

What information is available and what is missing for energy and water cycle process modeling efforts, and the ways to address the gaps

Eric F Wood

With help from

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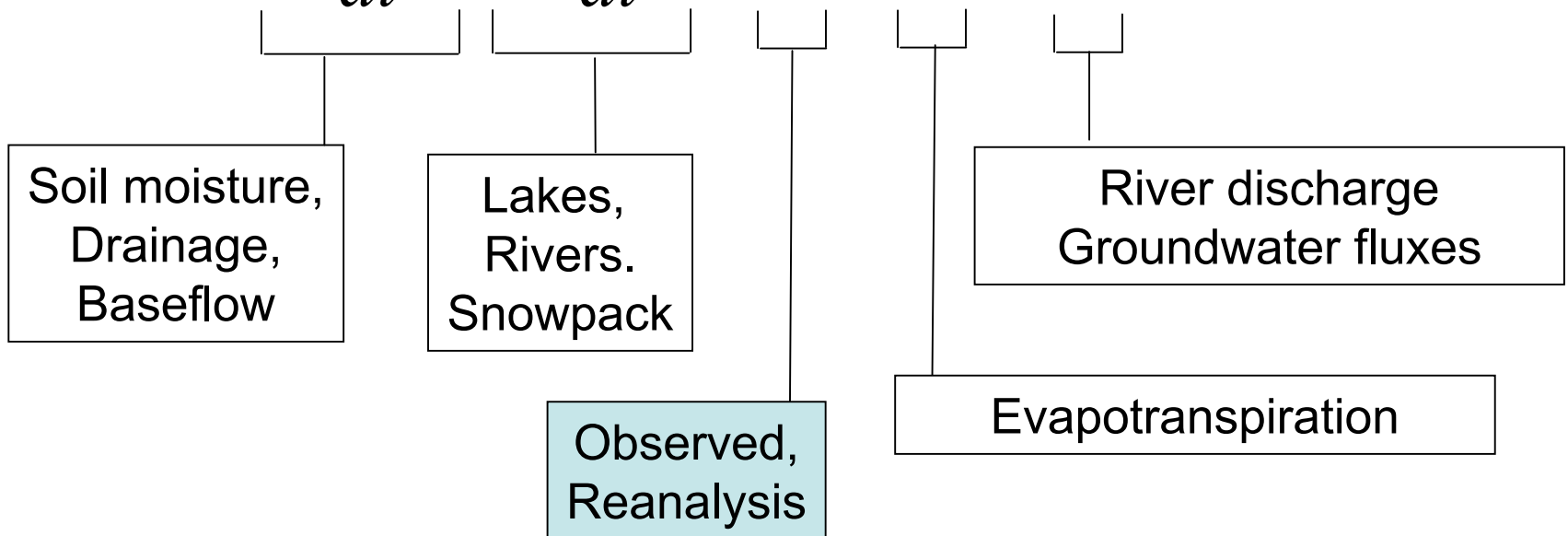


**NORTHERN EURASIA LANDSURFACE PROPERTIES AND
CHANGE AND ITS ROLE IN THE GLOBAL EARTH
SYSTEM AND MODELS**

August 12 – 17, 2007, Aspen, CO

Basic water budget equation

$$\frac{d\theta_{SM}}{dt} + \frac{d\theta_{SW}}{dt} = P - ET - Q$$

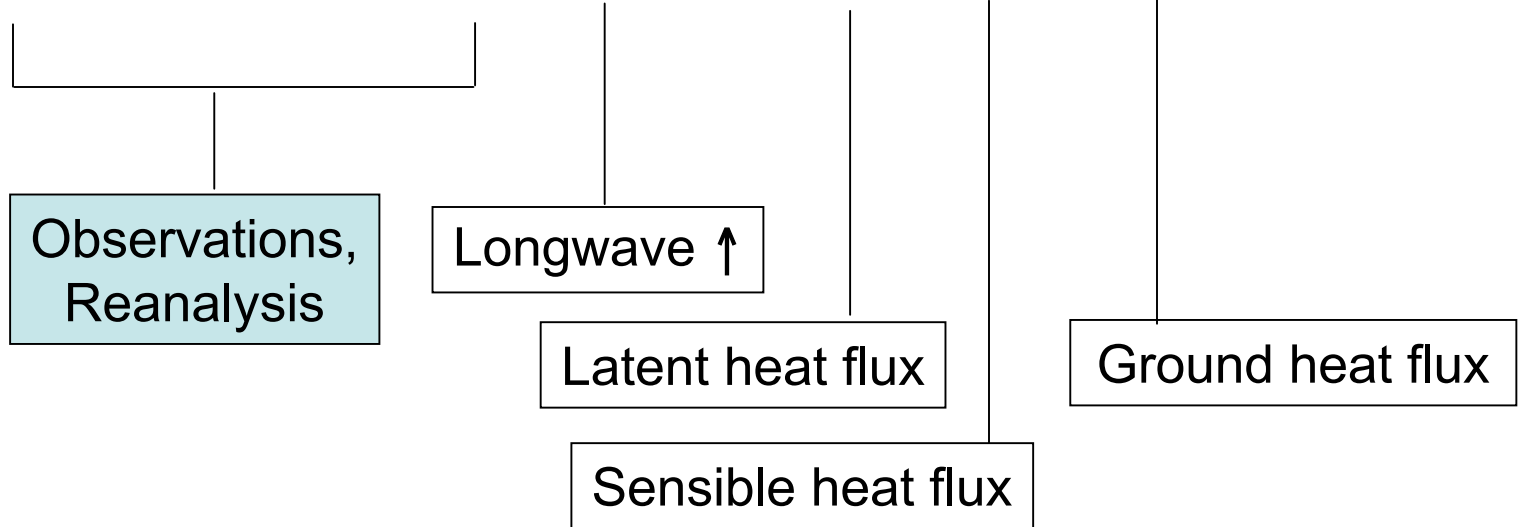


Observations

Parameterized

Basic energy budget equation

$$R_S(1 - \alpha) + R_L - \epsilon k T_S^4 = \lambda E + H + G$$



Observations

Parameterized

What information is available and what is missing for energy and water cycle process modeling efforts, and the ways to address the gaps

“What information is needed for energy and water cycle modeling?”

1. Forcing data: Surface Meteorology, incoming/downward radiation
2. Data used in process parameterizations:
 - Land cover and, for vegetation, its related phenology
 - Soil data including composition, hydraulic and thermal properties
 - Topography, including micro-topography that controls fens and bog areas
3. Data used to validate process parameterizations:
 - River discharge
 - Snow cover extent
 - Lake and reservoir levels
 - Soil moisture, groundwater levels
 - Bore hole temperature logs for permafrost and seasonal freezing
 - Evapotranspiration

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Some examples of the needed data sets

Precipitation Gauge Network

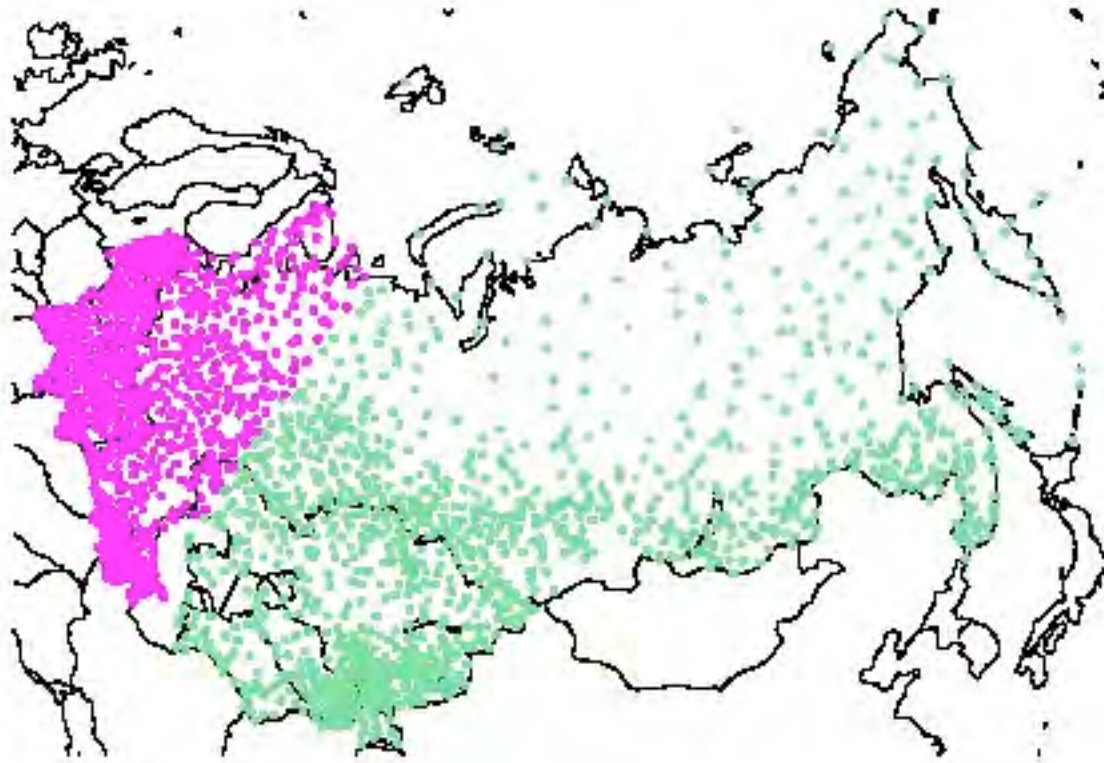
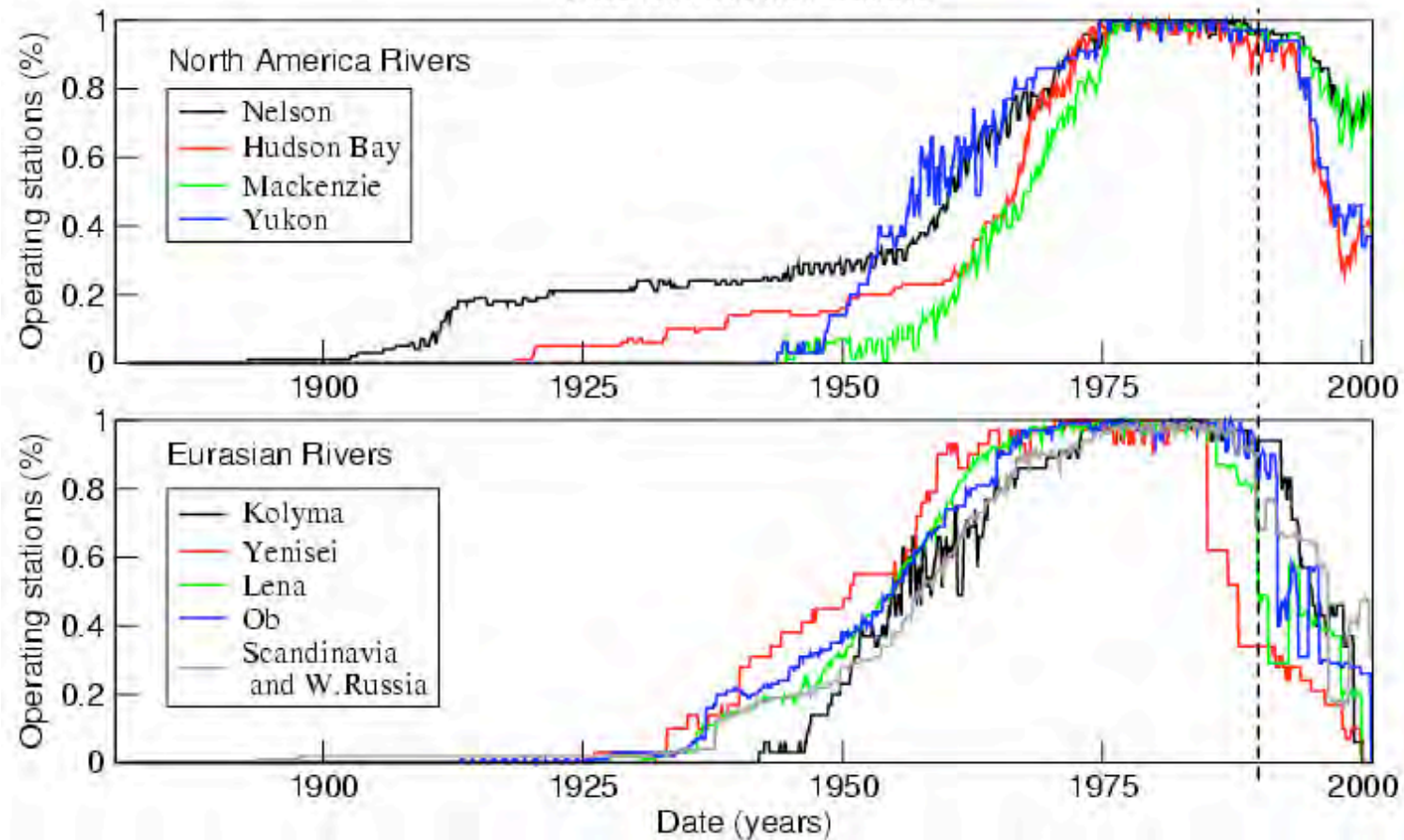


Figure taken from NCDC Documentation for Dataset 9813:
Daily and Sub-daily Precipitation for the Former USSR.

Discharge Gauge Network

Operating stations in Arctic basins (% of the maximum # stations)

