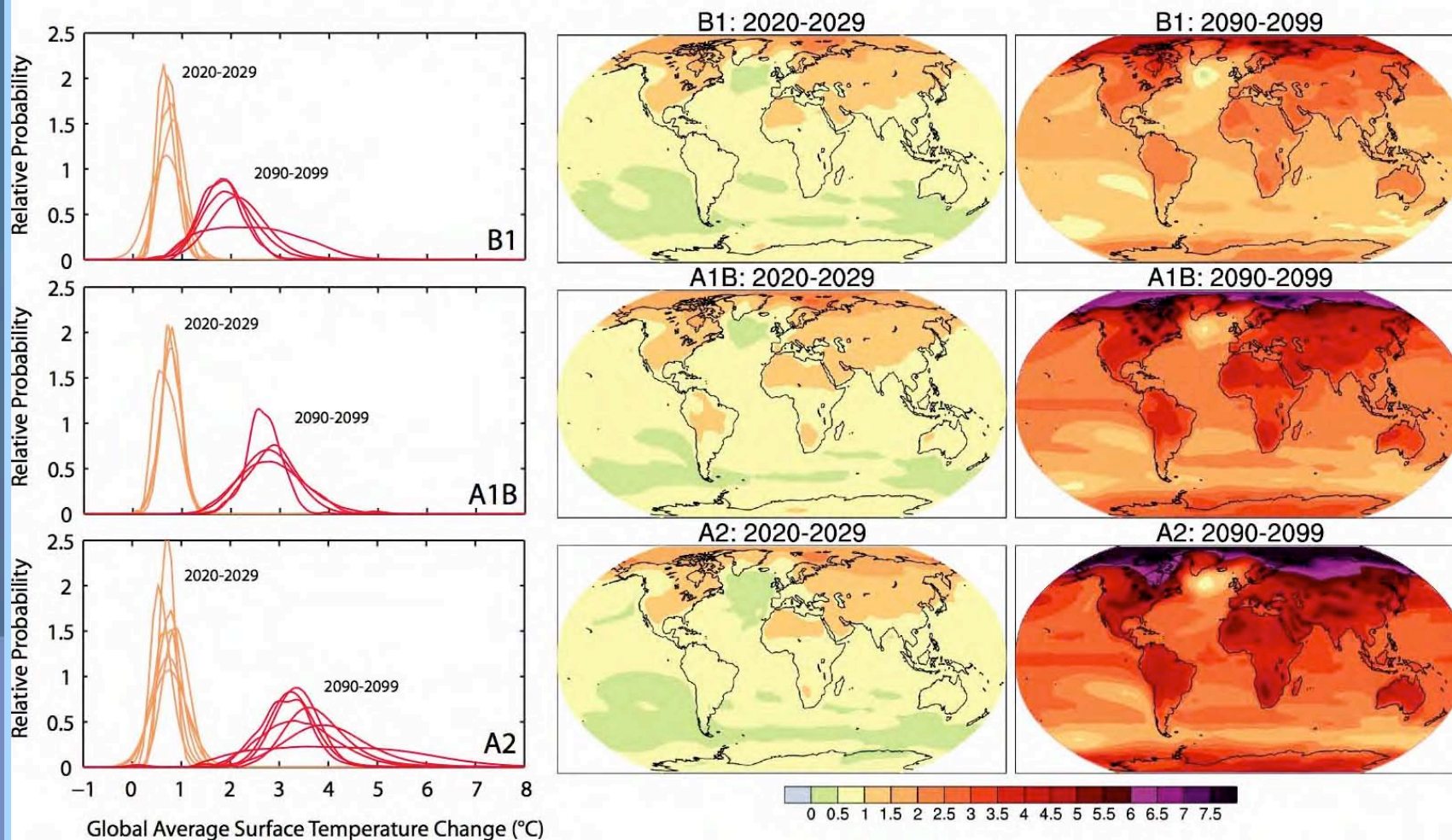


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JJA

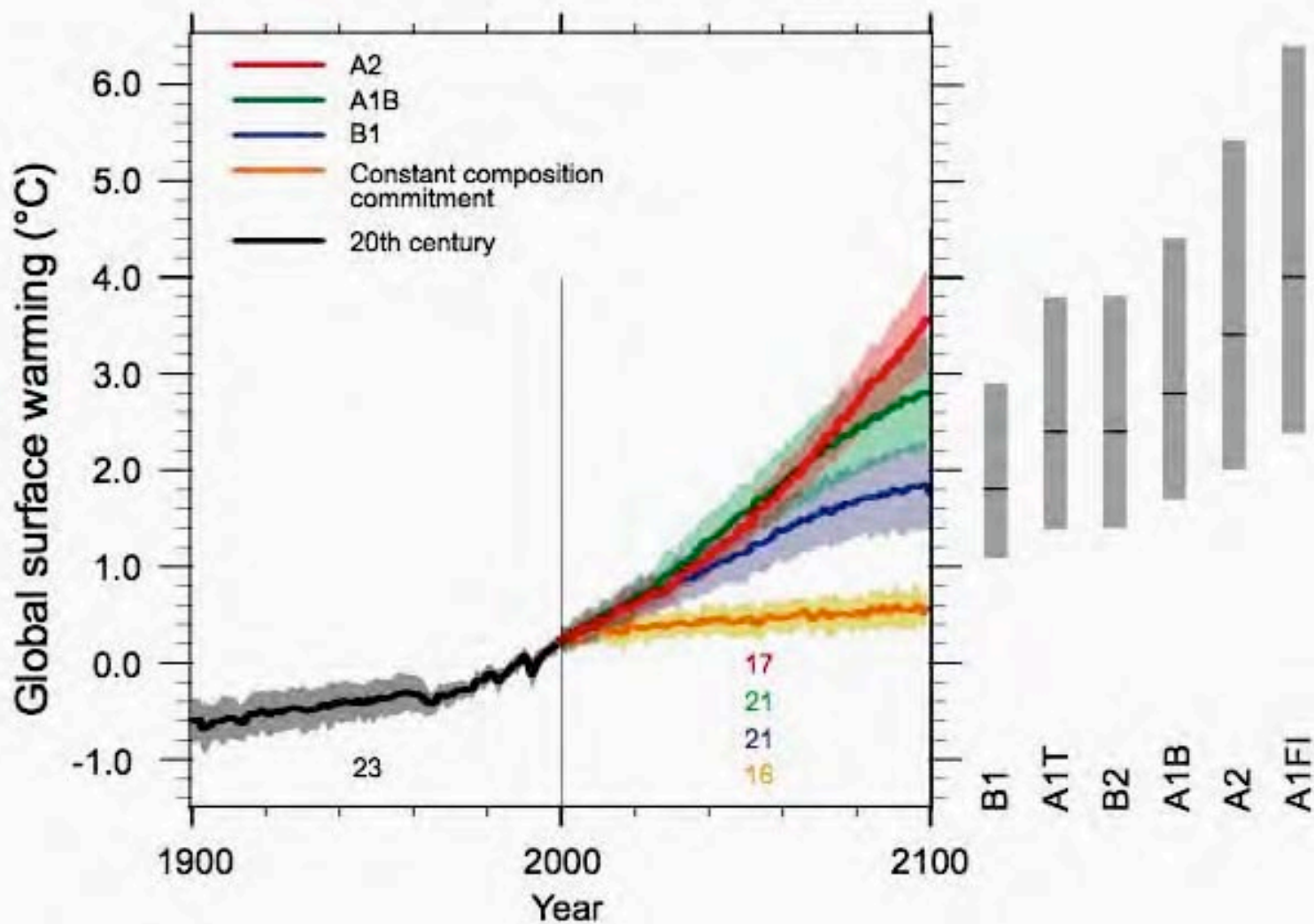
