

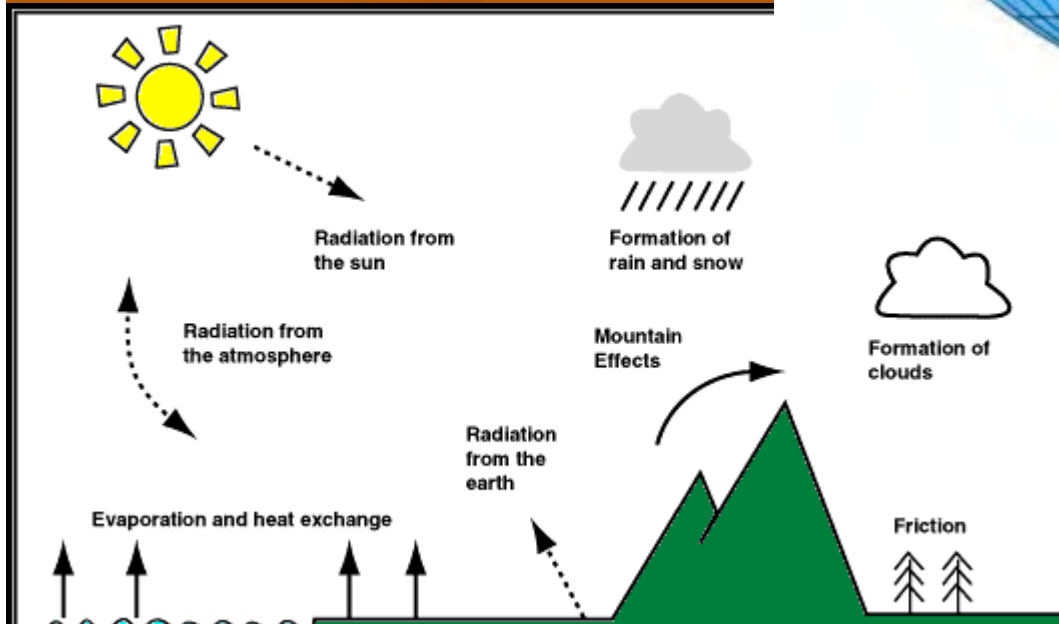
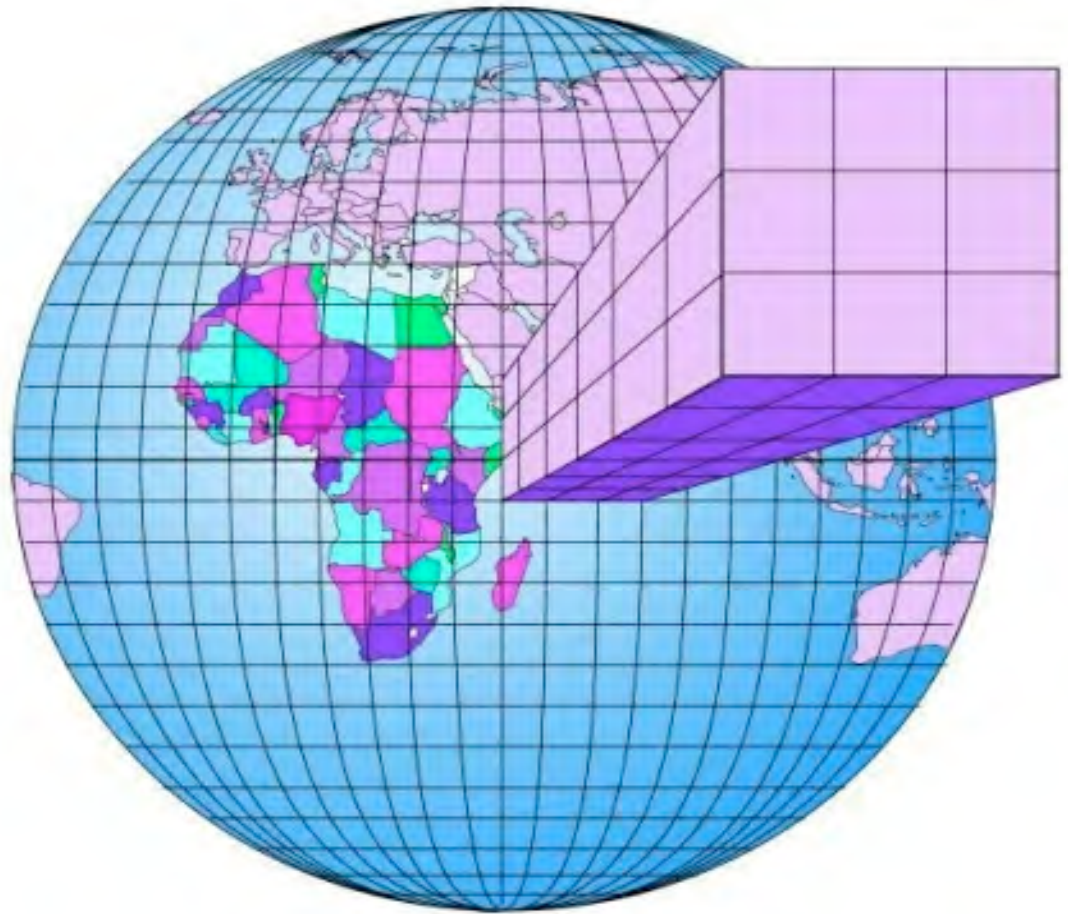
Land Surface Modeling in the Extra-tropics