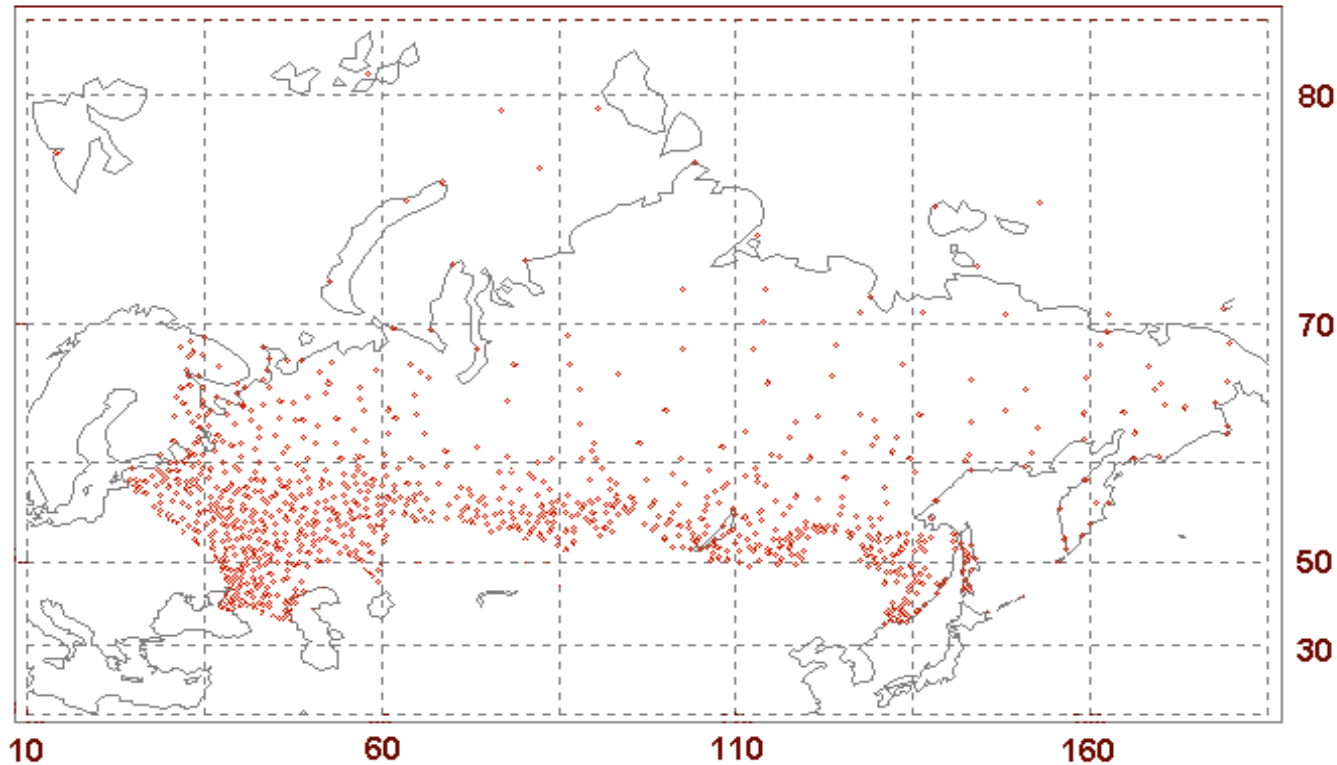
1198 R-Arctic Net stations



R-ArcticNET



In Situ Observations

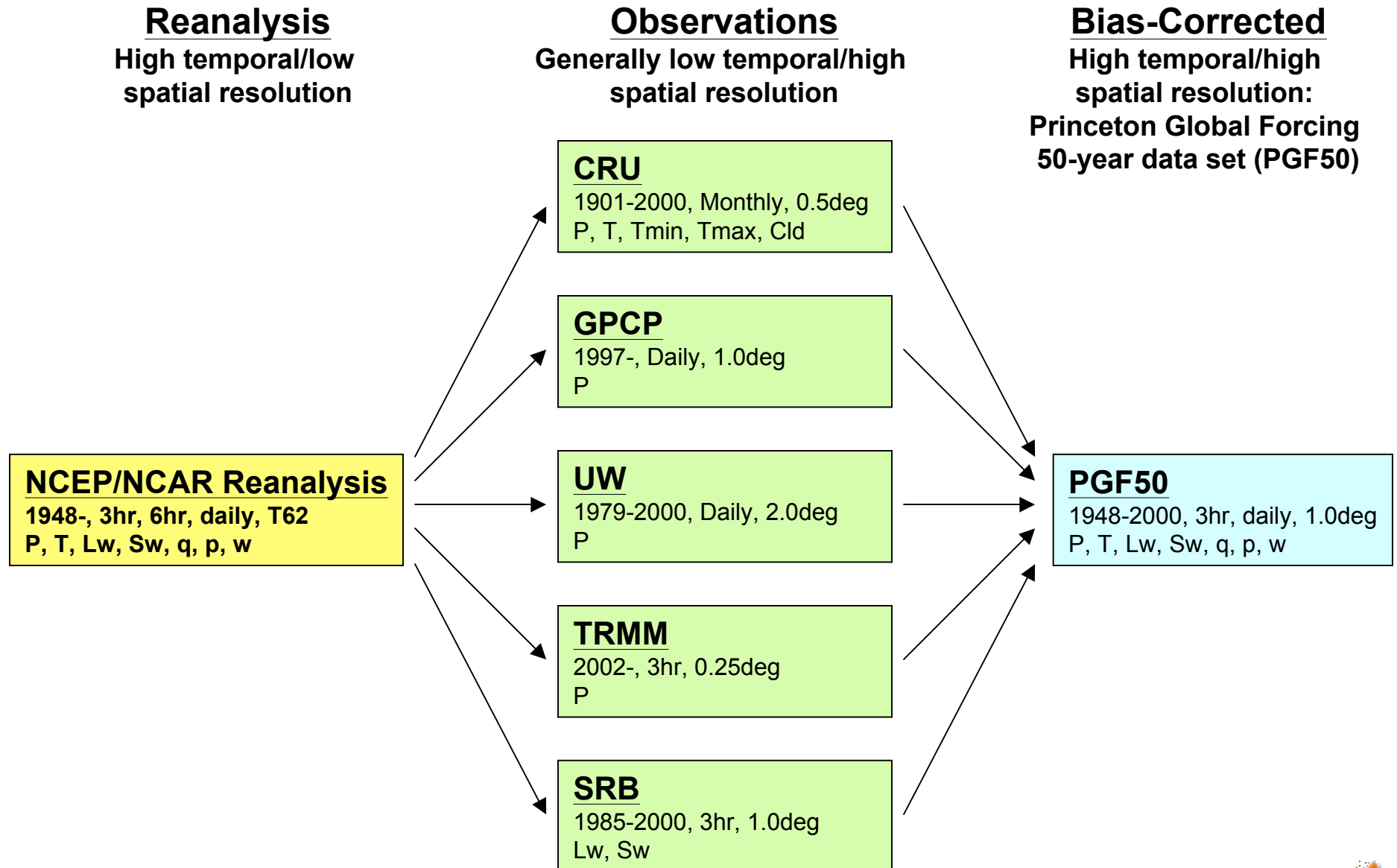


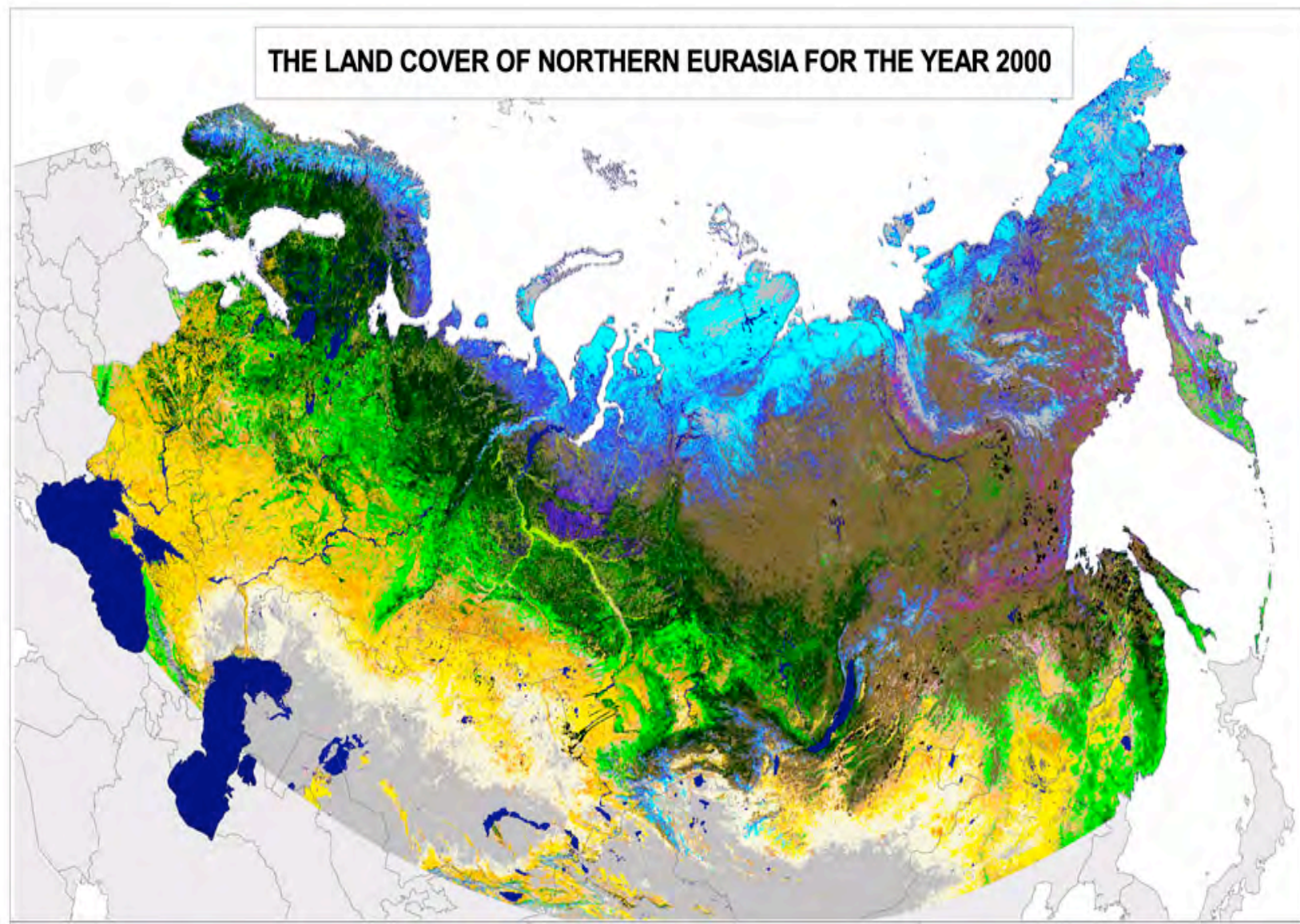
Data on snow cover characteristics includes the following information:

- snow cover thickness along snow courses over different landscapes,
- degree of the area covered by snow cover along a snow course;
- snow cover density
- Snow Water Equivalent

Source: Nina Speranskaya, State Hydrological Institute, Russia

Global Forcing Dataset (see Sheffield et al 2006, *J Climate*)

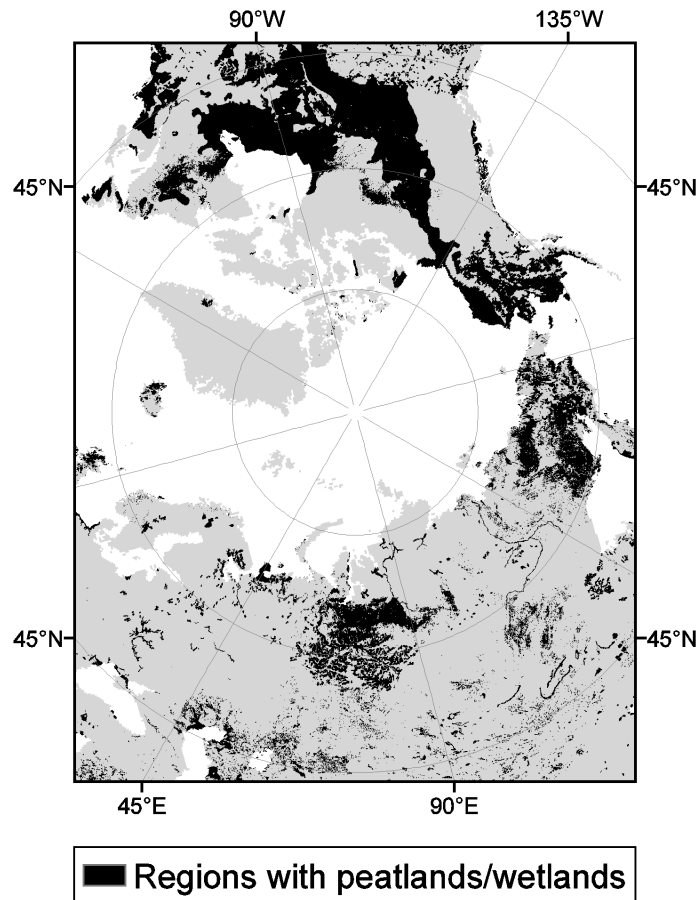




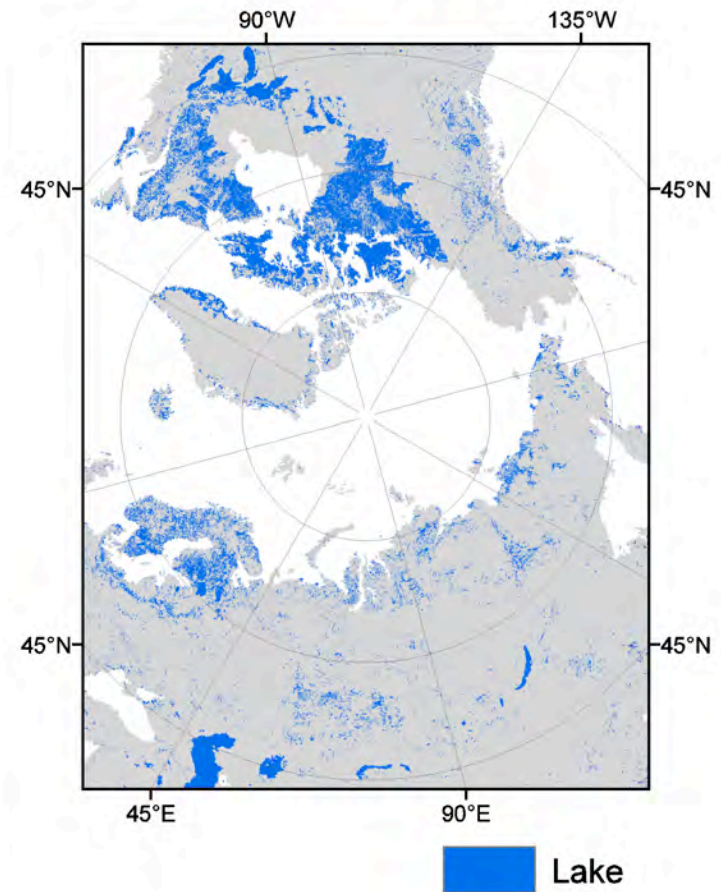
Is the accuracy and resolution sufficient?
Is the accompanying phenology available?

Where do we find peatlands, what is the distribution of lakes?

Distribution of peatlands/wetlands
in the Northern Hemisphere



Distribution of lakes
in the Northern Hemisphere



Lehner, B., and P. Doll (2004), Development and validation of a global database of lakes, reservoirs and wetlands, *Journal of Hydrology*, 296, 1-22.

What information is available and what is missing for energy and water cycle process modeling efforts, and the ways to address the gaps

“...energy and water cycle process modeling efforts?”

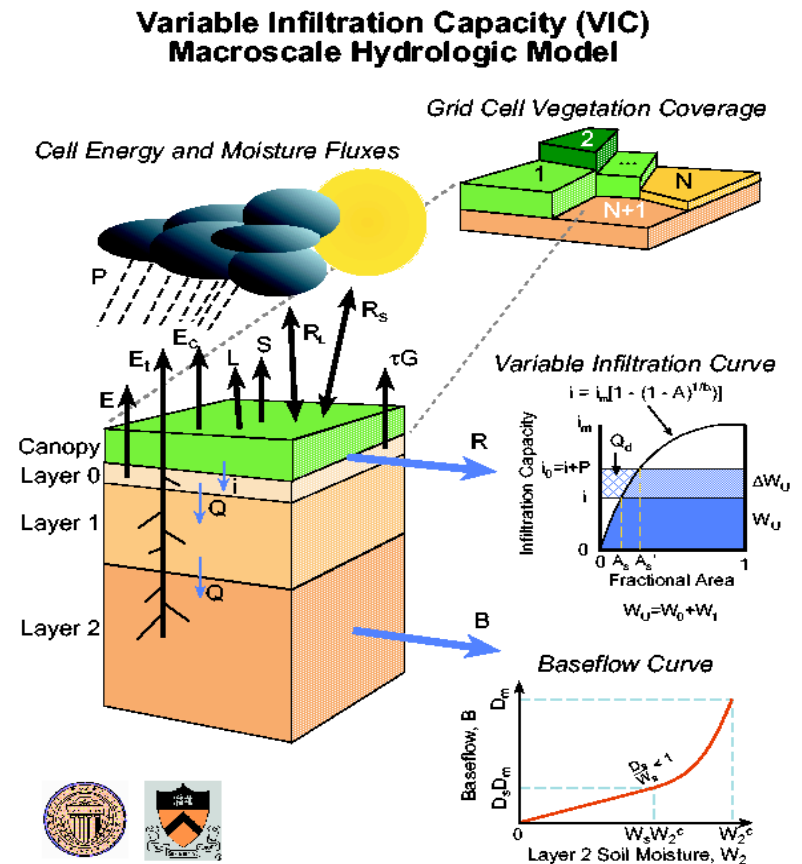
In-situ observations are too sparse to support land surface modeling forcing data sets having spatial resolutions finer than about 1-degree.

Thus, to what extent are we capturing the temporal and spatial dynamics of energy and water cycle dynamics across the NEESPI domain?

Are the energy and water cycle process parameterizations sufficiently accurate that they can be used to attribute LULCC and climate change attribution?

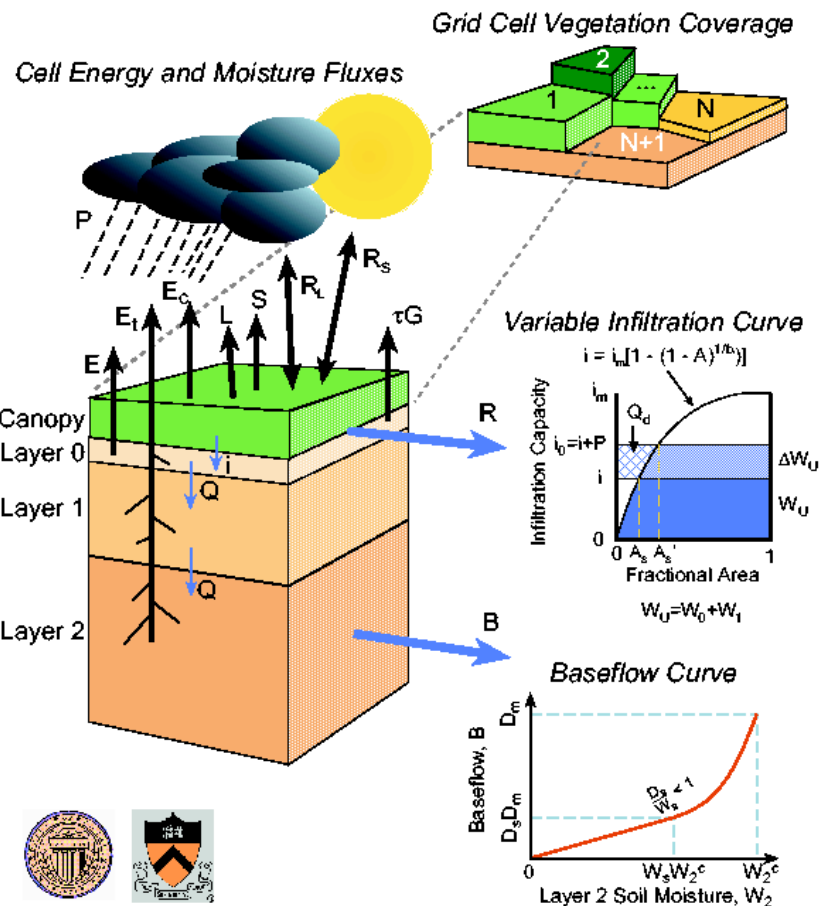
Typical Land Surface Modeling Framework

- VIC hydrology model
 - Large, “flat” grid cells (e.g. 100x100 km)
 - Land cover “tiles” (vegetation types)
 - On hourly to 3-hrly time step, simulate:
 - Soil moisture/water table
 - Snow pack
 - Runoff
 - Lake/surface water
 - Evapotranspiration/Latent heat flux
 - Soil temperature
 - Sensible heat flux
 - Ground heat flux