Temperature

AOGCM Projections of Surface Temperatures



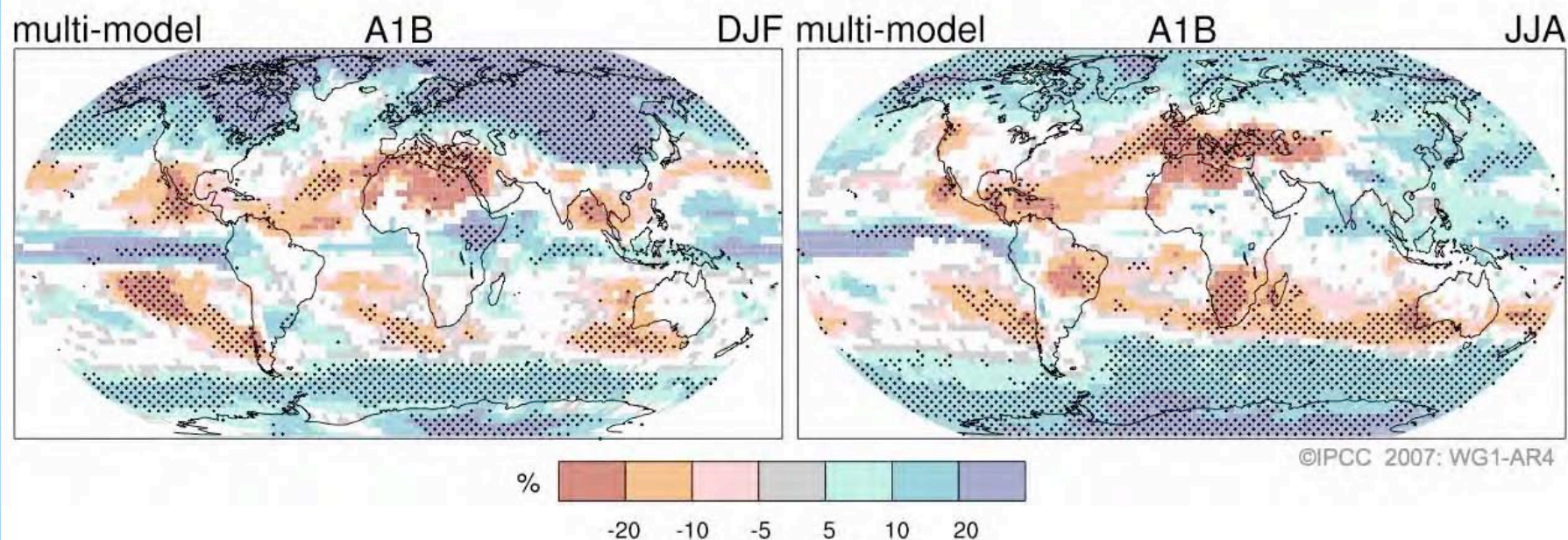
©IPCC 2007: WG1-AR4

Projections