Robert E. Dickinson

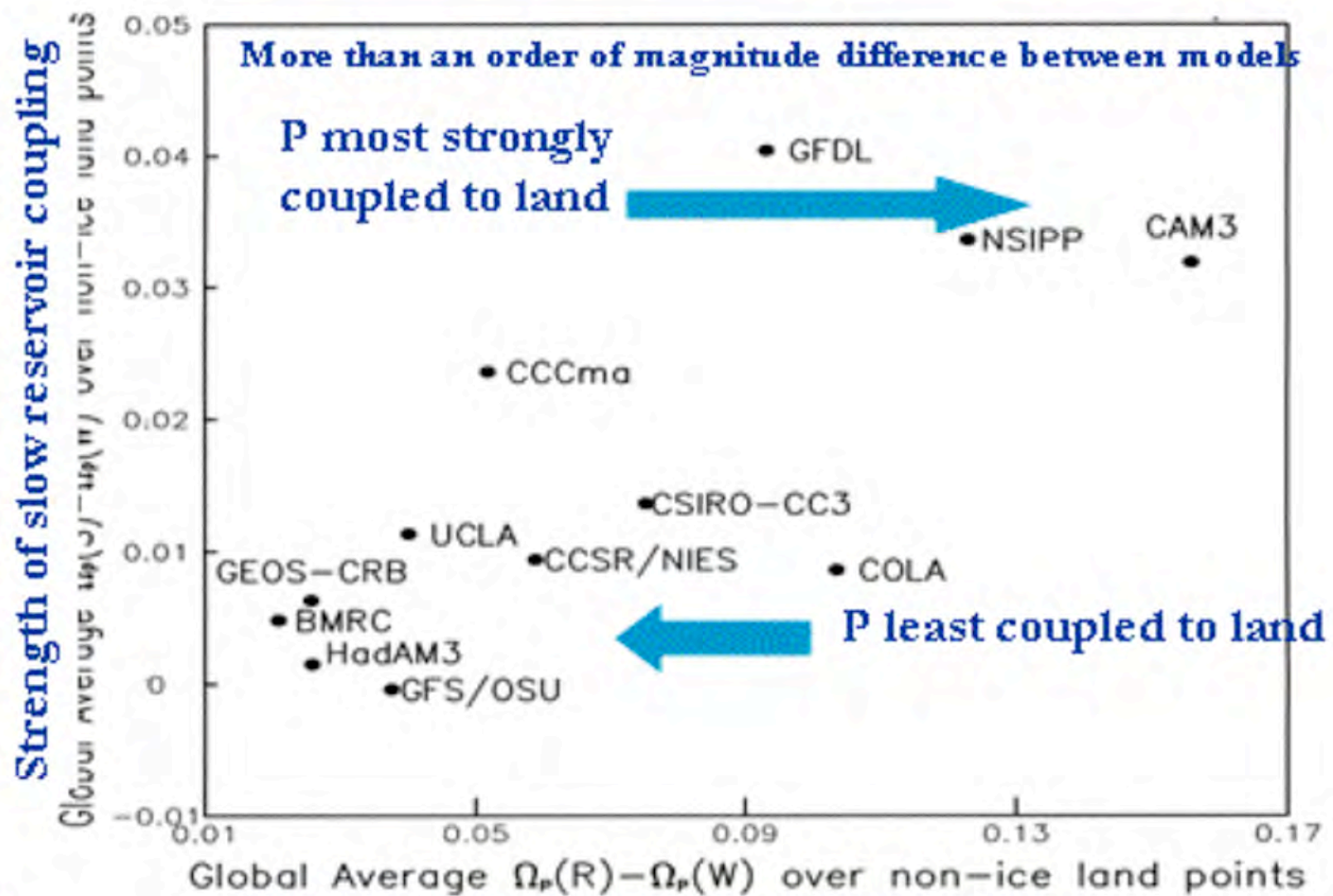
Aspen Global Change Institute

14 August, 2007

Climate Model



Scaling/parameterization
Need to describe details
within the grid boxes



1

2 Fig. 7 Global average strength of all reservoir coupling for all

3 twelve models.

4

Outline

- Processes to be modeled
- What's right and wrong with current state of the art?
- Emphasis on: a) canopy- soil or snow interaction; b) treating complexity of land surface.

Processes to be modeled?

- Basic Physical Constraints
 - Energy Conservation
 - Water Conservation
- What comes in and goes out?
- What interacts with and is stored at land surface?
- Where can substantial errors be made?

Solar –highly variable input

- Geometry
 - Latitude
 - Time of year
 - Diurnal
- Cloud properties
- Aerosols
- Atmospheric Composition (H_2O , O_3 , CO_2)

Interaction of solar input with “land surface”

- Land surface is a complex system of surfaces
 - Canopy consisting of leaves, stems and their geometric properties
 - Soil surface –can be overlain by snow
 - Other surfaces – water, ice
- Each of these surfaces has its own energy and water balance.
- How much to include-what to lump together?

Lumping together

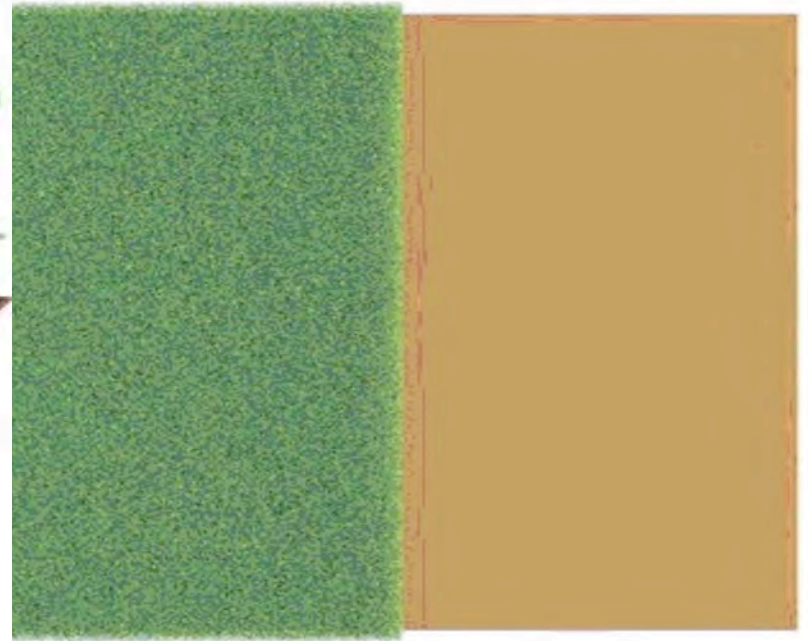
- Simplifies but leads to error – how much?
- Albedo describes the total system reflected solar –
 - models now attempt to construct from individual elements
 - leaves, canopy and their geometries
 - -large compared to wavelengths so logic is mostly that of ray-tracing between objects with partial absorptions when a ray hits an object
 - leaves assigned single scattering albedos.

How now done?