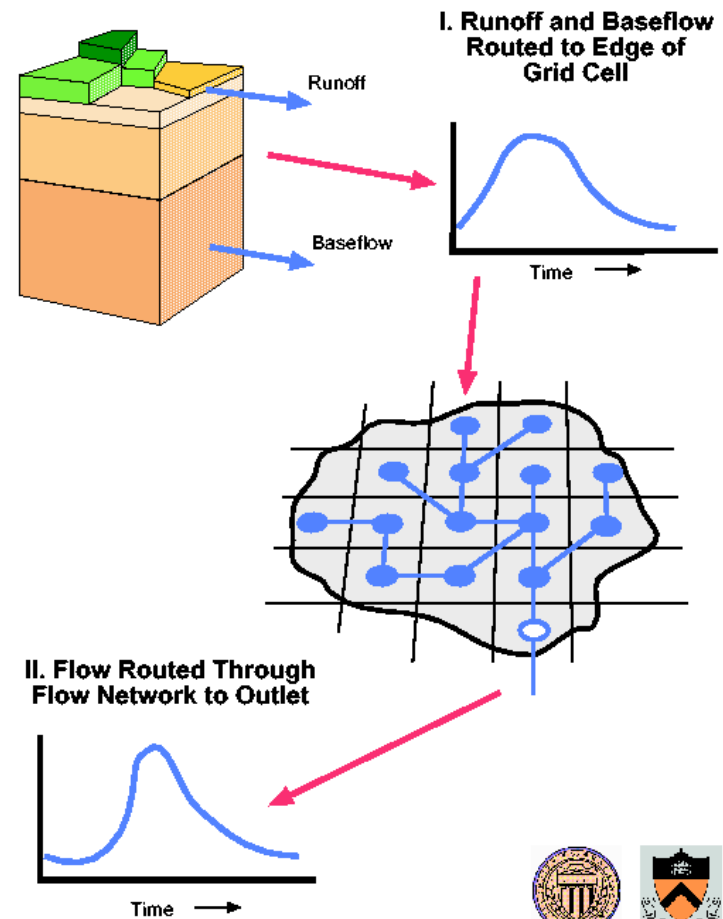


The Variable Infiltration Capacity (VIC) Model

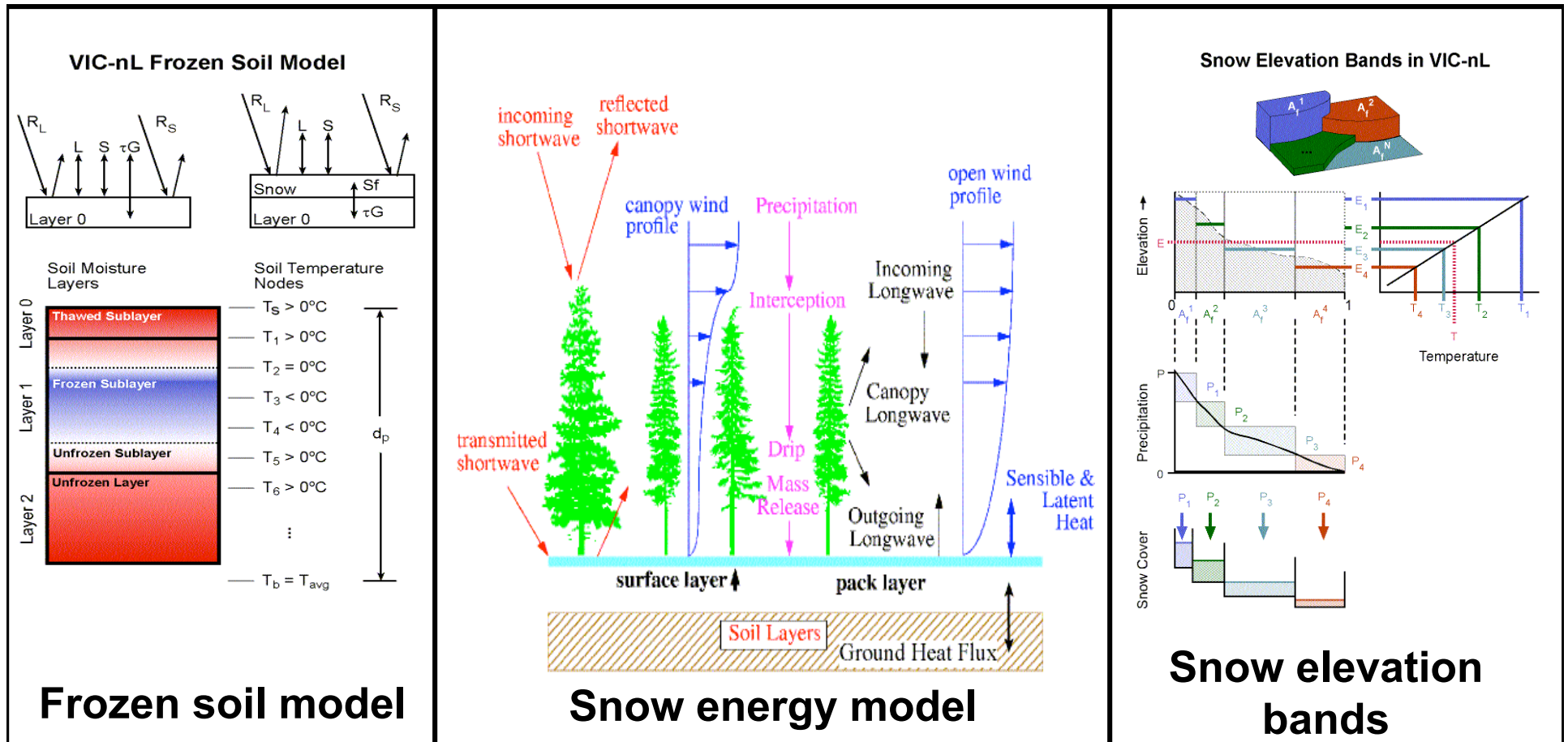
Variable Infiltration Capacity (VIC) Macroscale Hydrologic Model

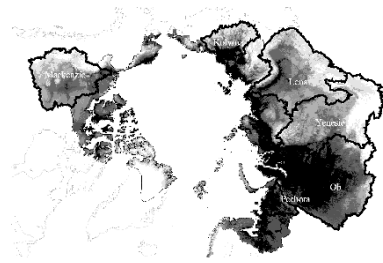


VIC River Network Routing Model



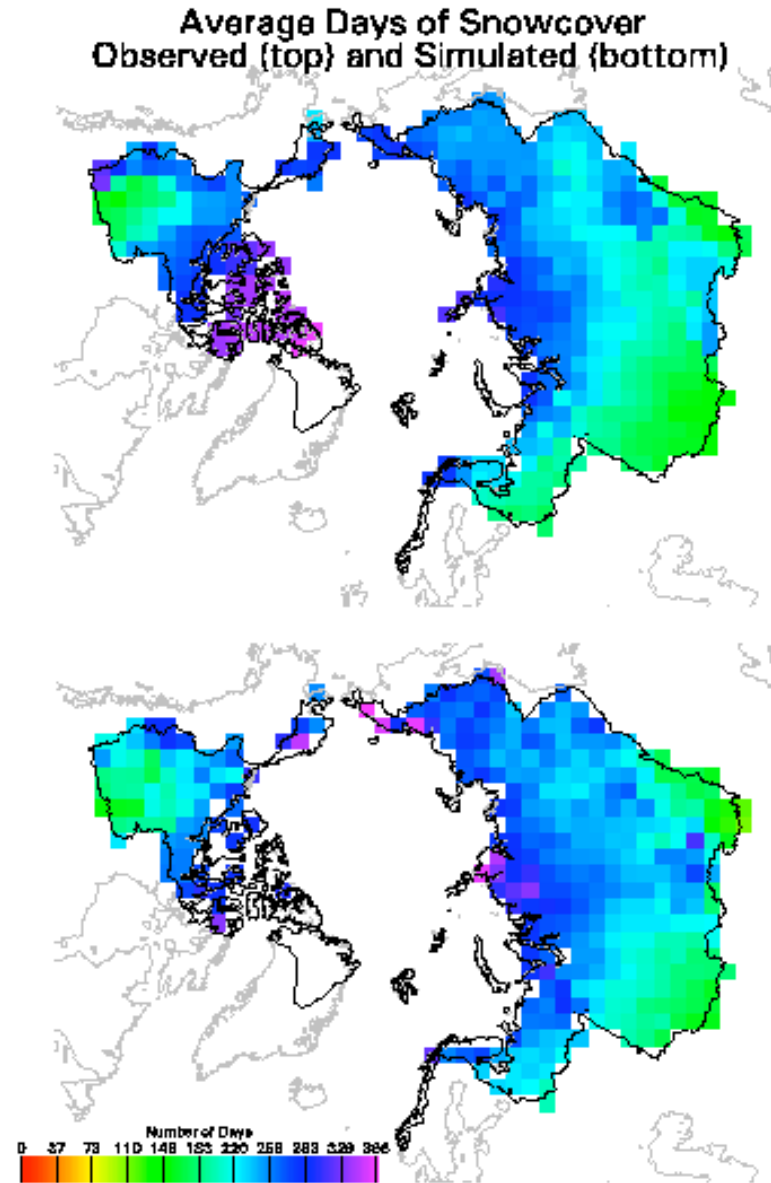
VIC Cold Season Schematic

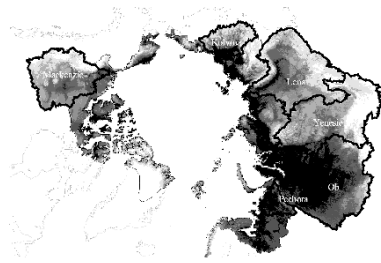




VIC Arctic Modeling

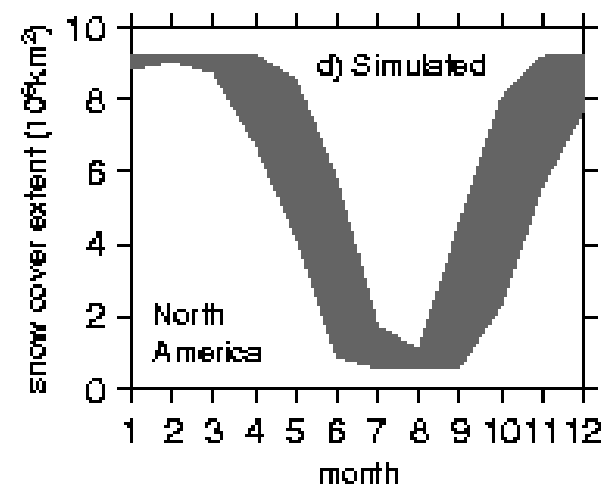
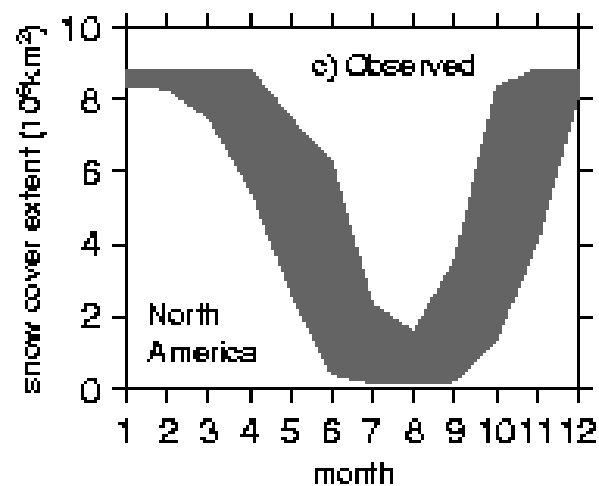
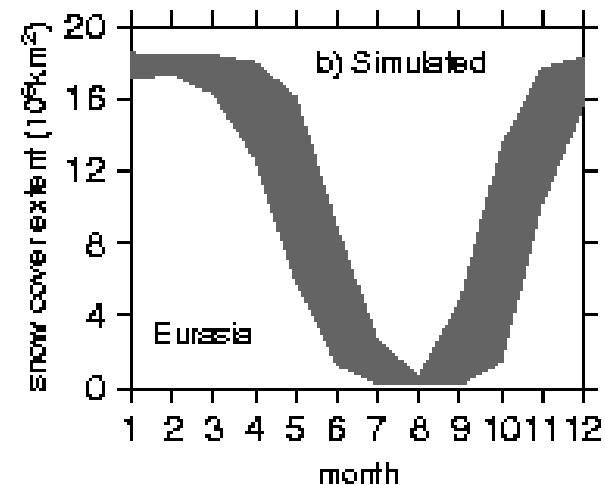
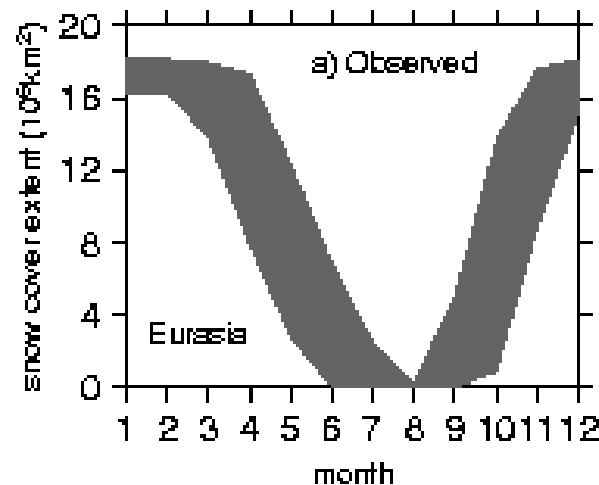
Annual average days
of snow cover for the
entire Arctic basin





VIC Arctic Modeling

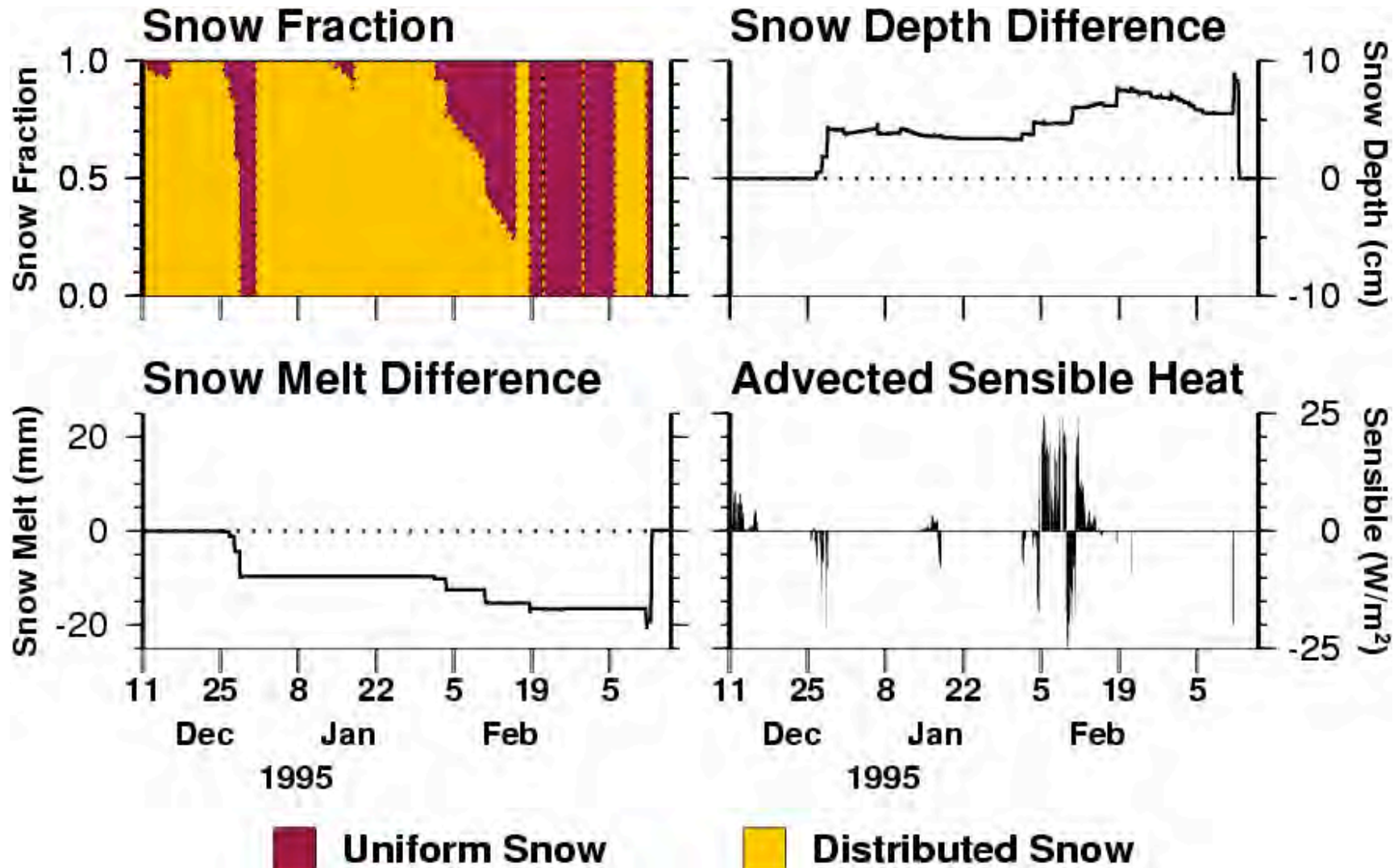
Monthly variation in snow cover extent
(1980-1993)

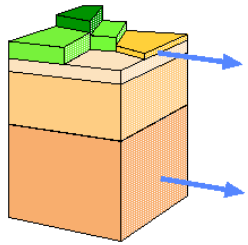


Spatial versus Uniform Snow Algorithm

Winter 1994-1995 – Rosemount, Minn.