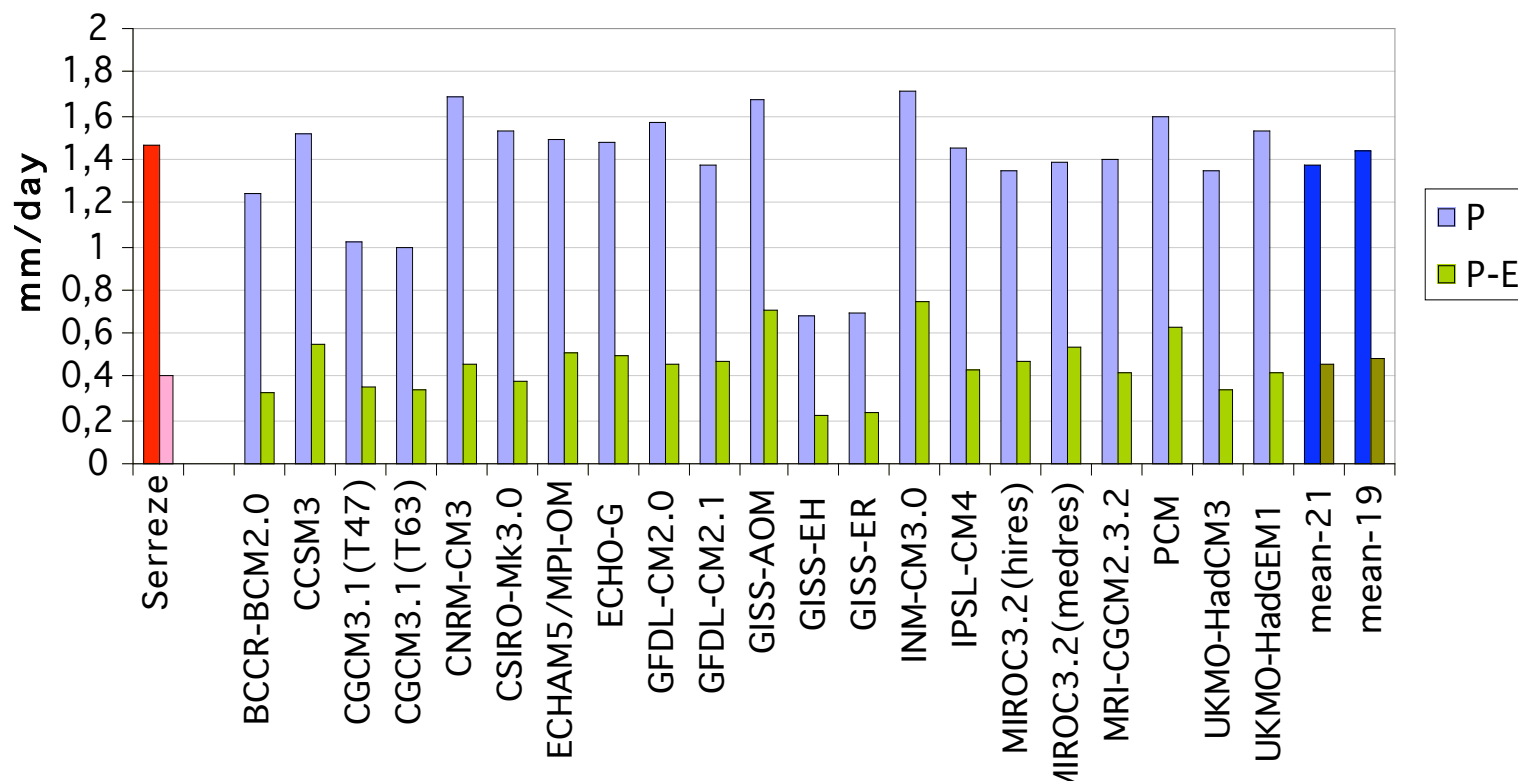


Projected Patterns of Precipitation Changes

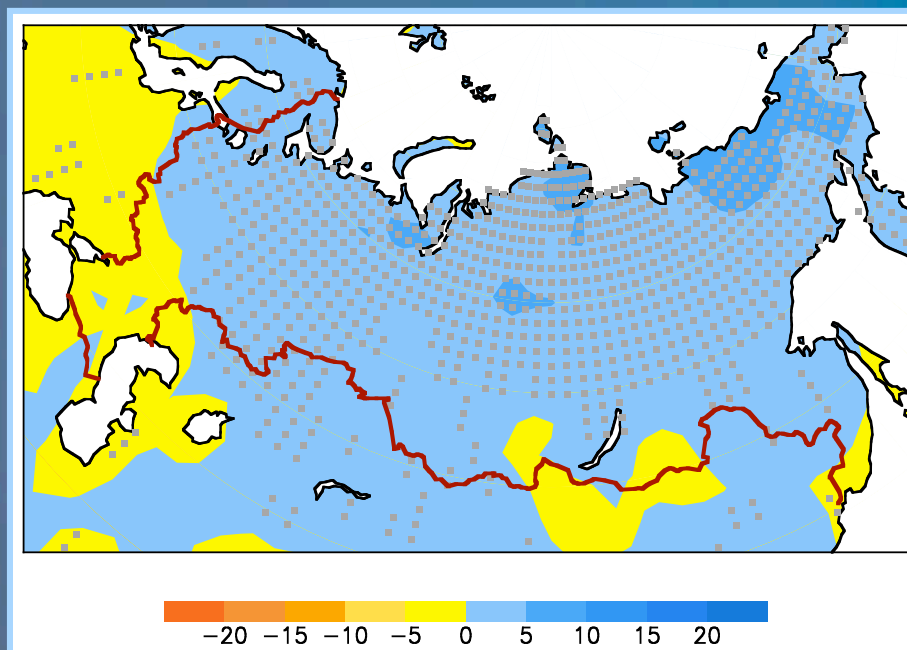