- Describing mutual shading of leaves a key element.
- Leaves treated as a plane parallel scattering cloud – this adds up to canopy albedo – combines with albedo of surface under canopy by adding principle.
- Soil or other underlying surface not vertically under canopy is treated as radiatively separate



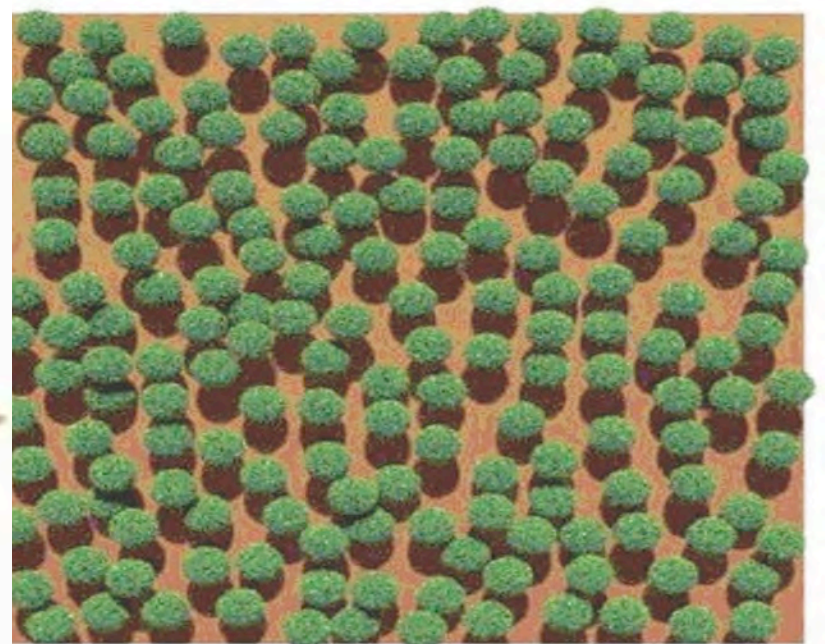
Current models



Thanks to B. Pinty for fig.



Reality



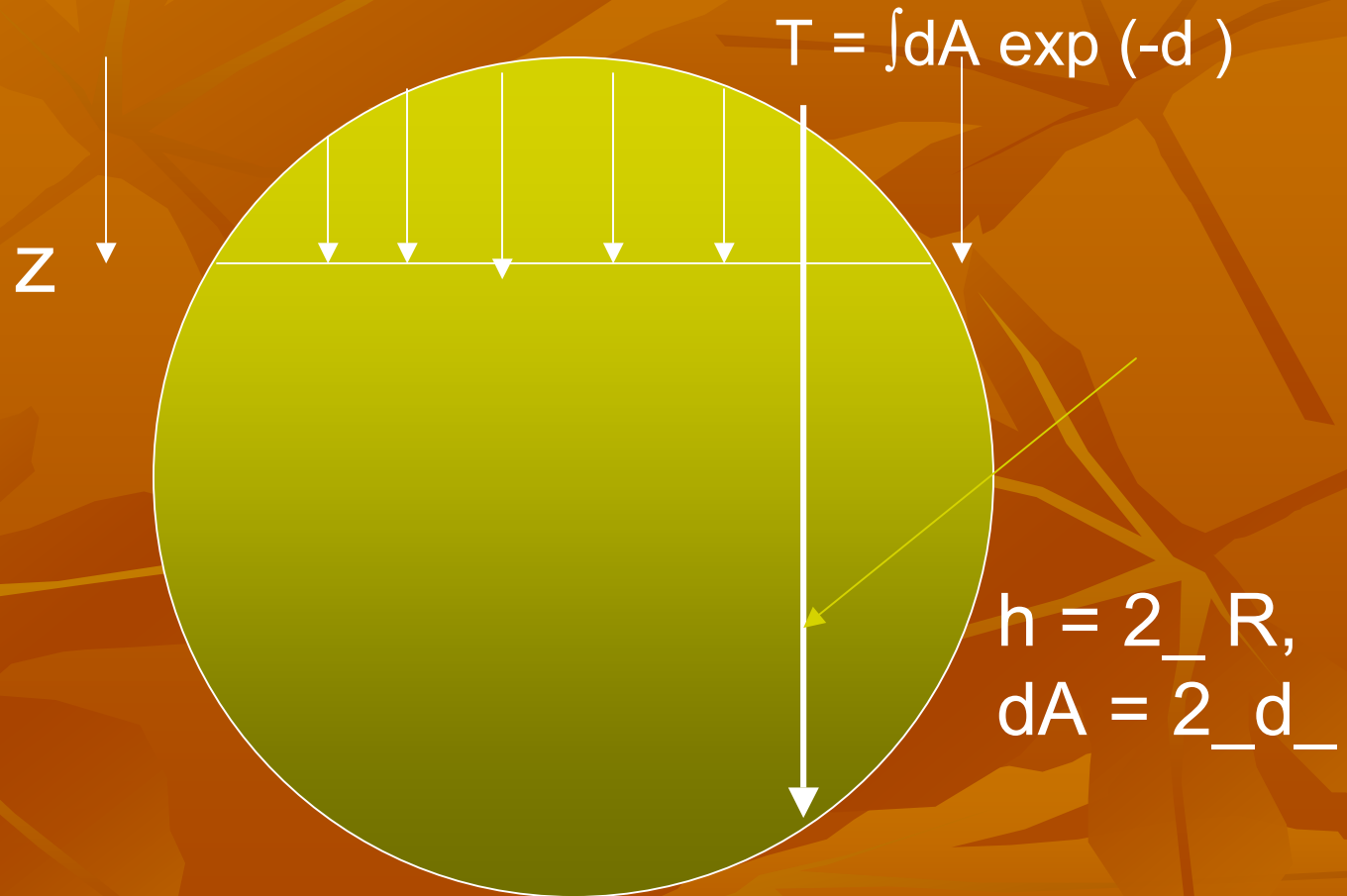
Thermal Infrared – “Long-wave”

- Downward spectral from atmosphere – atmospheric model lumps into total.
- All land surface objects have commonly been treated as black bodies but spectral emissivity of soils far from black body
- If emissivity included, overlying leaves treated with same logic as in solar – i.e a cloud.