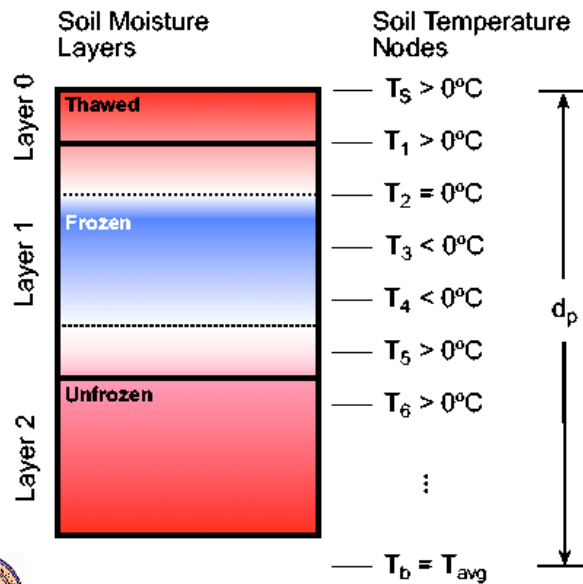
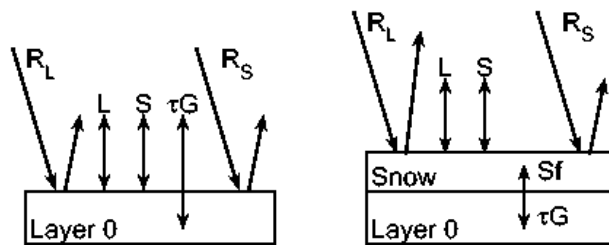
Algorithm Development



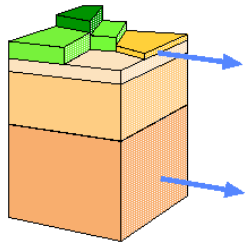


Spatially-distributed frozen soils

VIC Frozen Soil Algorithm

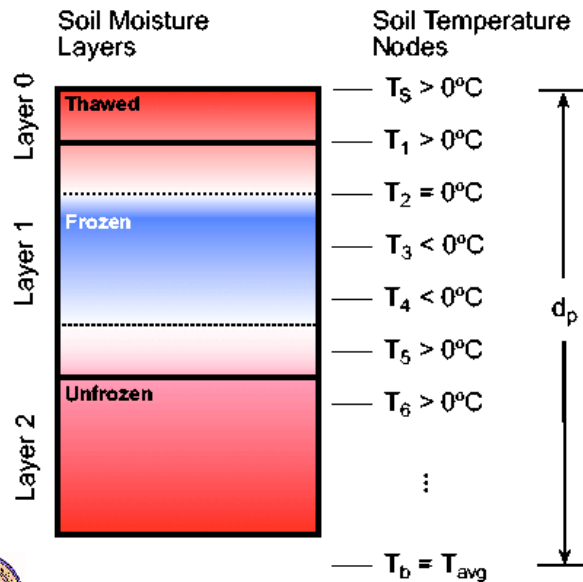
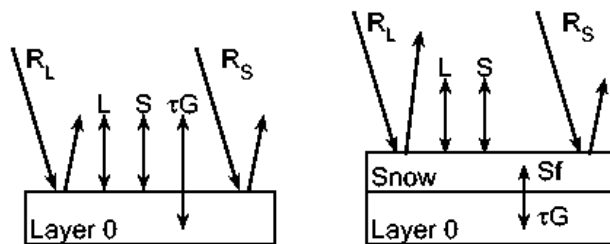


- Soil node temperatures solved via heat diffusion equation
- Ice content, infiltration rate and heat capacity calculated at nodes
- Assumed uniform temperature distribution across the grid cell allows spatial variation of infiltration capacity

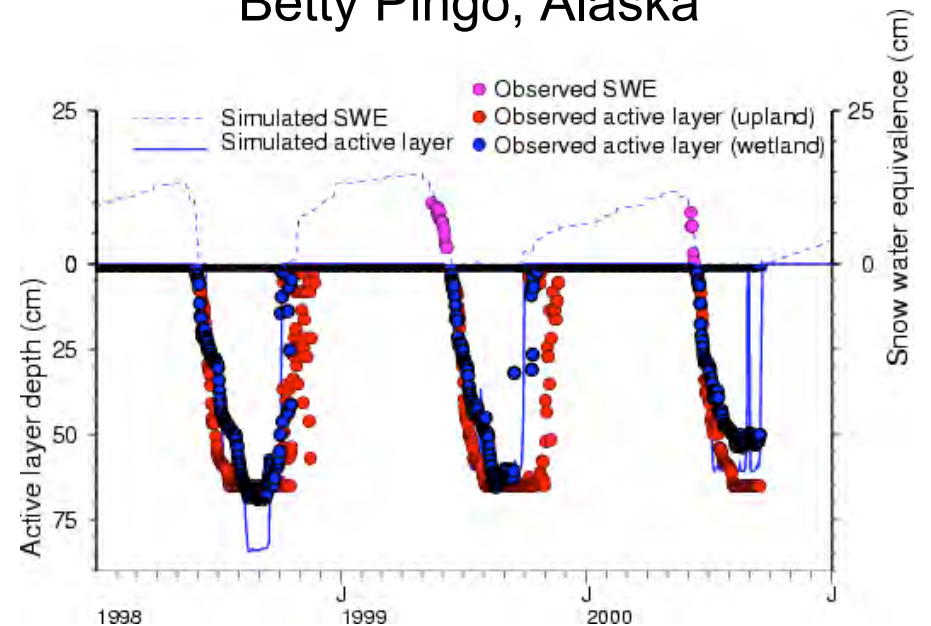


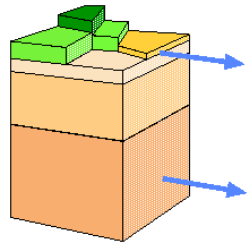
Spatially-distributed frozen soils

VIC Frozen Soil Algorithm



Betty Pingo, Alaska

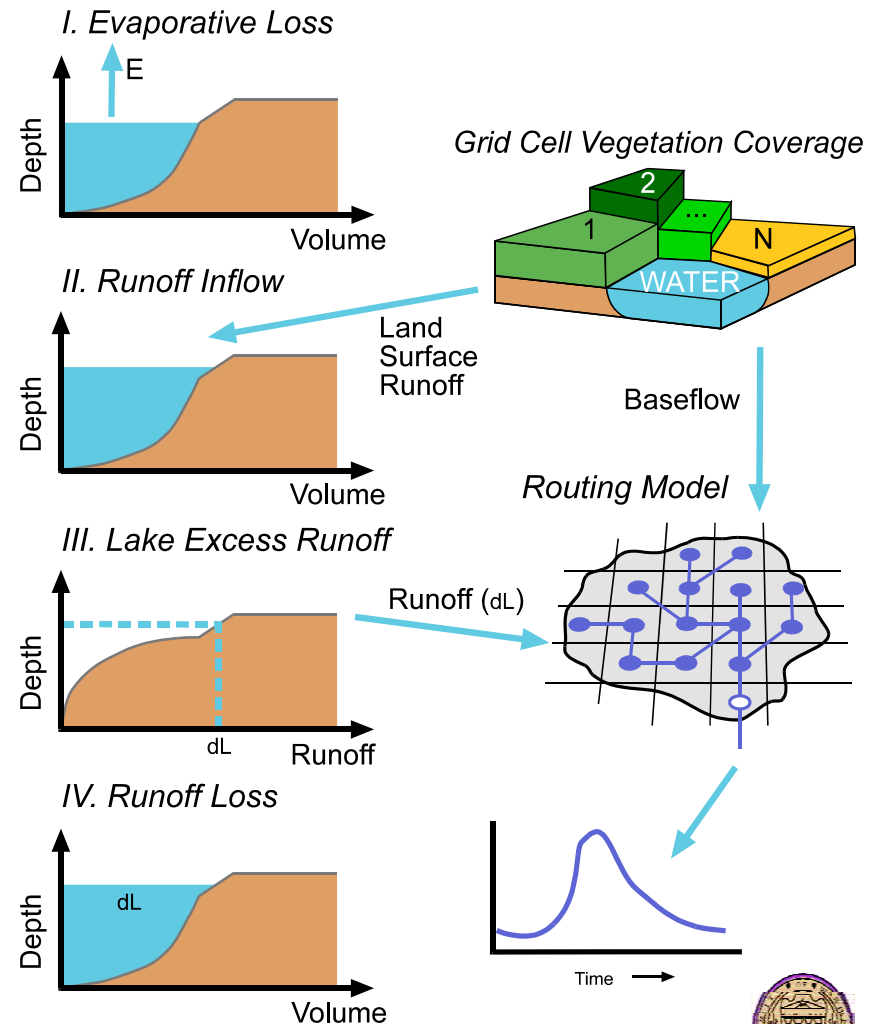


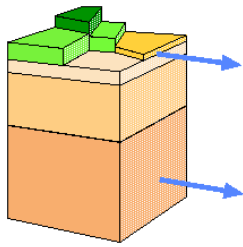


Lakes and Wetlands

- Lake energy balance based on:
 - Hostetler and Bartlein (1990)
 - Hostetler (1991)
- Assumptions:
 - One “effective” lake for each grid cell;
 - Laterally-averaged temperatures; and
 - Non-linear rule curve for release of runoff from grid cell lake.

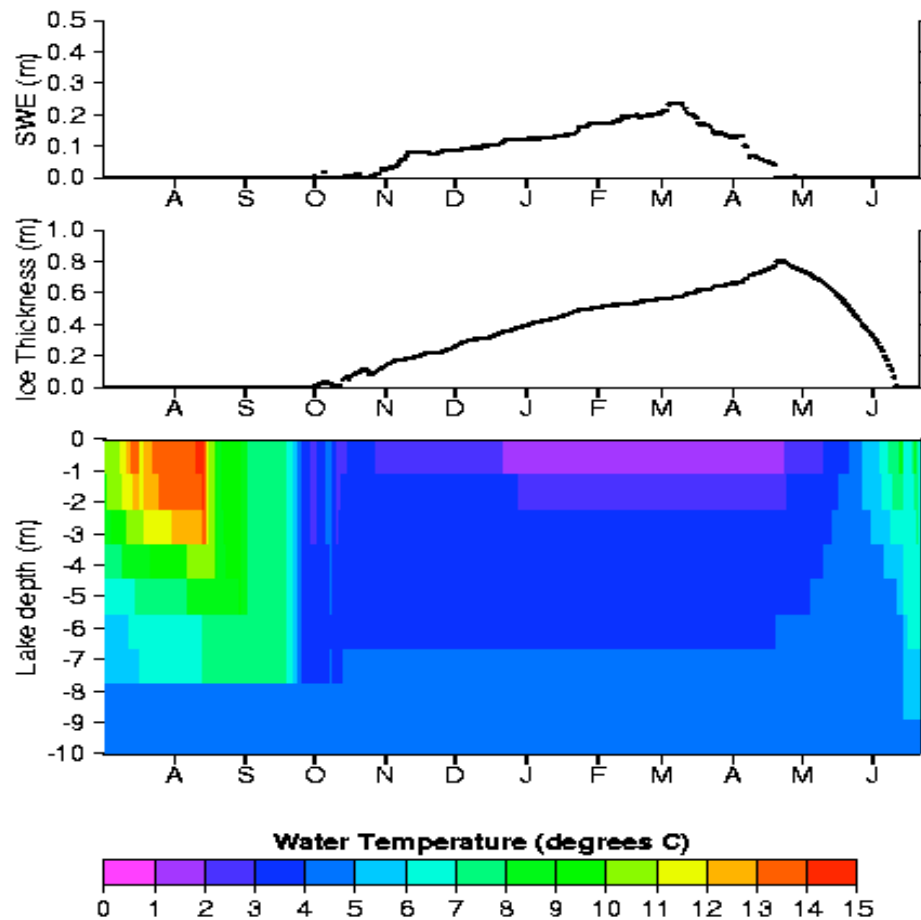
VIC Lake Algorithm



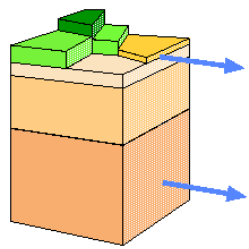


Lakes and Wetlands - Snow and Ice

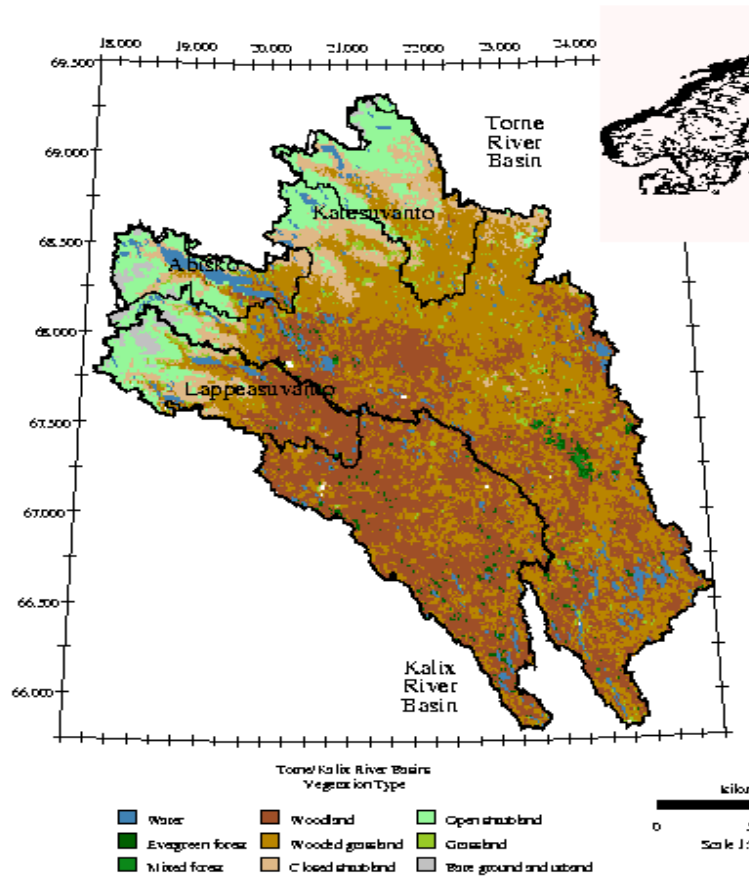
Location: 68.125 N, 19.125 E
July 1, 1985 - June 30, 1986



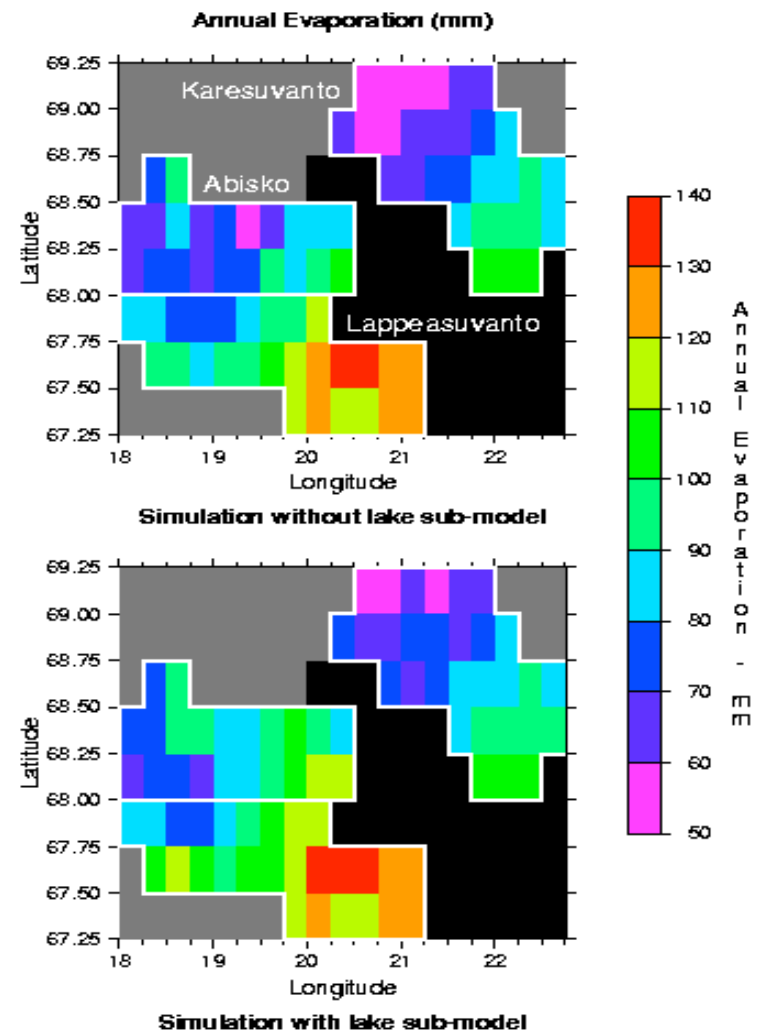
- ◀ Snow depth simulated on top of lake ice
- ◀ Lake ice - stays too long?
- ◀ Temperature profile - complete mixing in fall and spring