Ob watershed: annual mean precipitation and P-E

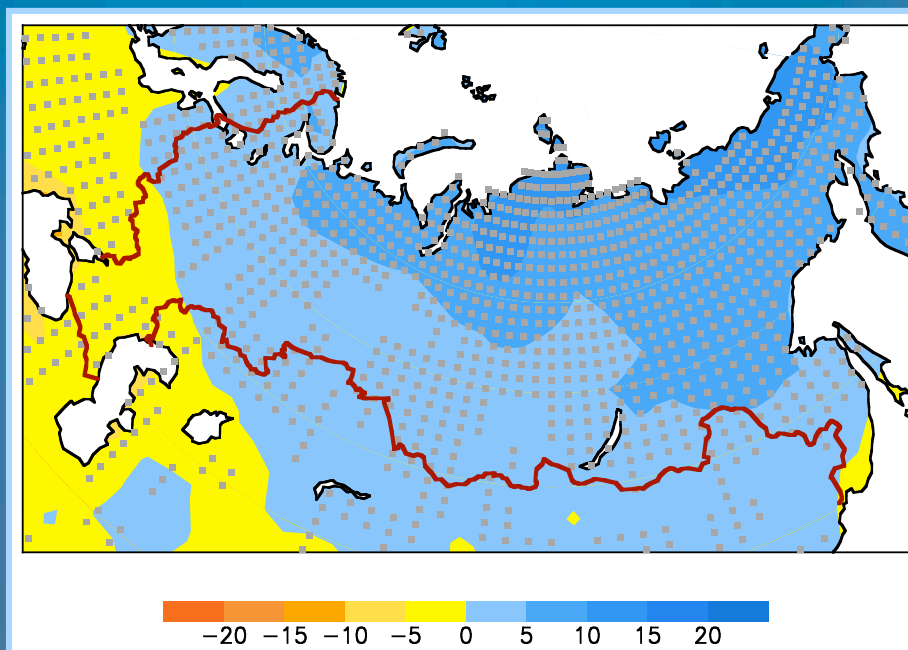
Ob: P and P-E annual means (1960-1989)



Runoff change (%)