Spherical Bush – an alternative to cloud of leaves

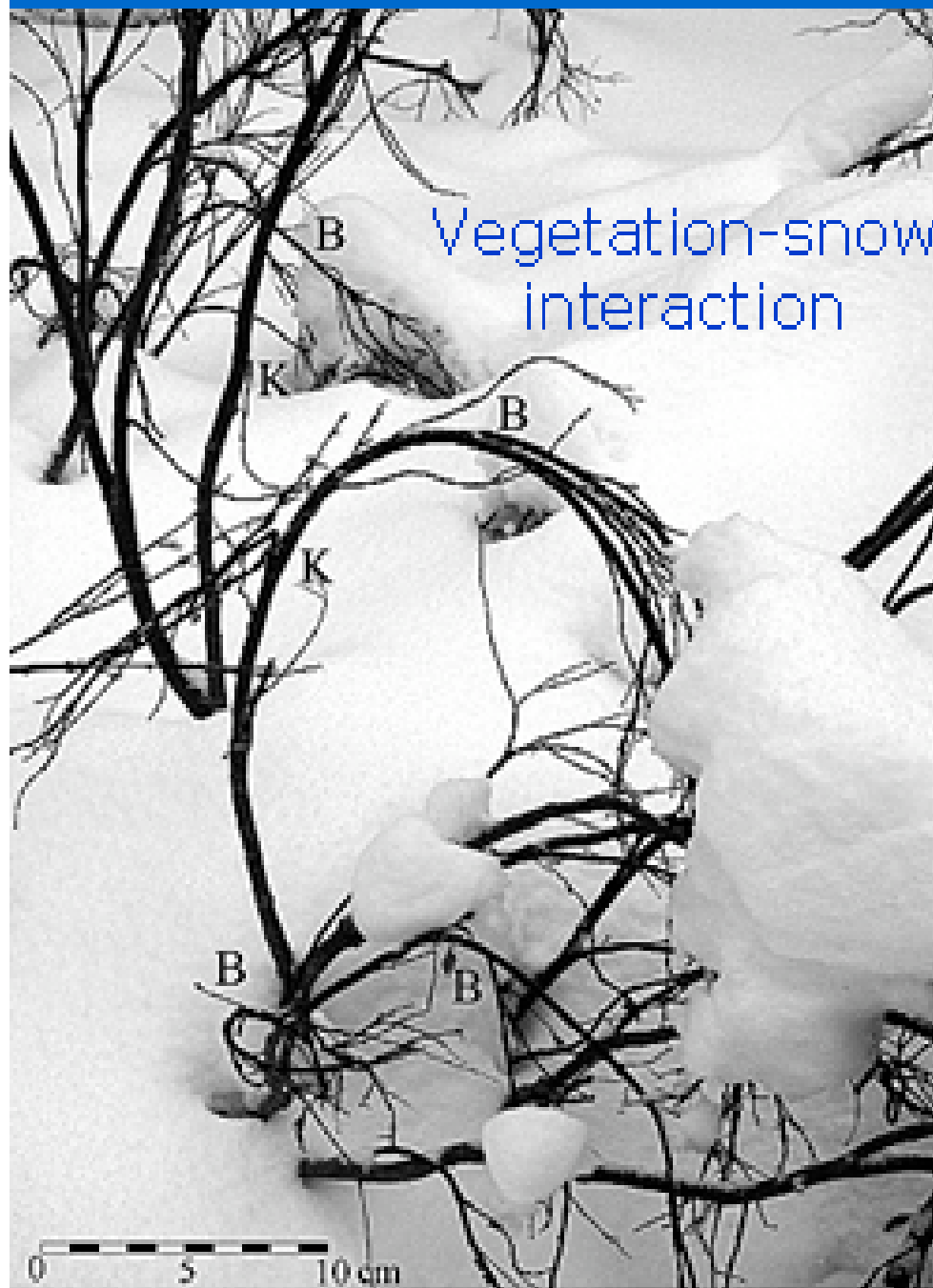
- Spherical geometry nearly as simple as planar.
- Can see the strong analytic differences between this idealization of a canopy and the cloud of leaves view.
- Through statistical modeling, can treat interactions between bushes and in limit of closed canopy recover the homogeneous cloud view.

Bush distribution of paths –linear.



What and where makes a difference?

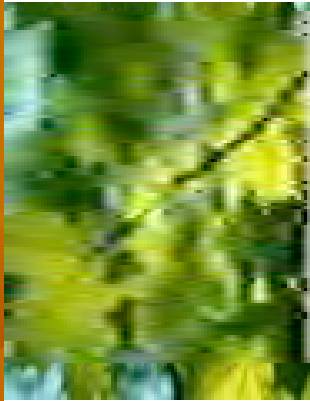
- Most important is the simple geometric effect of shadows – the radiative transfer issues within canopy alone make less difference but also important.
- Canopy open and underlain by bright soil –semiarid – or by snow.
- Factors of two or more differences and description of energy loading between canopy and soil/snow