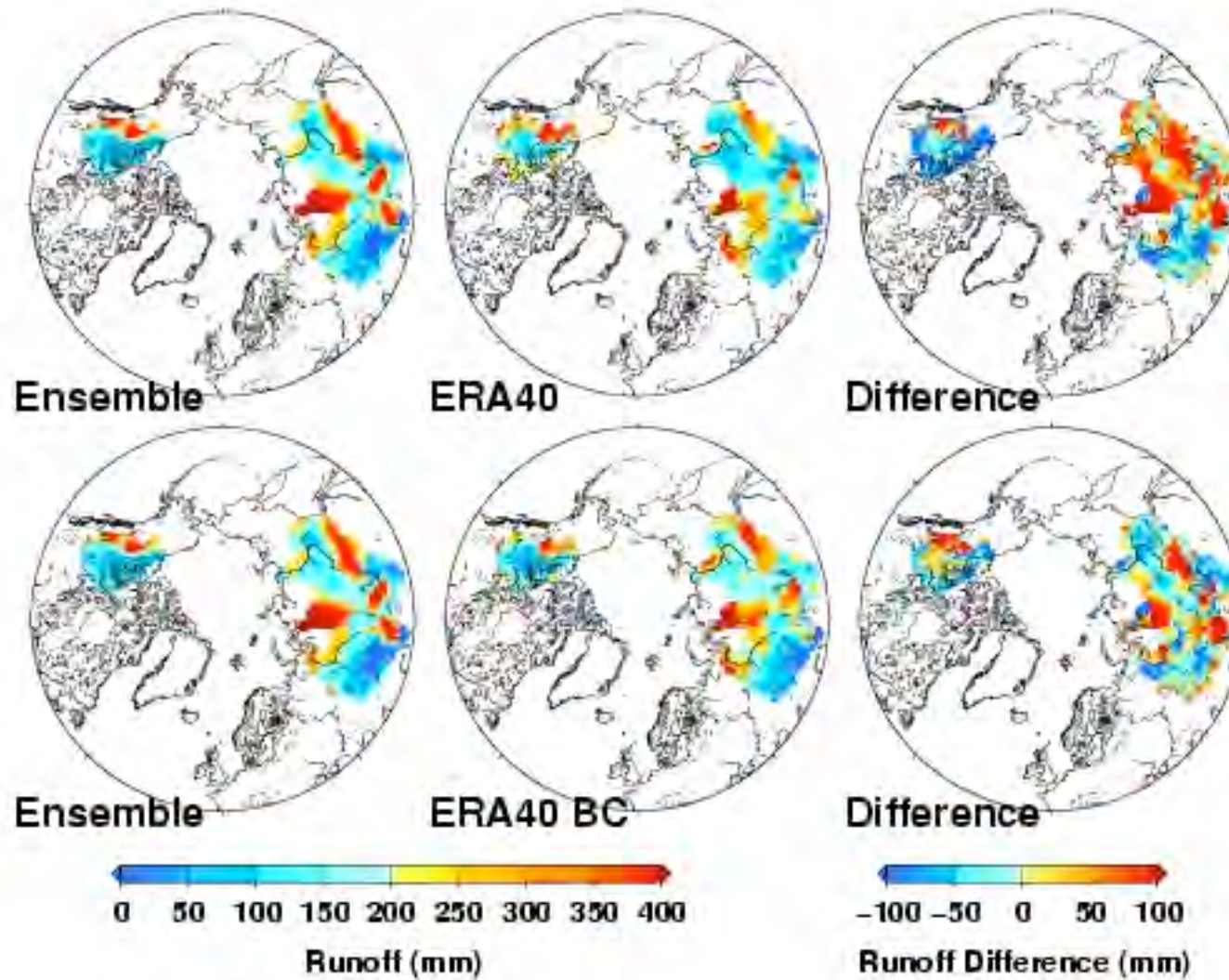
Lakes and Wetlands Evaporation



Torne River basin, headwater catchments

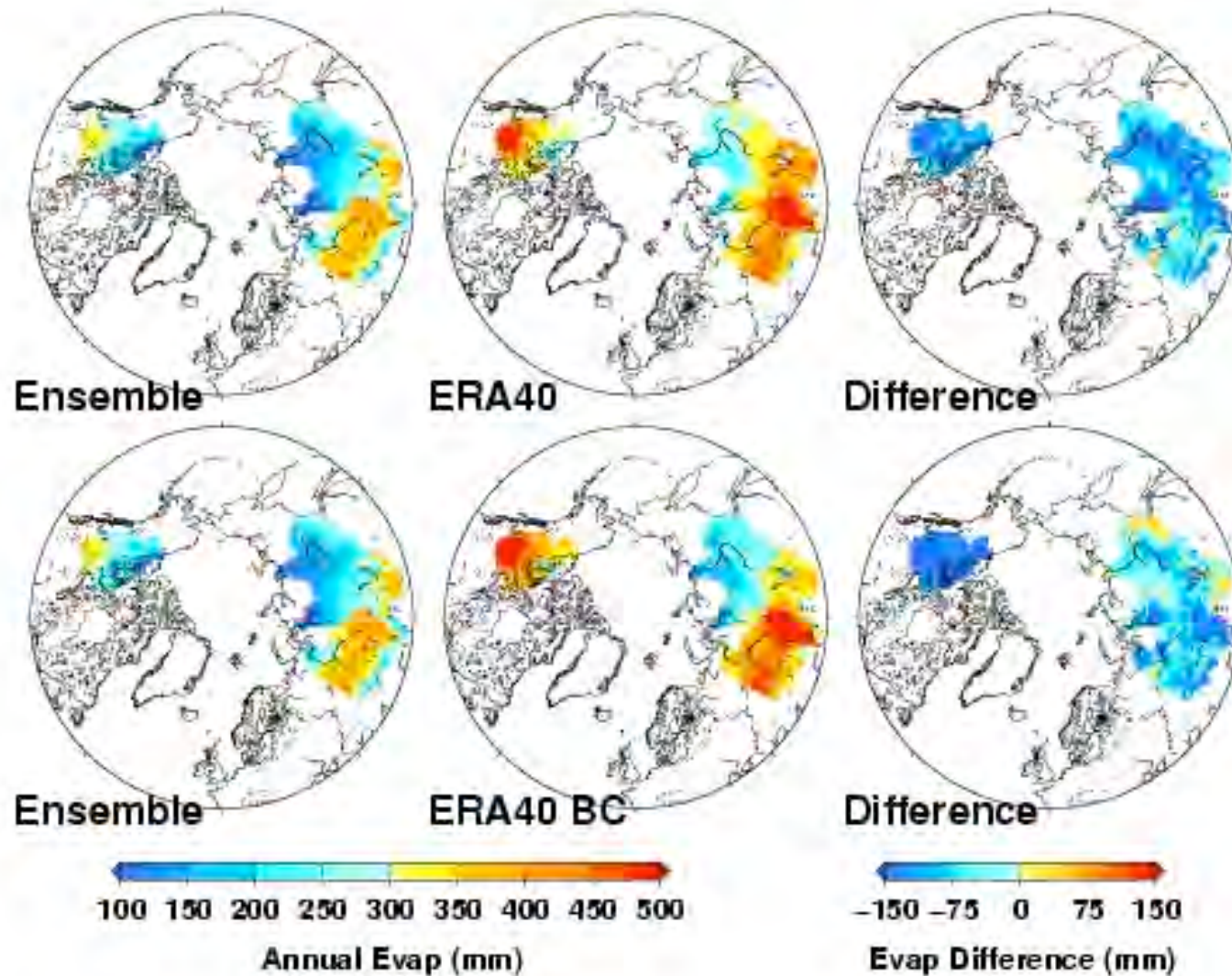


Annual runoff across the pan-Arctic



From Bohn et al. 2005 (Multi-model estimates of Arctic land surface conditions, AGU Dec 2005)

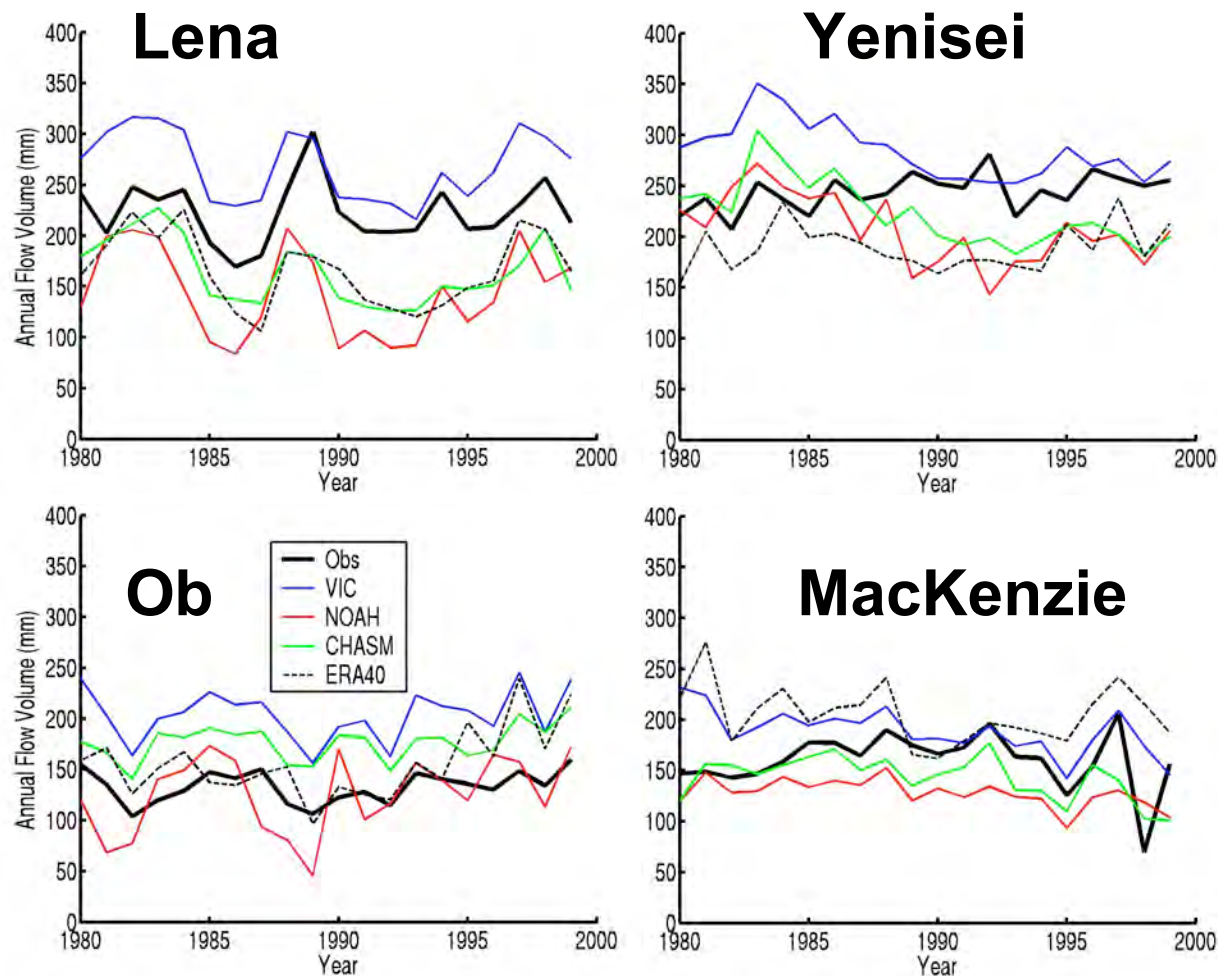
Annual evapotranspiration across the pan-Arctic



From Bohn et al. 2005 (Multi-model estimates of Arctic land surface conditions, AGU Dec 2005)

Multi-model runoff across the pan-Arctic

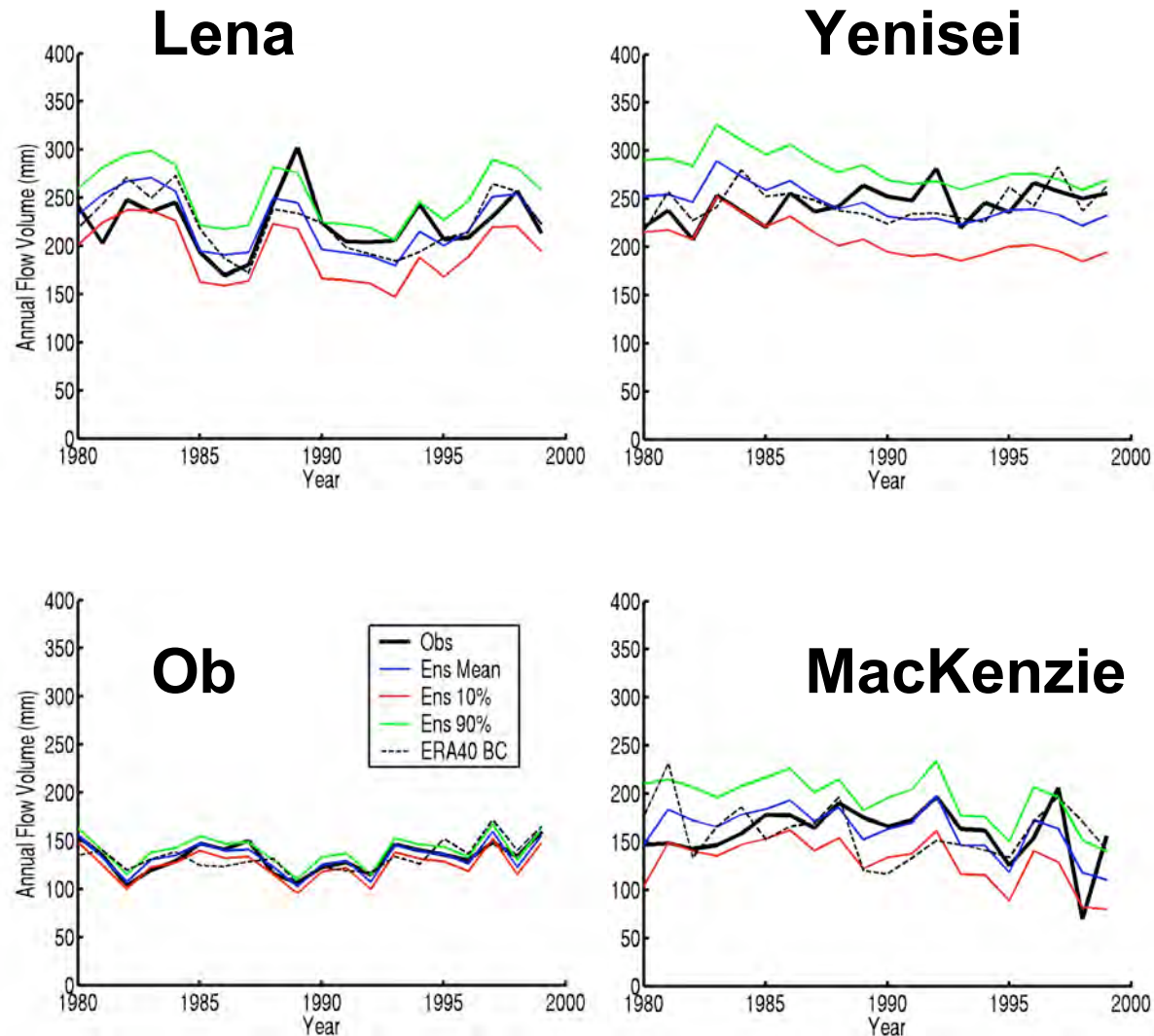
Raw Annual Discharge, 1980-1999



From Bohn et al. 2005 (Multi-model estimates of Arctic land surface conditions, AGU Dec 2005)