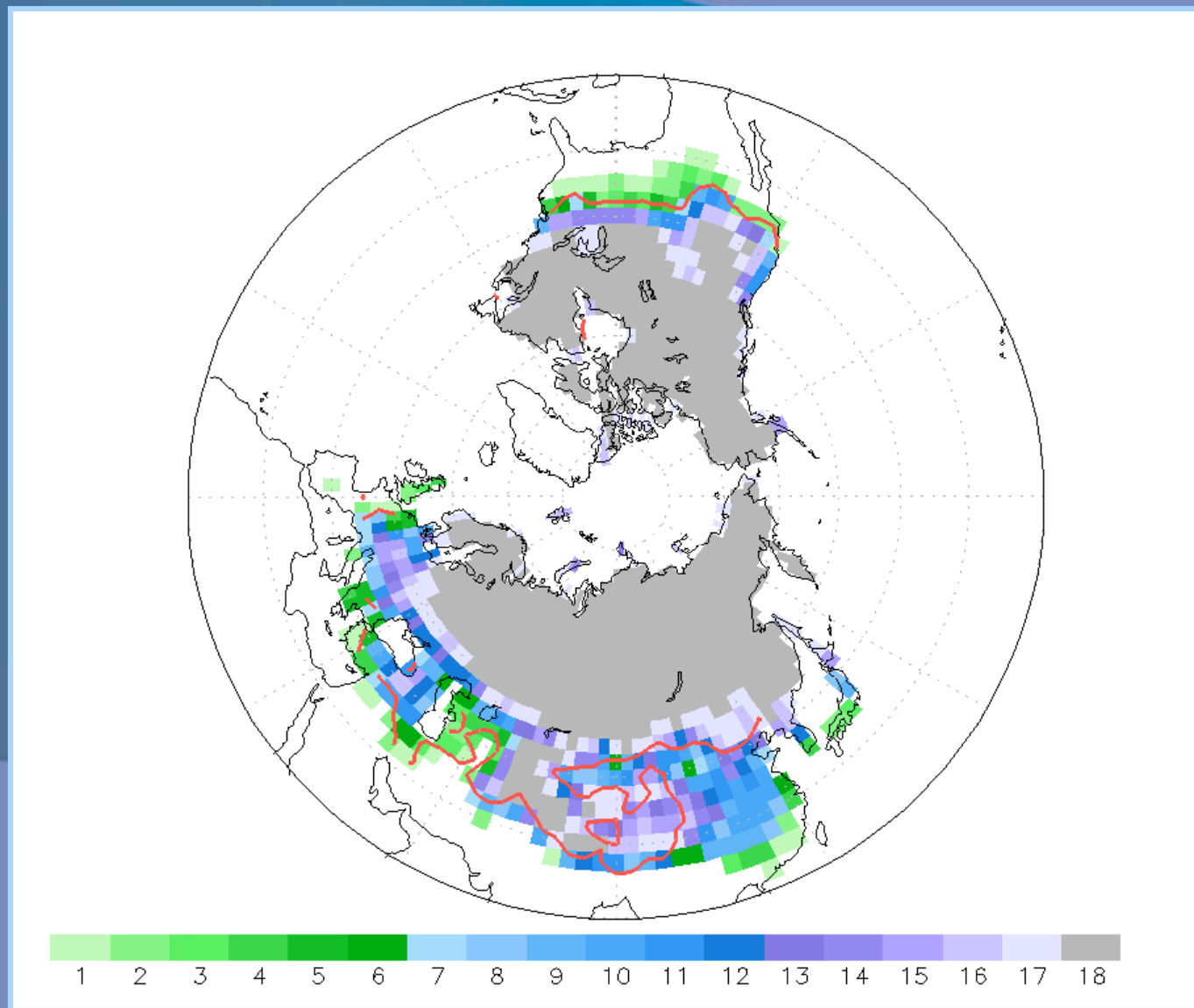


2011-2030



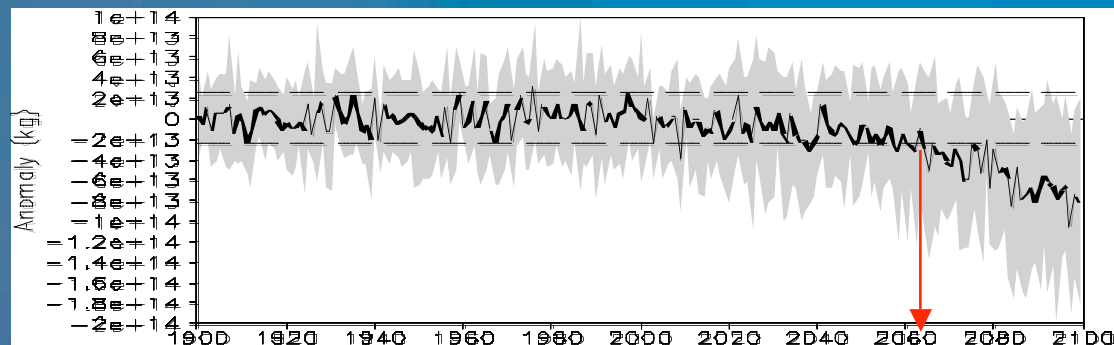
2041-2060

Snow

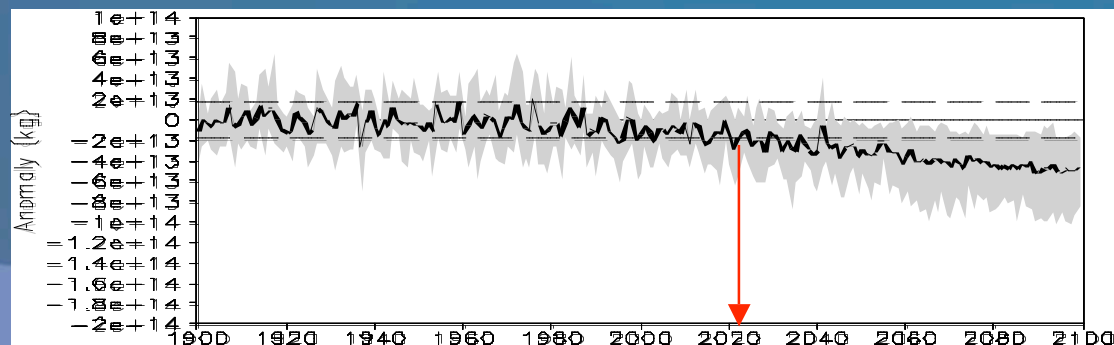


The Ob catchment: Snow mass anomaly (kg) evolution as simulated by the ensemble of IPCC AOGCMs (_2)