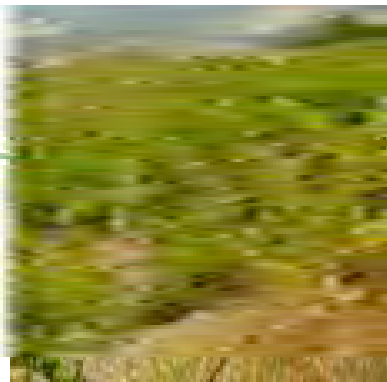
End of winter







Leaf Scale



Plot Scale



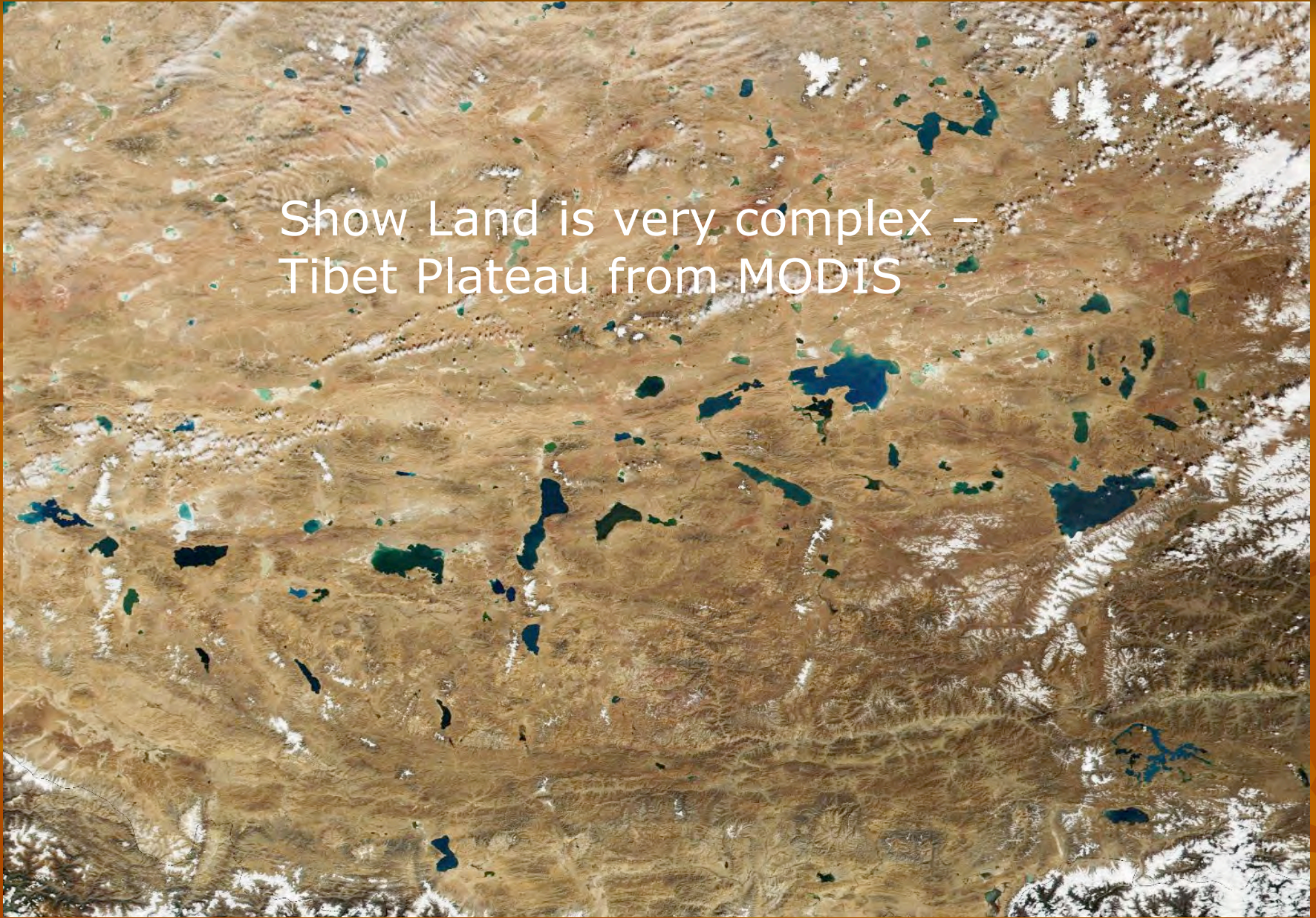
PFT fractions



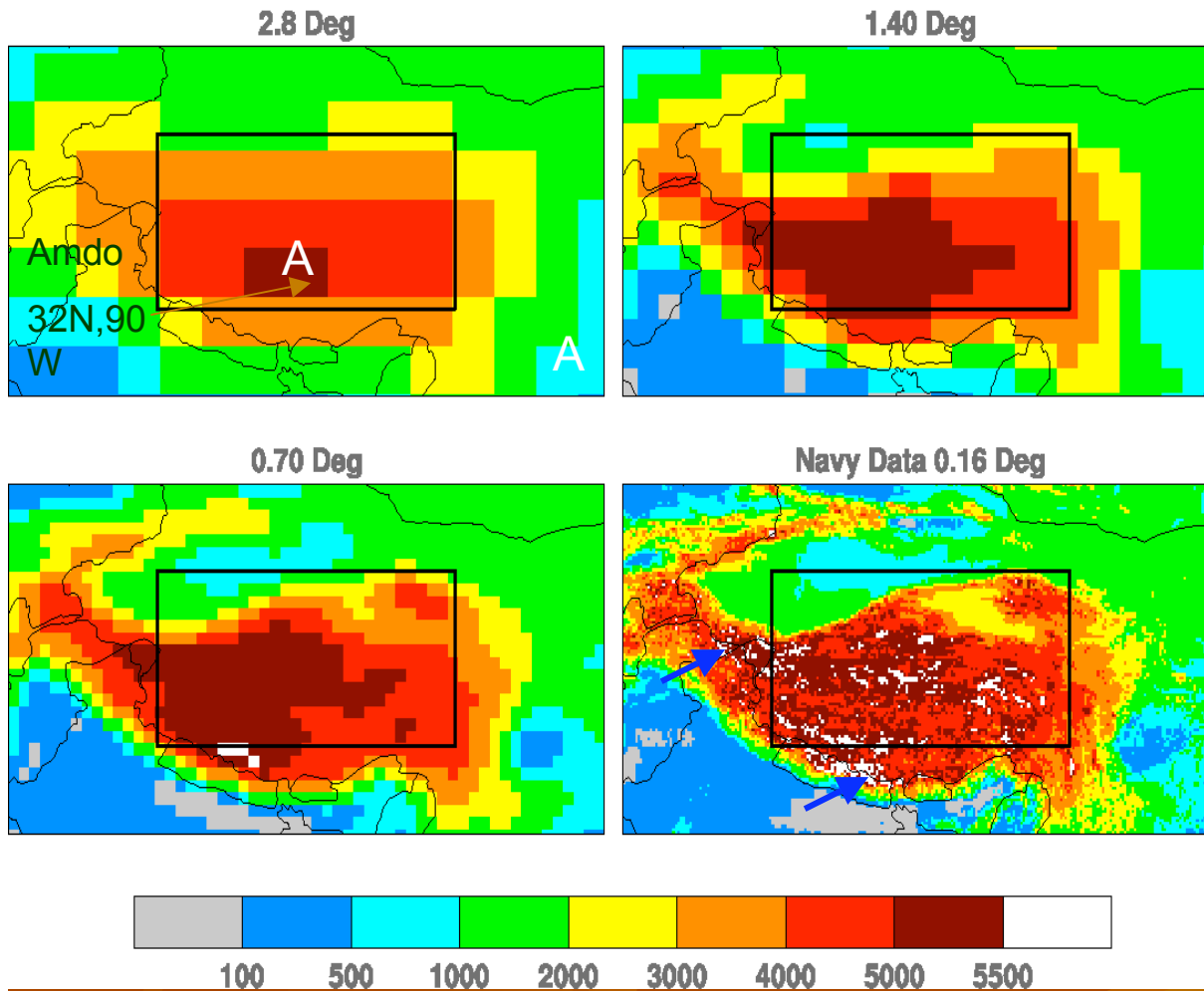
Major source of error?

- Models do separate energy budgets for canopy and underlying surface but details have large errors.
- Very difficult to get the turbulent flux energy exchanges with better than order of magnitude accuracy without a lot of research, e.g. 3-D fluid modeling- that appears not to have been done.
- Statement that adding trees decreases snow cover probably right but as a quantified feedback –very flimsy.