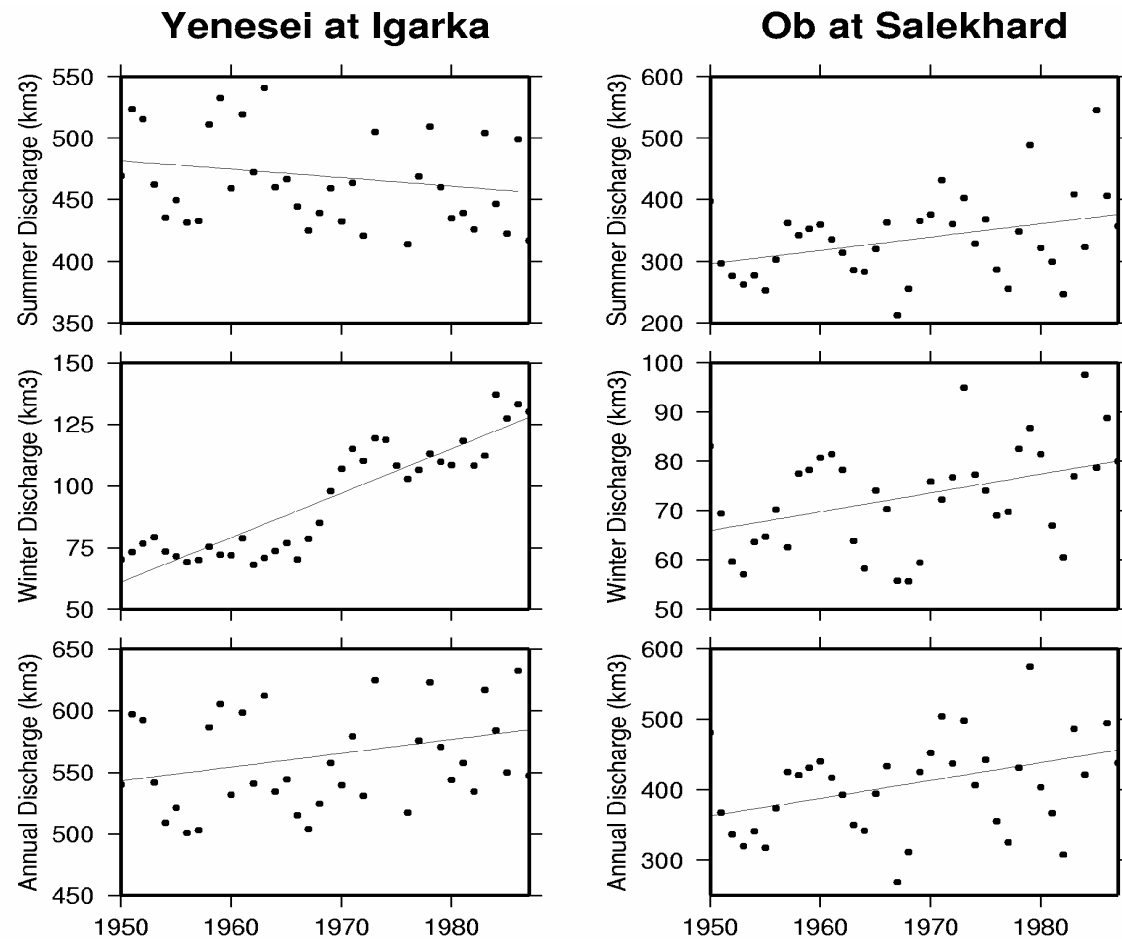
Multi-model runoff across the pan-Arctic

Ensemble Discharge, 1980-1999



From Bohn et al. 2005 (Multi-model estimates of Arctic land surface conditions, AGU Dec 2005)

Time series of discharge of the Yenesei and Ob Rivers over the last 50 years.



Conflicting Explanations for Discharge Trends

Increased northward atmospheric moisture transport	Nijssen et al. (2001), Wu et al. (2005), Arnel (2005)
Human effects (reservoir construction)	Yang et al. (2004), Ye et al. (2003), McClelland et al. (2004)
Release of water from permafrost degradation	Frauenfeld et al. (2004), Zhang et al. (2003), Ye et al. (2003)
Climate-induced changes to the land surface (increased fire frequency)	McClelland et al. (2004), Conrad and Ivanova (1997)
Changes in lake areal extent and storage	Smith et al. (2005)
Change in evapotranspiration	Gedney et al. (2006)
Change in snow accumulation / ablation patterns	Brown (2000), Groisman et al. (1994), Robinson et al. (1990)

Snow (Re)Distribution

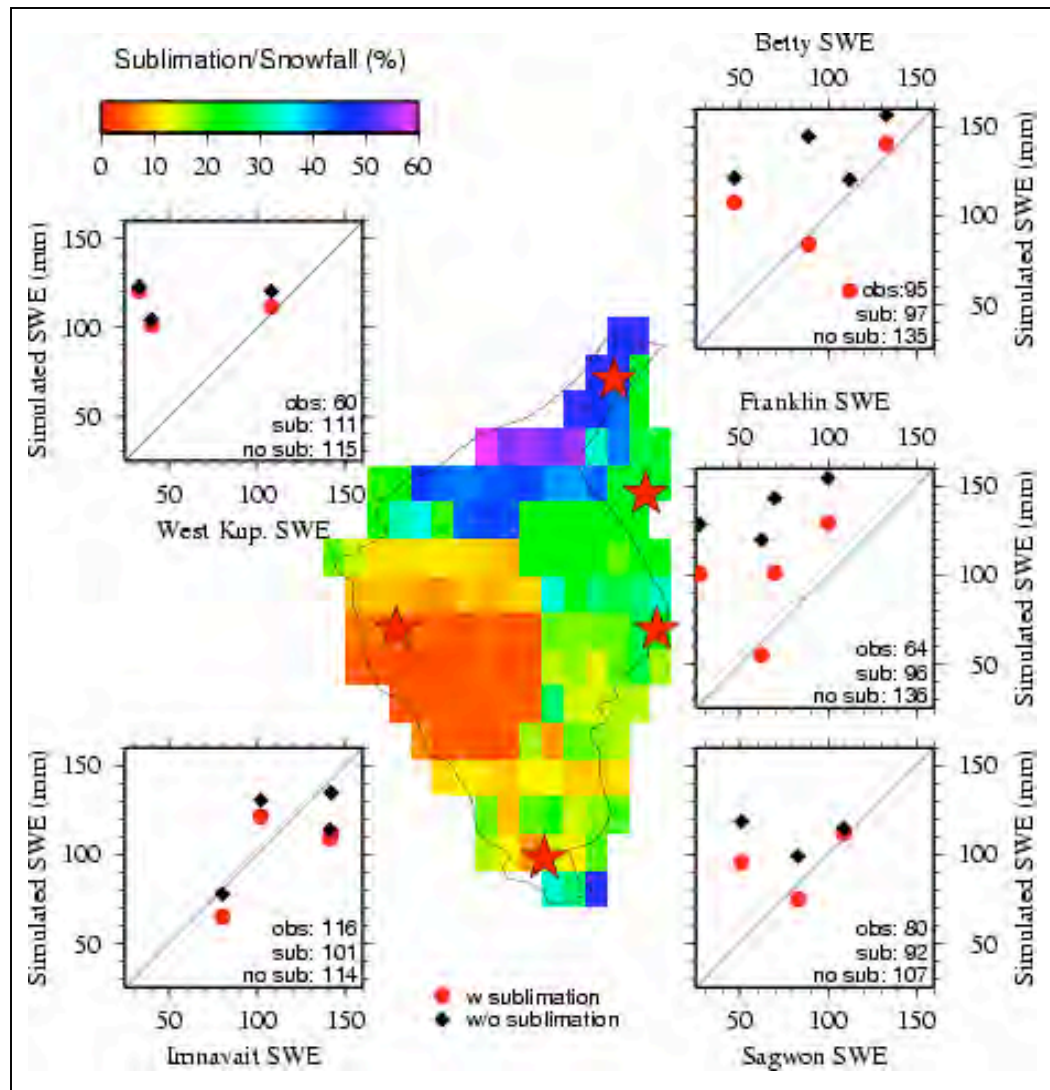


- Sublimation during blowing snow
 - Increase in winter latent heat
 - Decrease in snow available for melt

- Variable snow coverage
 - Variable melt water availability
 - Advection of sensible heat from bare areas



Sublimation from transported snow



What information is available and what is missing for energy and water cycle process modeling efforts, and the ways to address the gaps

“...energy and water cycle process not well understood”

- Permafrost processes (how to model at large scales)
- Surface runoff and baseflow in permafrost environments
- Lake/wetland extent at large scales
- Fire/land cover change effects
- Peatlands – significant in the NEESPI region
- Inertia of the system – some of the processes, such as permafrost dynamics, occur over large time scales that we may or may not be capturing in our modeling

We have made progress on process-based modeling:

- Lake extent and processes
- Blowing snow
- Active layer development
- Water - Carbon cycle interactions

Extent of shallow Lakes and wetlands are highly variable over snow free season

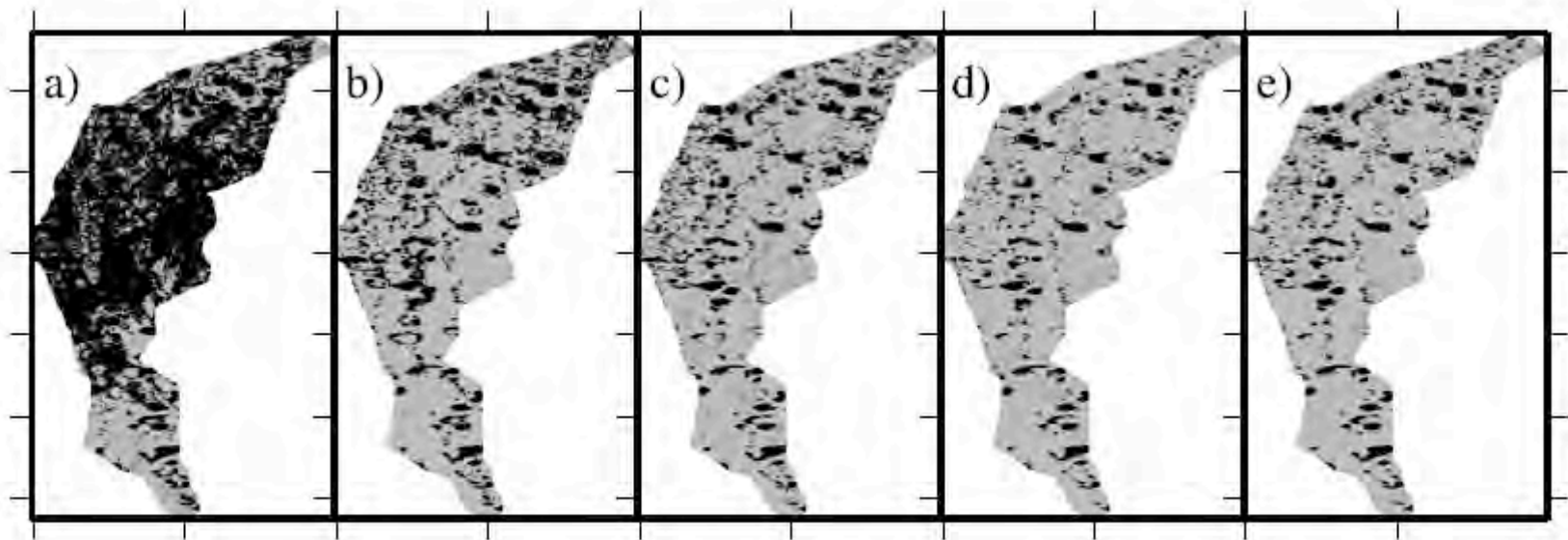
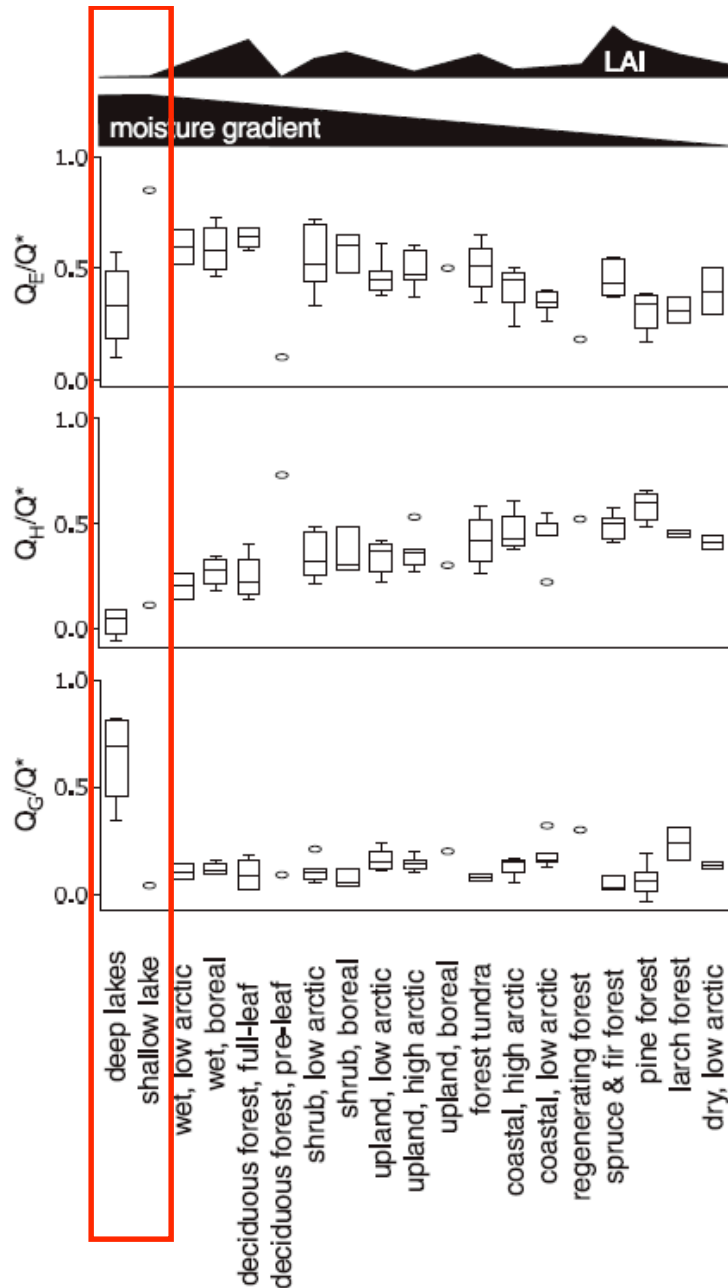


Figure 7. Change in saturated extent (solid areas) of the Putuligayuk River basin from classified SAR images: (a) 14 June 2000, (b) 21 June 2000, (c) 5 July 2000, (d) 22 July 2000, and (e) 7 September 2000.

Bowling, L. C., et al. (2003), The role of surface storage in a low-gradient Arctic watershed, *Water Resources Research*, 39, -.



Eugster, W., et al. (2000), Land-atmosphere energy exchange in Arctic tundra and boreal forest: available data and feedbacks to climate, *Global Change Biology*, 6, 84-115.

Shallow and Deep lakes very different

Evapotranspiration in Kuparuk Basin, AK
Eddy covariance measurements of H₂O flux between 1994-1996 (note: data below may be from different years!)