March

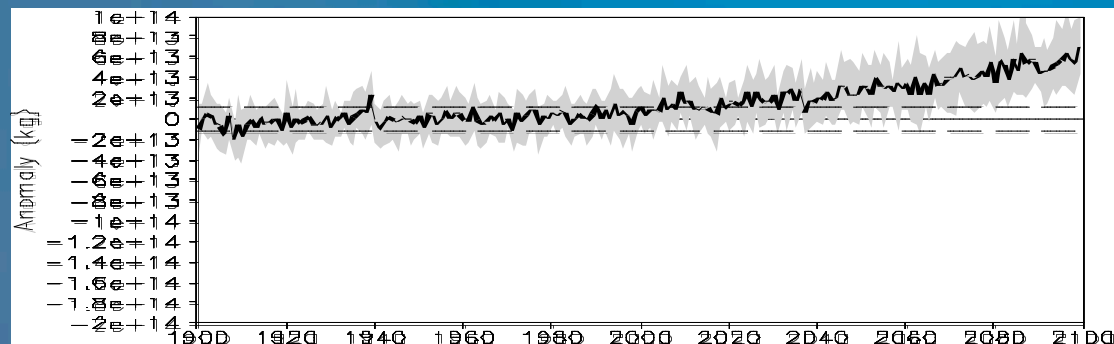


May

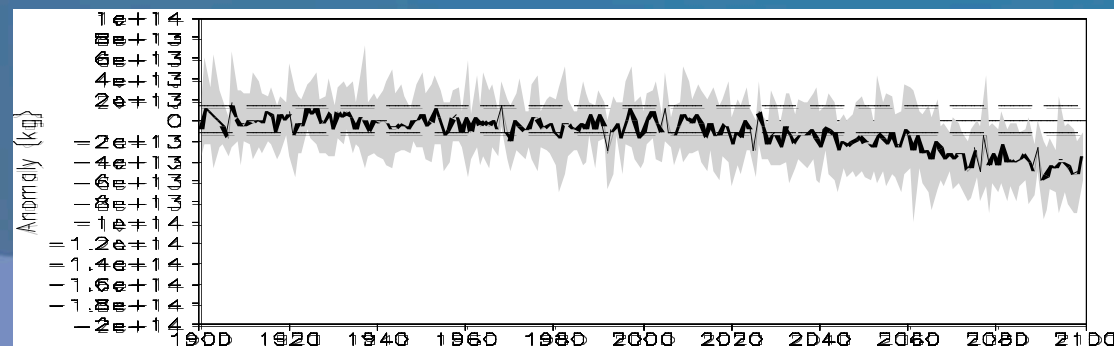


The Lena catchment: Snow mass anomaly (kg) evolution as simulated by the ensemble of IPCC AOGCMs (_2)