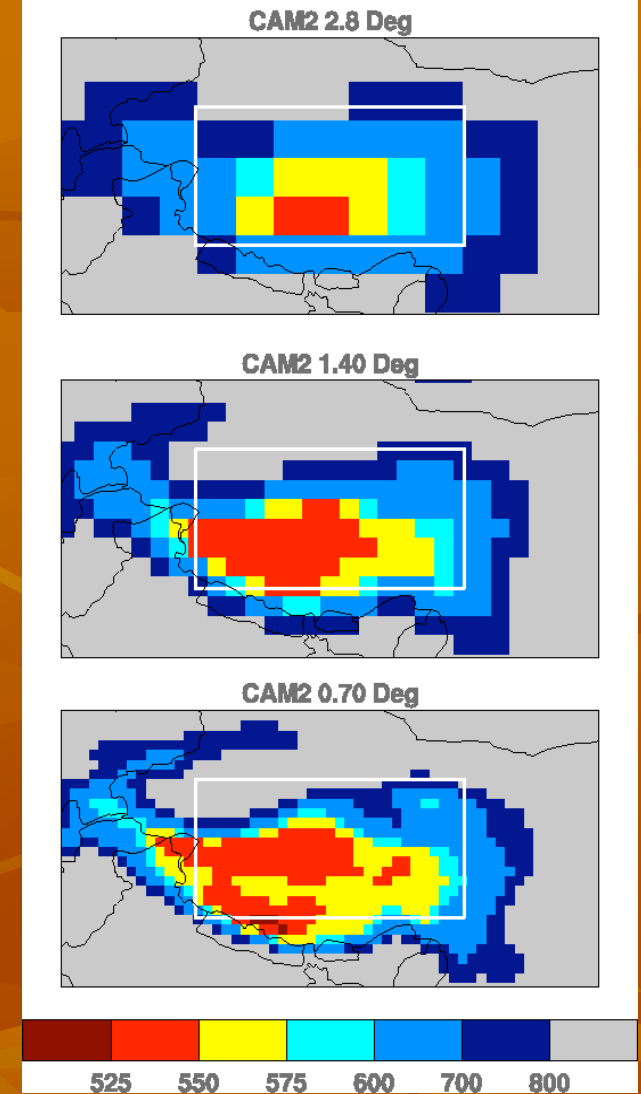
Show Land is very complex –
Tibet Plateau from MODIS



Tibetan Plateau (30-40N 80-100E) Surface Elevations

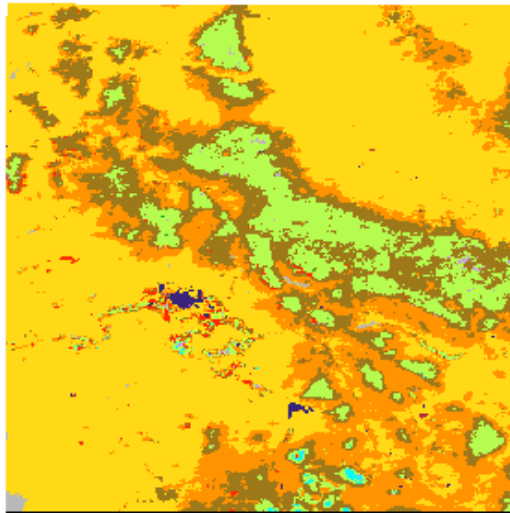


July Surface Pressure (mb)