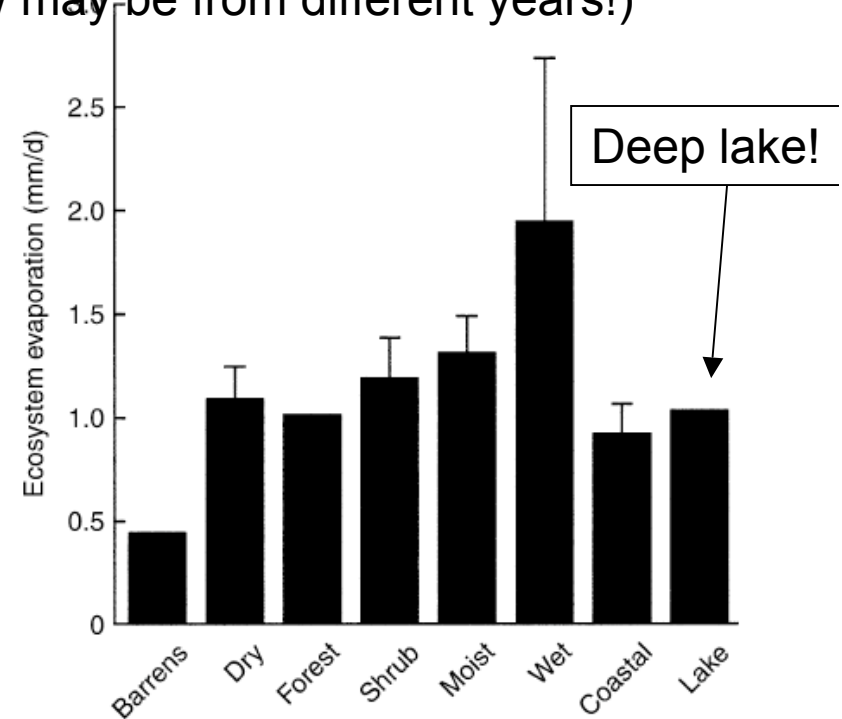
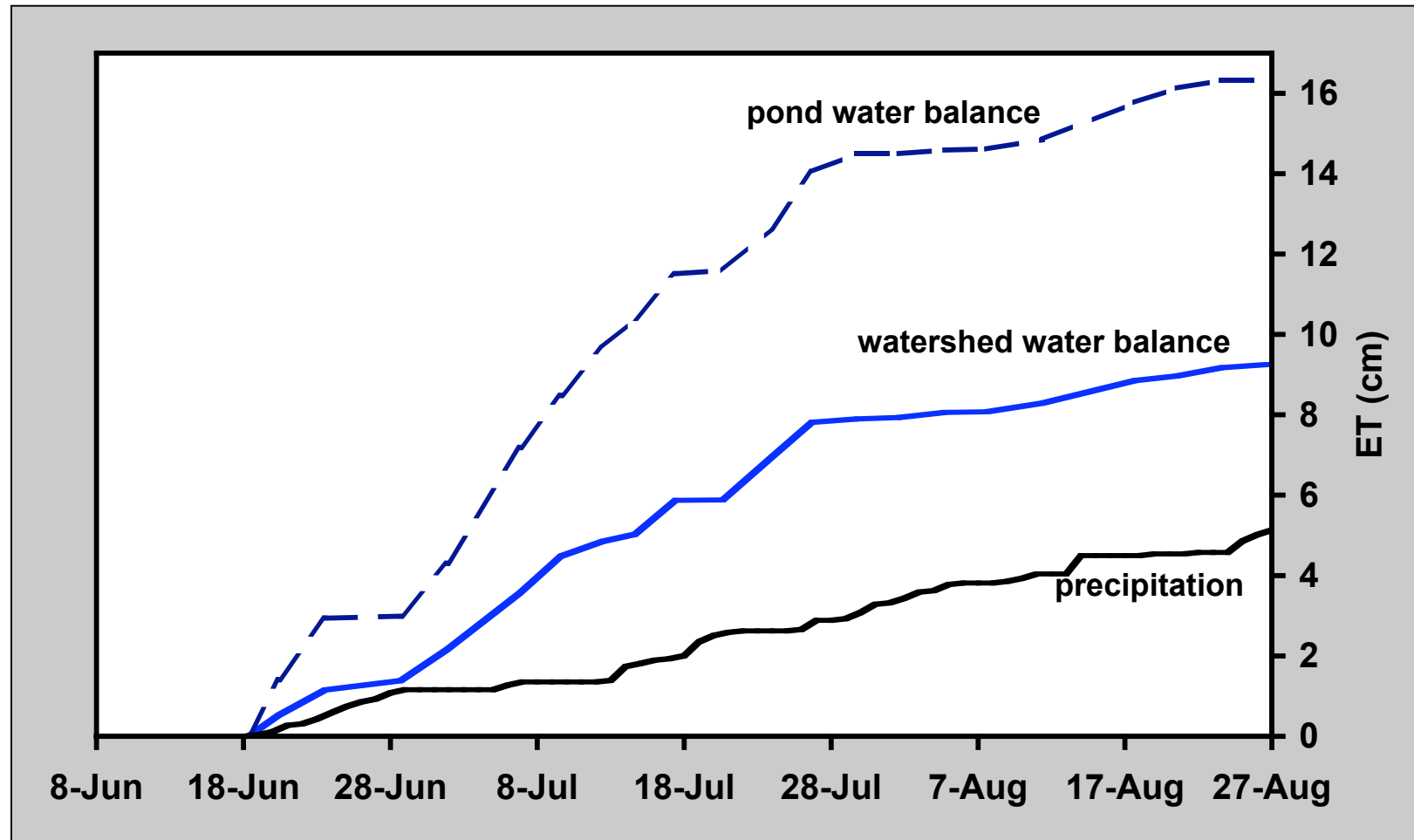


FIG. 7. Total ecosystem evaporation in seven arctic ecosystem types and a lake. Data are means, and error bars represent +1 SE.

McFadden, J. P., et al. (2003), A regional study of the controls on water vapor and CO₂ exchange in arctic tundra, *Ecology*, 84, 2762-2776.

Cumulative wetland complex evapotranspiration, Summer 1996



Lake dynamics

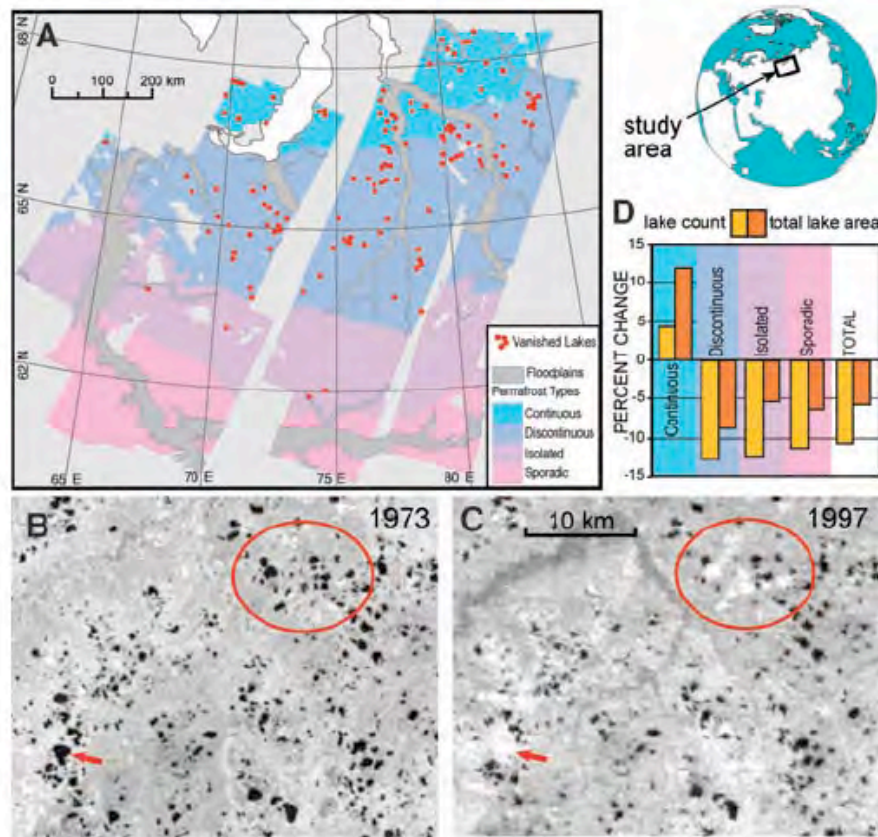
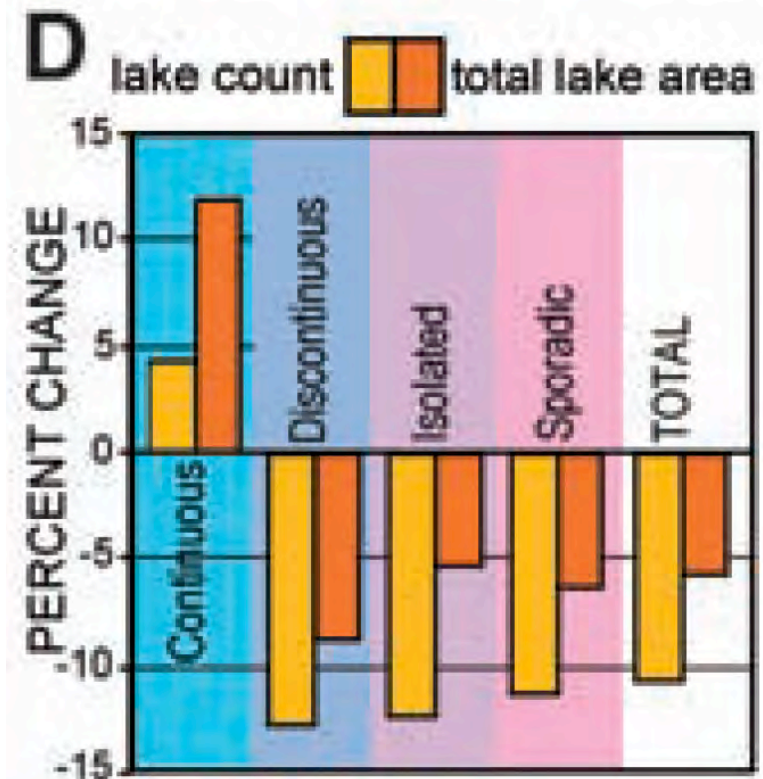
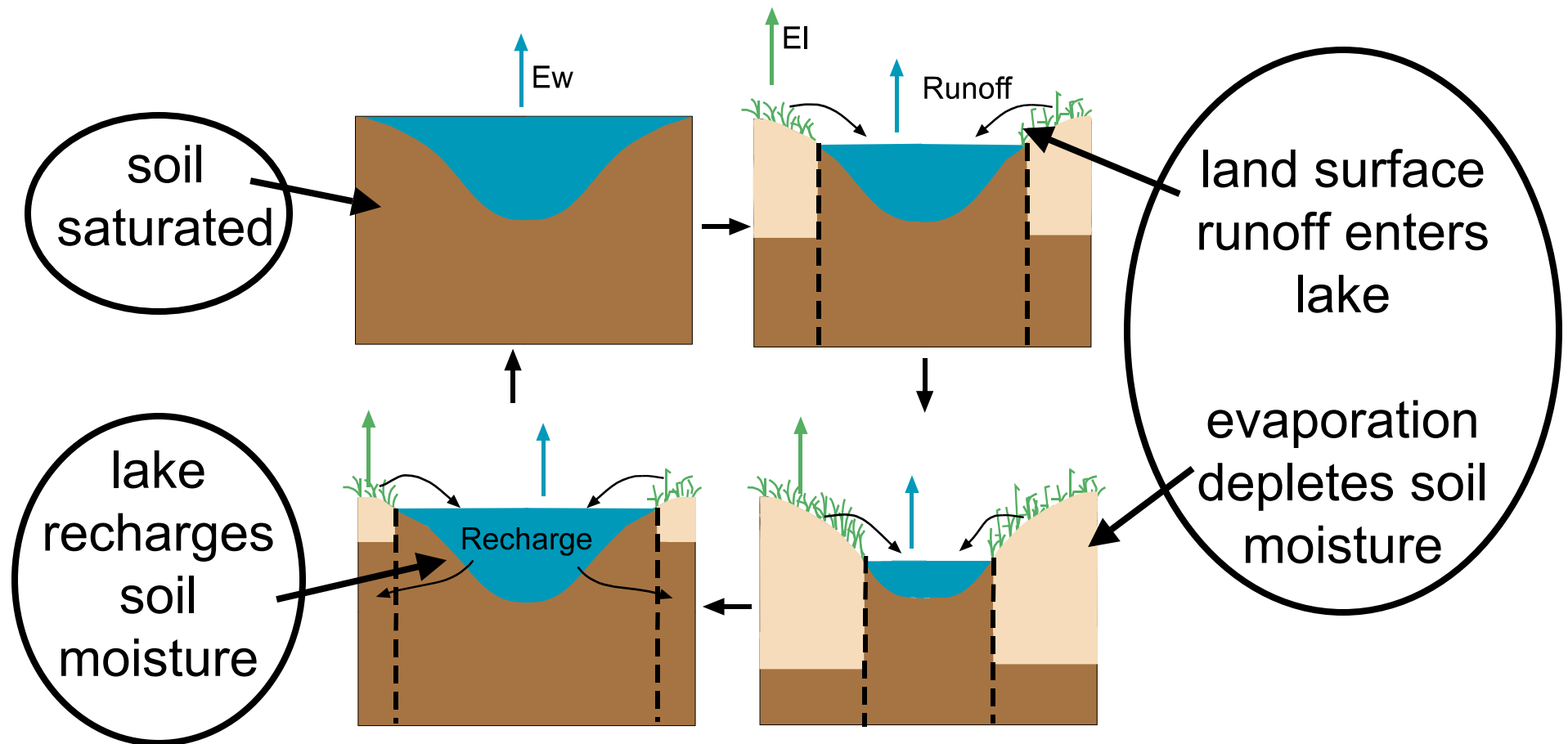


Fig. 1. (A) Locations of Siberian lake inventories, permafrost distribution, and vanished lakes. Total lake abundance and inundation area have declined since 1973 (B), including (C) permanent drainage and revegetation of former lakebeds (the arrow and oval show representative areas). (D) Net increases in lake abundance and area have occurred in continuous permafrost, suggesting an initial but transitory increase in surface ponding.



Smith, L. C., et al. (2005), Disappearing Arctic lakes, *Science*, 308, 1429-1429.

Lakes and wetland modeling (VIC Lake/Wetland Algorithm)

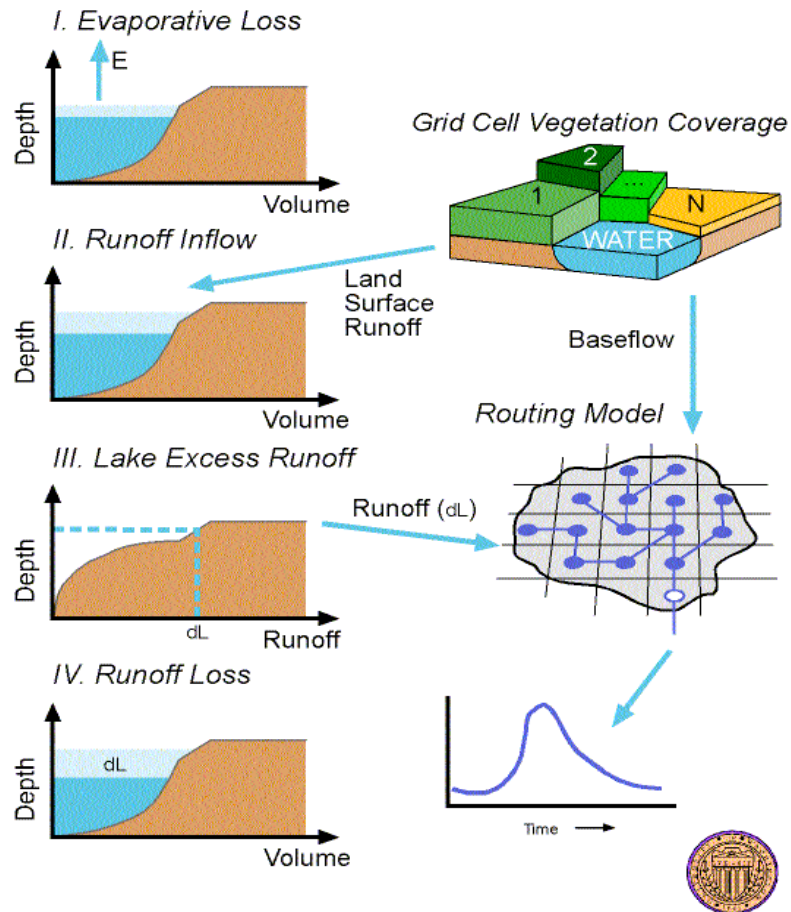


Lake drainage = $f(\text{water depth, calibration parameter})$

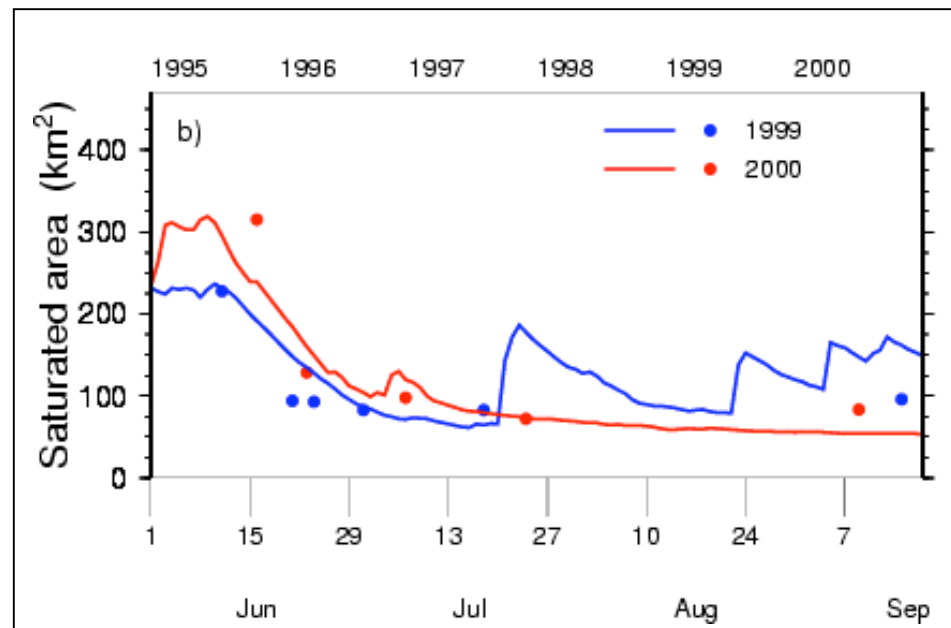
(From T. Bohn, U. Washington)

Seasonal variation in surface water extent

VIC Lake Algorithm



Putuligayuk River, Alaska



Lake and wetland extent

- Problem:
 - Evaporation/carbon flux from shallow water lake and wetland systems is an important component of the energy balance/carbon cycle
 - Water extent is changing due both to ET changes and permafrost thaw, but direction of change depends on the permafrost dynamics
- Needs:
 - Spatial distribution of depth of permafrost
 - Ground ice content
 - Model algorithm development

How does peatland differ?

Model parameters describing the soil. The peat soil values are the median values found for fibric (little decomposed) and sapric (highly decomposed) peat found by Letts et al. (2000) in their literature review of parameter values, except heat capacity and heat conductivity that were calculated according to Williams and Smith (1989) and Campbell and Norman (1998). The values for the third and fourth soil layers represent coarse mineral soil texture and are taken from Table 2 in Cox et al. (1999).