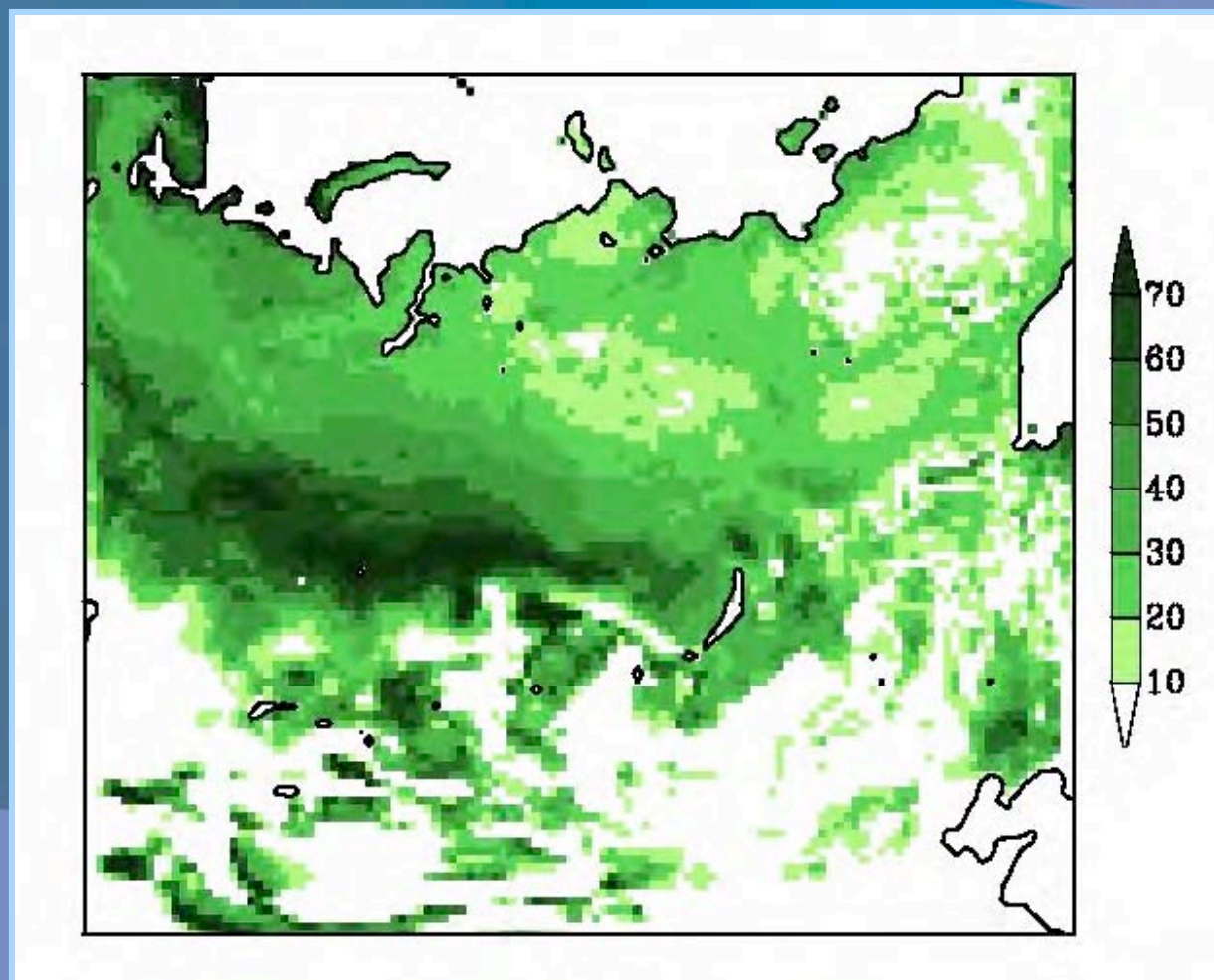
March



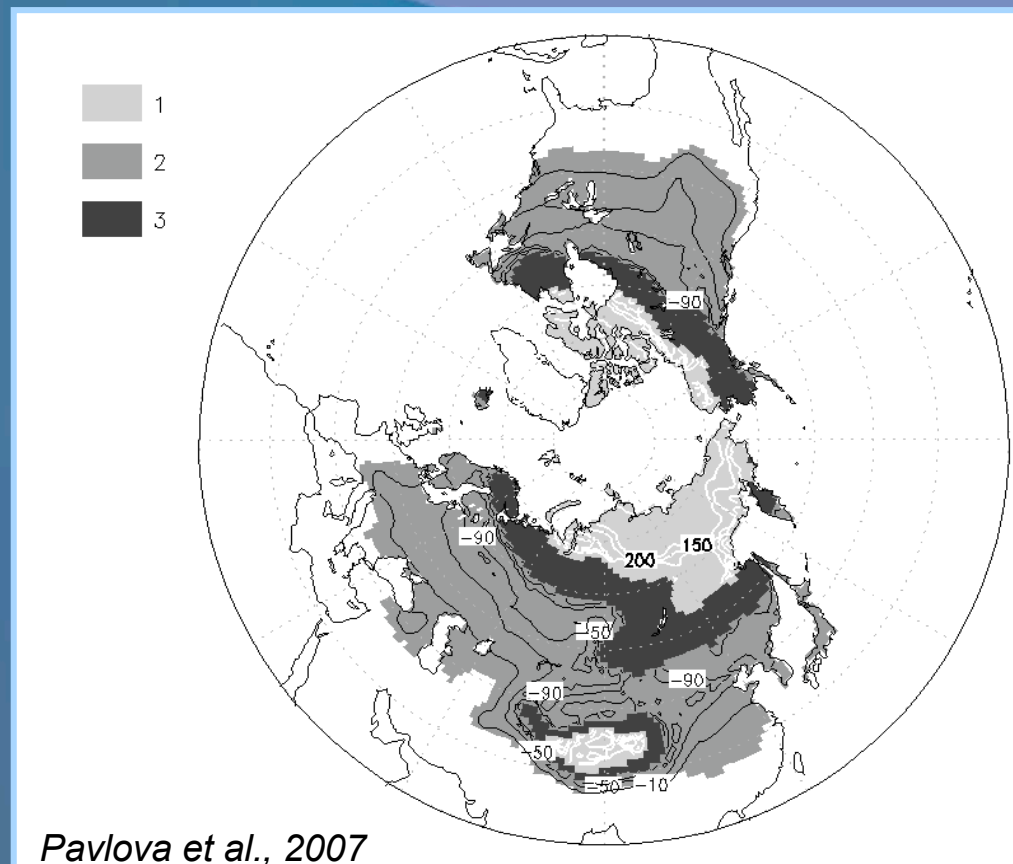
May



Change of stable snow cover duration by the end of the 21st century (days, A2)



Permafrost



Depth changes of seasonally thawing layer (white lines, cm) and seasonally frozen layer (black lines, cm) in 2080-2099 (A2) relative to 1980-1999.

- 1- region of seasonal thawing retained by the end of the 21st century,
- 2- same but for seasonal freezing,
- 3- region where seasonal thawing is replaced by seasonal freezing in the upper 3 m layer/



Among IPCC AR4 key uncertainties:

Changes in key processes that drive some global and regional climate changes are poorly known (e.g., ENSO, NAO, blocking, MOC, **land surface feedbacks**, tropical cyclone distribution). {11.2–11.9}

The magnitude of future **carbon cycle feedbacks** is still poorly determined. {7.3, 10.4}

Some priorities



- ✓ Effects of vegetation dynamics and interaction with land-surface on NEA energy and water cycles (e.g. snow+vegetation=albedo+water flux seasonality)
- ✓ Effects of cryospheric and vegetation changes on the chemical composition of the atmosphere (e.g. GHG from thawing permafrost or black carbon from forest fires)
- ✓ Land surface mosaic and upscaling techniques

Shift of climatic zones in the warming climate

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We acknowledge the modeling groups for making their simulations available for analysis, the Program for Climate Model Diagnosis and Intercomparison (PCMDI) for collecting and archiving the CMIP3 model output, and the WCRP's Working Group on Coupled Modelling (WGCM) for organizing the model data analysis activity. The WCRP CMIP3 multi-model dataset is supported by the Office of Science, U.S. Department of Energy. Some of the results presented above were obtained under the projects supported by the US National Science Foundation via the International Arctic Research Center of the University of Alaska Fairbanks (subaward UAF05-0074 of OPP-0327664) and by the Russian Foundation for Basic Research (Grant 05-05-65093).