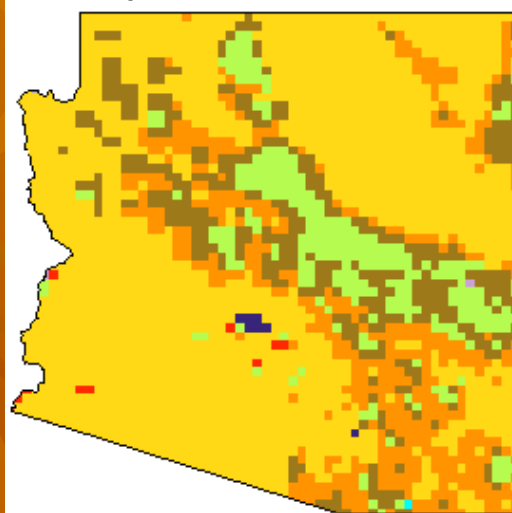


Example: land cover from satellite pixel scale to Climate model

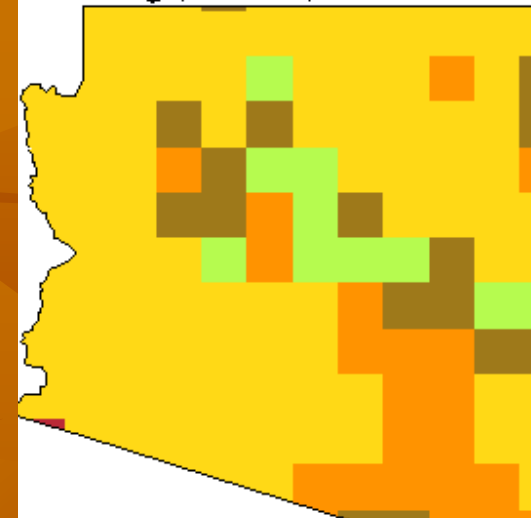
1-km Resolution



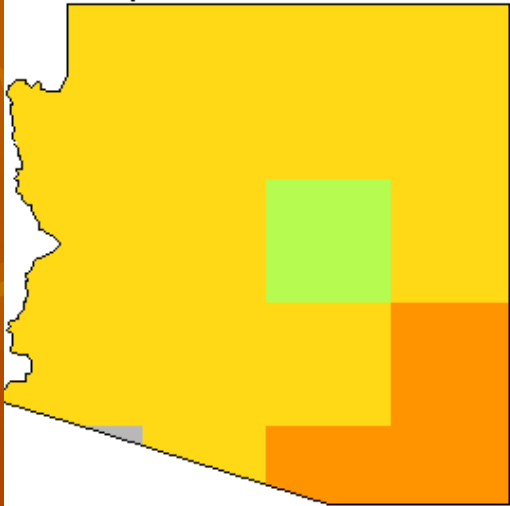
0.1 Deg (1800x3600) Resolution



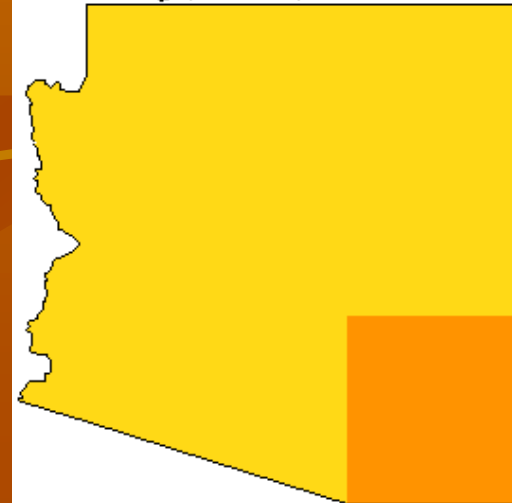
0.5 Deg (360x720) Resolution



1.4 Deg (128x256) Resolution

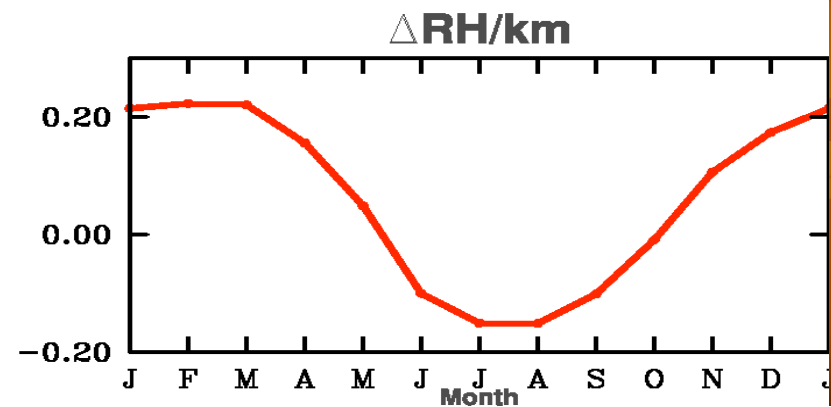
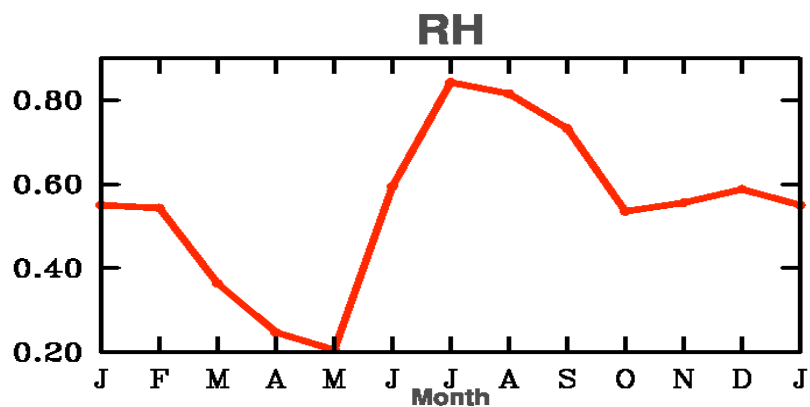
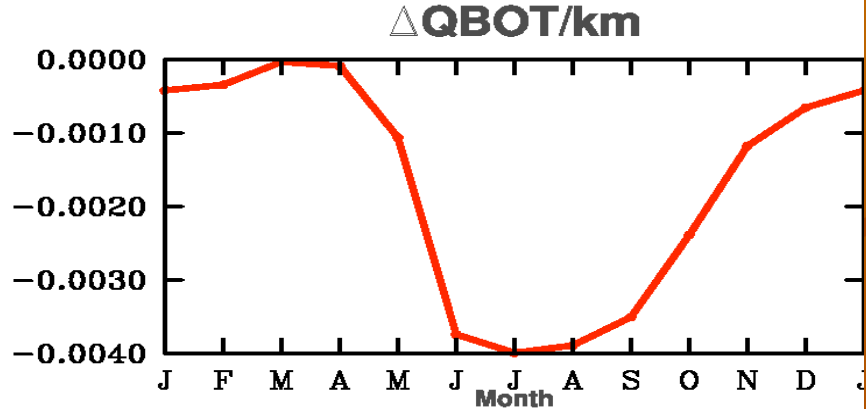
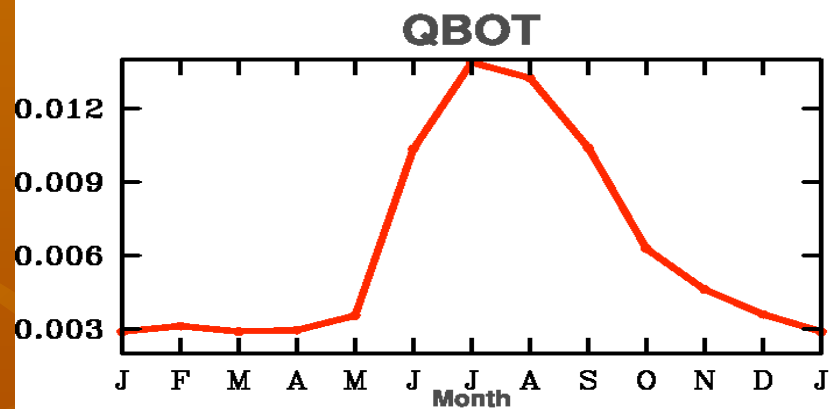
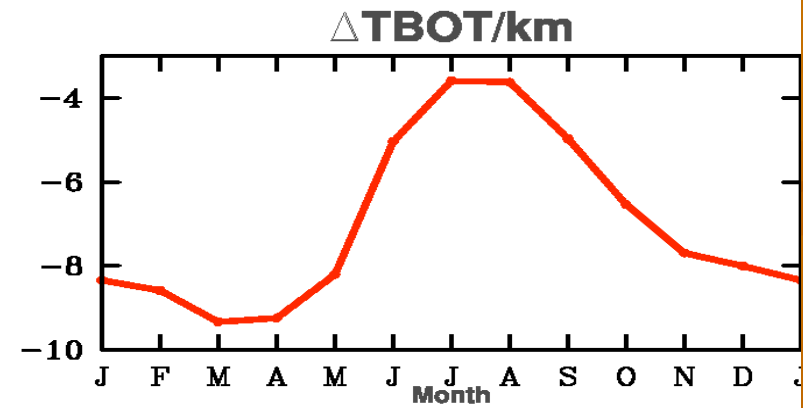
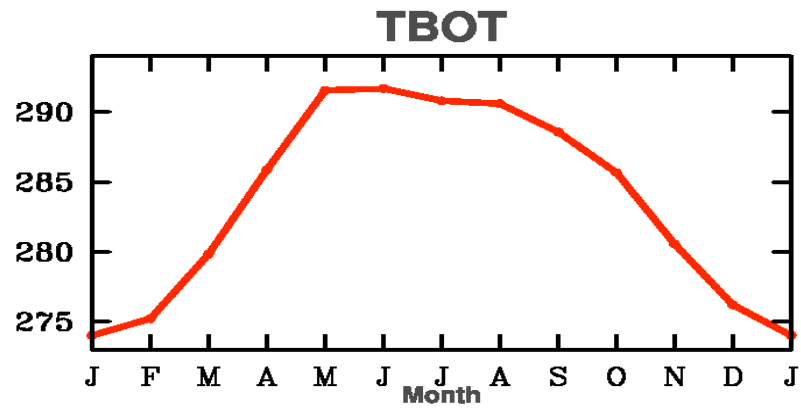


2.8 Deg (64x128) Resolution



**Arizona: 53.1%Open shrubs 19.6%Grasslands 15.8%Woody Savanna
9.7% Evergreen Needle-leaf 0.5%Croplands 0.27% Urban (900 km2)**