Parameter	Upper soil layer (peat)	Second soil layer (peat)	3 rd and 4 th soil layer (mineral soil)
Layer thickness	0.1 m	0.25 m	0.55 and 2 m, respectively
Saturated hydraulic conductivity	$2.8 \times 10^{-4} \text{ m s}^{-1}$	$1.0 \times 10^{-7} \text{ m s}^{-1}$	$7.57 \times 10^{-6} \text{ m s}^{-1}$
Saturated hydraulic suction	$1.03 \times 10^{-2} \text{ m}$	$1.01 \times 10^{-2} \text{ m}$	$3.29 \times 10^{-2} \text{ m}$
Clapp-Hornberger's B constant	2.7	12	4.5
Soil-moisture content at wilting point	$0.04 \text{ m}^3 \text{ m}^{-3}$	$0.15 \text{ m}^3 \text{ m}^{-3}$	$0.06 \text{ m}^3 \text{ m}^{-3}$
Soil-moisture content at critical point (field capacity)	$0.24 \text{ m}^3 \text{ m}^{-3}$	$0.70 \text{ m}^3 \text{ m}^{-3}$	$0.15 \text{ m}^3 \text{ m}^{-3}$
Soil-moisture content at saturation	$0.90 \text{ m}^3 \text{ m}^{-3}$	$0.80 \text{ m}^3 \text{ m}^{-3}$	$0.40 \text{ m}^3 \text{ m}^{-3}$
Dry-soil heat capacity	$0.25 \times 10^6 \text{ J m}^{-3} \text{ K}^{-1}$	$0.60 \times 10^6 \text{ J m}^{-3} \text{ K}^{-1}$	$1.25 \times 10^6 \text{ J m}^{-3} \text{ K}^{-1}$
Dry-soil heat conductivity	$0.03 \text{ W m}^{-1} \text{ K}^{-1}$	$0.04 \text{ W m}^{-1} \text{ K}^{-1}$	$0.272 \text{ W m}^{-1} \text{ K}^{-1}$

peat

Mineral soil

peat

peat

Mineral soil

Soil-moisture content at saturation

$0.90 \text{ m}^3 \text{ m}^{-3}$

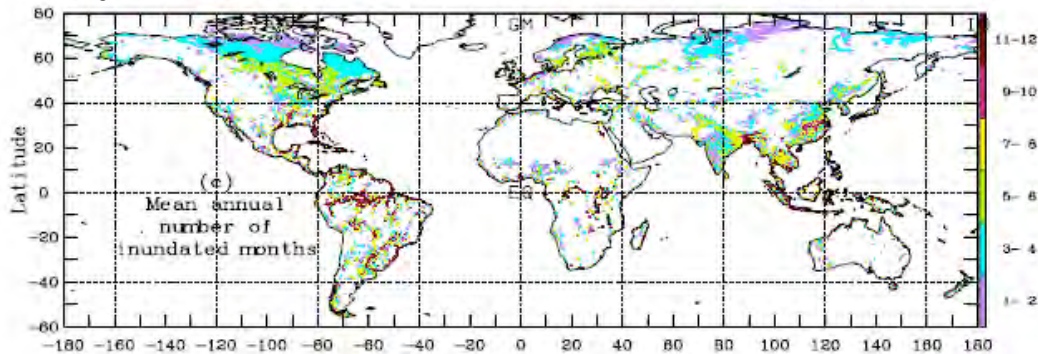
$0.80 \text{ m}^3 \text{ m}^{-3}$

$0.40 \text{ m}^3 \text{ m}^{-3}$

Rennermalm, A. K., et al. (2005), Interannual variability in carbon dioxide exchange from a high arctic fen estimated by measurements and modeling, *Arctic Antarctic and Alpine Research*, 37, 545-556.

How does peatlands alter the water and energy balance?

Wetland/peatlands in Northern Hemisphere Inundated 1-4 months a year



Prigent, C., et al. (2007), Global inundation dynamics inferred from multiple satellite observations, 1993-2000, *Journal of Geophysical Research-Atmospheres*, 112, -.

Wetland/peatlands:
- Higher evapotranspiration

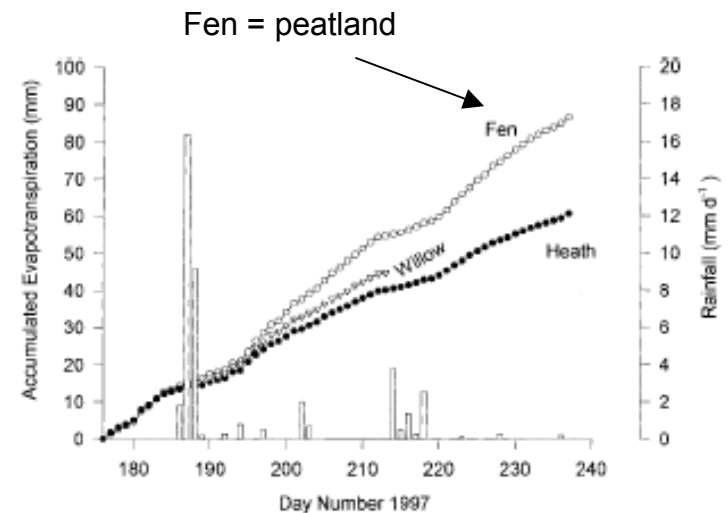


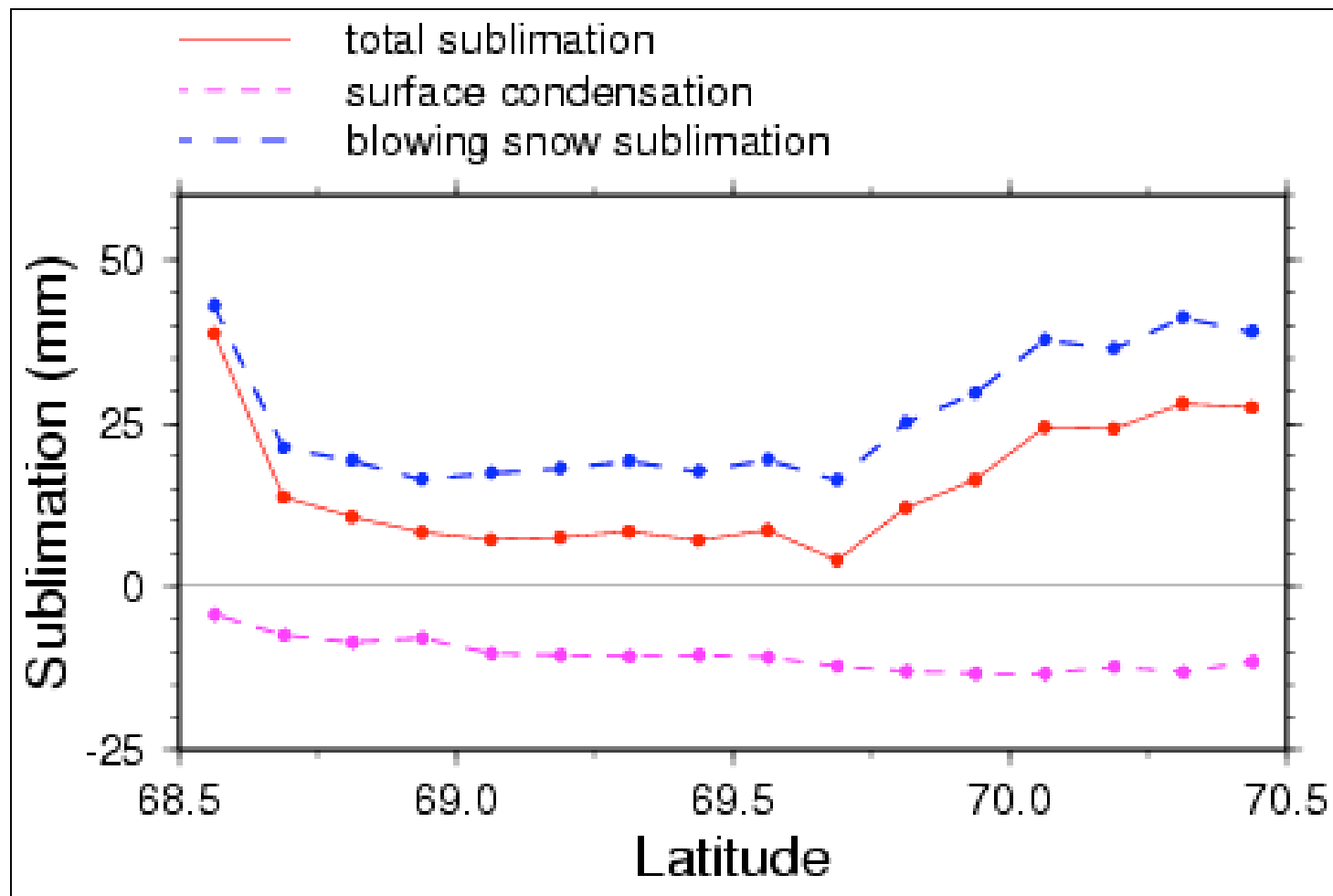
Fig. 4. Daily rainfall and accumulated evapotranspiration from three ecosystems throughout the 1997 growing season

Soegaard, H., et al. (2001), Surface energy- and water balance in a high-arctic environment in NE Greenland, *Theoretical and Applied Climatology*, 70, 35-51.

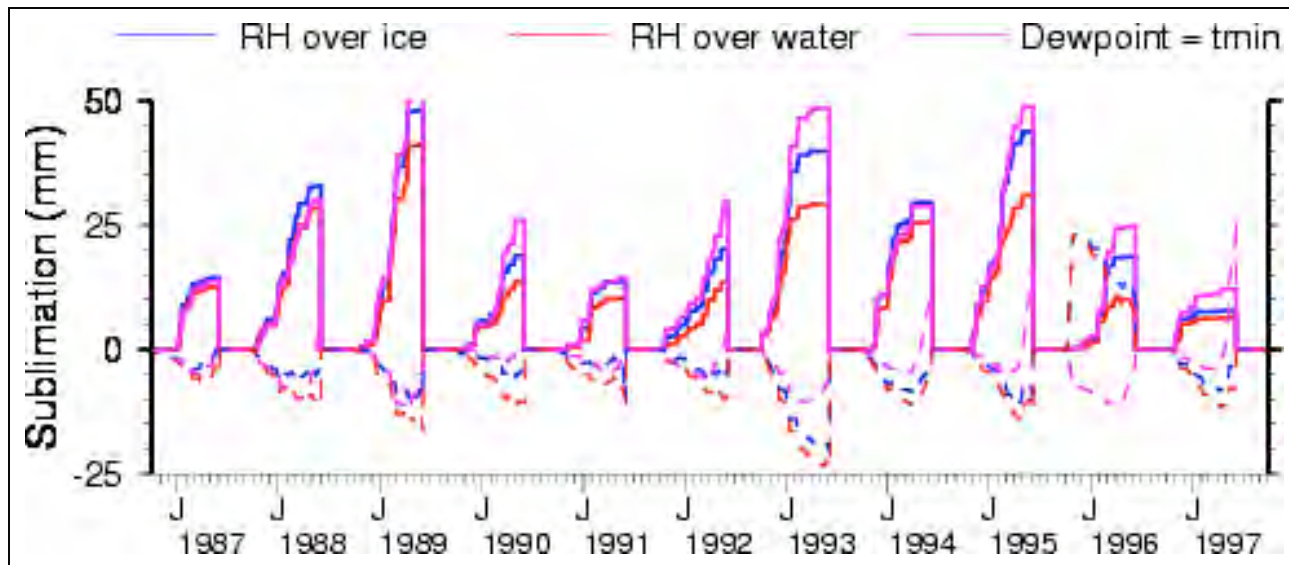
Vapor pressure uncertainty

- Problem:
 - Extremely low winter temperatures in the arctic lead to low saturated vapor pressure values
 - Existing parameterizations (e.g. $T_{\text{dew}} = \text{daily } T_{\text{min}}$) are not adequately tested in the arctic
 - In-situ observations are rare and uncertainty often exists whether the observations are taken assuming equilibrium with water or ice
- Implications:
 - Evaporation/sublimation may be limited by vapor pressure rather than water or energy
 - The direction of net over-winter vapor flux is unknown
- Needs:
 - Over-winter eddy covariance and humidity measurements?

Annual average sublimation Kuparuk River



Sensitivity to vapor pressure

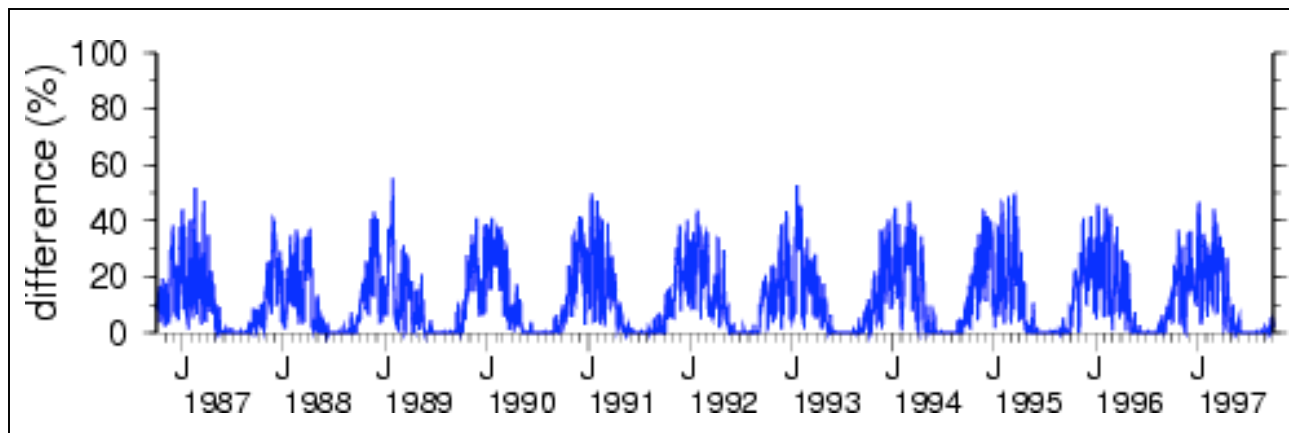


Average annual
surface sublimation:

RH w.r.t. ice
-5.3 mm

RH w.r.t. water
-9.2 mm

Tdew = Tmin
+6.6 mm



“What’s missing..”

New Directions and Research Needs

1) Data -

Primary need is to assemble a best quality retrospective precipitation data set, duration at least one decade, with best current attempts at gauge catch correction.

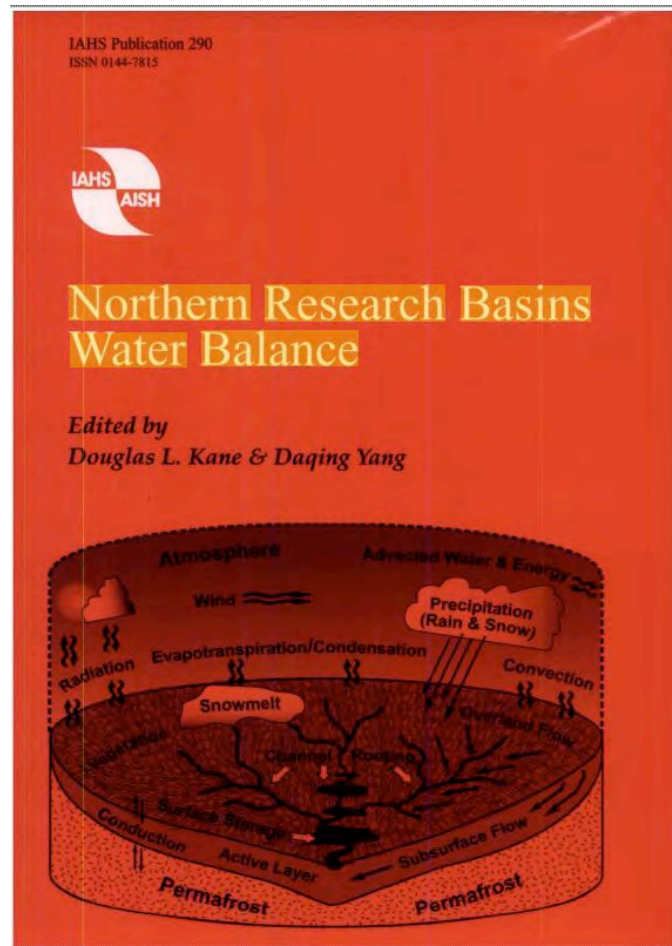
What is the potential of remote sensing to supplement the sparse in situ observation network?

2) Modeling

- We tend to focus on an entire basin. Process and validation tends to be at ‘experimental basin scale. Issues of process scaling
- Continue improving and testing process representations
 - Energy balance - - snow, ET
 - Subgrid snow distribution, partial coverage, sublimation from blowing snow
 - Frozen soils and permafrost
 - Lake and wetland storage, and linkages to the carbon cycle.
- Incorporate data assimilation approaches in off-line modeling?

“What’s missing..”

New Directions and Research Needs



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Thank you