All models that treat the coupling of the carbon cycle to climate change indicate a positive feedback effect with warming acting to suppress land and ocean uptake of CO₂, leading to larger atmospheric CO₂ increases and greater climate change for a given emissions scenario, but the strength of this feedback effect varies markedly among models.



Future concentrations of many non-CO₂ greenhouse gases and their precursors are expected to be coupled to future climate change. Insufficient understanding of the causes of recent variations in the CH₄ growth rate suggests large uncertainties in future projections for this gas in particular.

Future emissions of many aerosols and their precursors are expected to be affected by climate change.



IPCC AR4 robust findings:

Natural processes of CO₂ uptake by the oceans and terrestrial biosphere remove about 50 to 60% of anthropogenic emissions (i.e., fossil CO₂ emissions and land use change flux). Uptake by the oceans and the terrestrial biosphere are similar in magnitude over recent decades but that by the terrestrial biosphere is more variable. {7.3}

Future warming would tend to reduce the capacity of the Earth system (land and ocean) to absorb anthropogenic CO₂. As a result, an increasingly large fraction of anthropogenic CO₂ would stay in the atmosphere under a warmer climate. This feedback requires reductions in the cumulative emissions consistent with stabilisation at a given atmospheric CO₂ level compared to the hypothetical case of no such feedback. The higher the stabilisation scenario, the larger the amount of climate change and the larger the required reductions. {7.3, 10.4}