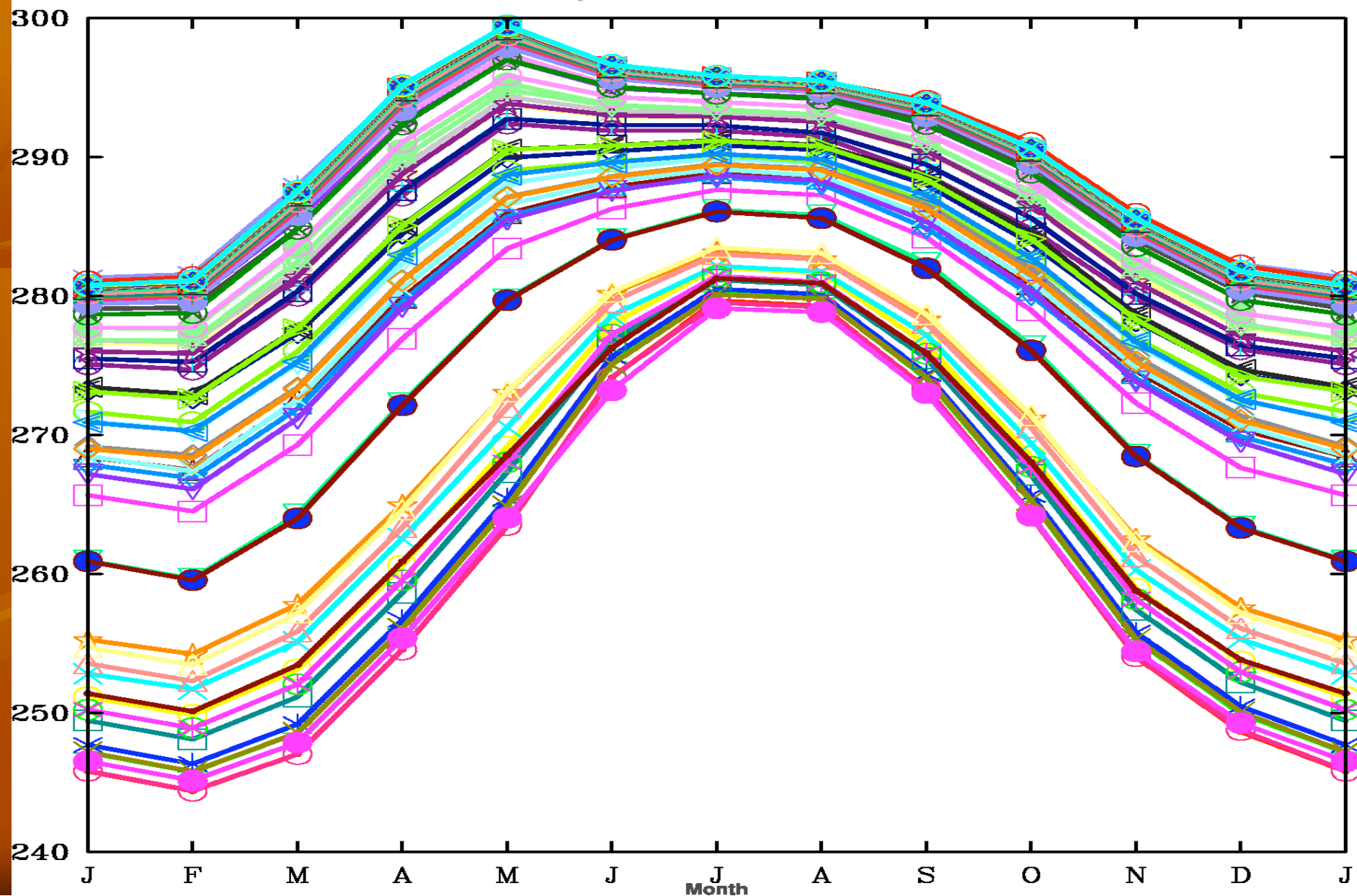
22 Year Average :: Lowest Layer Model Fields and Vertical Gradients
Himalaya :: 32.09, 75.94



1980-86 2-m Surface Air Temperature (°C) :: 32.09N, 75.94E Grid

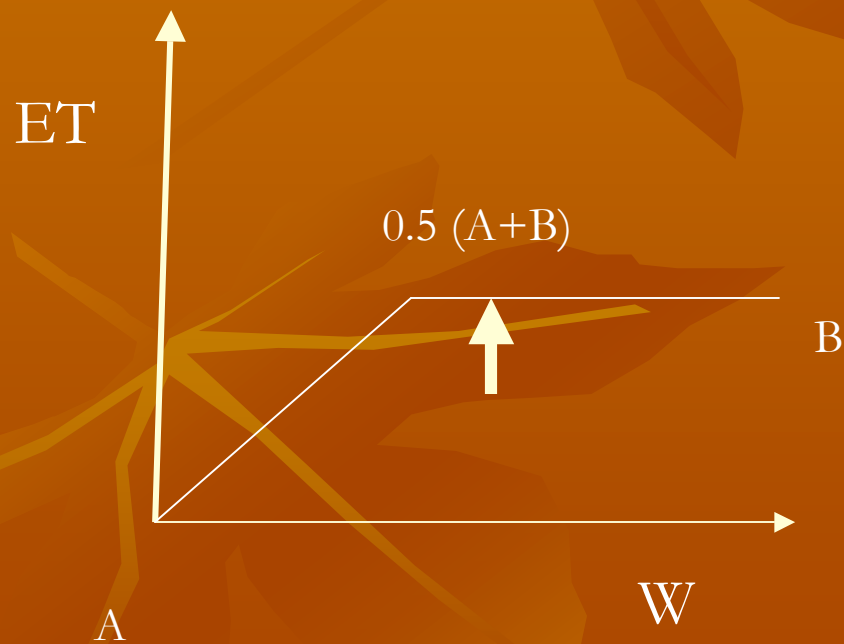
△ TSA= 244.38,299.49 △Z= 193m, 5069m △Patch= 0.54%, 2.94%

Total Patches= 65 Vegetation: 59, Baresoil: 4 Glacier: 2



Simple example of scaling issue

- Wet versus dry surface



$$ET(A/2 + B/2) = ET(B)$$

$$(ET(A) + ET(B)) = 0.5 ET(B)$$

Conclusions

- Land surface processes in climate models is still a frontier area.
- Many different ideas as to what is most important and not enough quantification as to what is wrong and what are the most important errors to address.
- Climate models have become easier to use but the land component still have pitfalls for the inexperienced